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Dear Abd. Rasyid Jalil:

Thank you for submitting the manuscript, "CENDRAWASIH HOT POOL: STUDY OF THE HIGH-FREQUENCY OCCURRENCE OF HIGH SST PHENOMENA AT CENDRAWASIH BAY, PAPUA" (ID 2652) to GEOGRAPHY, ENVIRONMENT, SUSTAINABILITY. With the online journal management system that we are using, you will be able to track its progress through the editorial process by logging in to the journal web site:

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CENDRAWASIH HOT POOL: STUDY OF THE HIGH-FREQUENCY OCCURRENCE OF HIGH SST PHENOMENA AT CENDRAWASIH BAY, PAPUA

ABSTRACT. The warm pool has been used to define a water body with the characteristic of SST exceeding 28°C within a particular area and a relatively long period in an annual circle. However, there are regions with an annual mean SST measured above 30°C, and we classified them as Hot Pools because of the conditions of intense solar radiation and low wind speed. One of the Hot Pool spots was found in Indonesia, in Cenderawasih Bay. In the present study, the existence of the Cendrawasih Hot Pool is examined using long-term observation of satellite SST data. We also used surface wind data, surface heat flux, and surface current to investigate their mechanisms. The results show that for more than half a year, SSTs in Cendrawasih Bay can reach more than 30°C. The location is also dominated by low wind speed, i.e., 80% wind speed of fewer than four m/s occurred during 2003-2015, and these caused the low latent loss in Cendrawasih Bay. In surface current, during the dry and wet seasons, Cendrawasih Bay is fully isolated since the easterly subsurface water flow does not enter the bay. The absence of strong currents prevents the mixing process from maintaining the high temperature in the surface layer. Those processes are figured out and become strong evidence to justify Cendrawasih Bay as one of the Hot Pool areas within the Indonesian seas.

KEYWORDS: Global climate; Sea Surface Temperature; Hot Pool Spot; Cendrawasih Bay.

INTRODUCTION

The western equatorial Pacific significantly influences the global climate. The warm pool, a region with average sea surface temperatures (SST) above 28°C, has a specific

impact on the Earth's circulation (e.g., Wyrski 1989; Yan et al. 1992; Clement and Seager 1999; Chongyin et al. 1999; Pierrehumbert 2000; Clement et al. 2005; Thoron et al. 2005; Herweijer et al. 2005).

On the other hand, high sea surface temperature (SST) in tropical regions has attracted researchers to investigate the mechanisms since the formation of high SST requires a particular atmospheric process (e.g., Ramanathan and Collins 1991; Wallace 1992; Arking and Ziskin 1994). This process is depicted by Waliser and Graham (1993), which shows the relation between SSTs and deep convection. They utilized highly reflective cloud information from an arbitrary examination of monthly SST data with a grid spacing of 2° (produced from combined satellite observation and in situ data) and daily visible and infrared satellite image data. The highly reflective cloud increases along with the SST as it rises from 26°C to 29.5°C . In contrast, the highly reflective cloud diminishes with increased SST in the temperature from 29.5°C to 32°C . As a result, the analysis proved that several atmospheric processes impacted SSTs below and above 29.5°C .

By taking advantage of high temporal and spatial resolution sea surface temperature (SST) products derived from satellite observations (i.e., daily and $\leq 25 \text{ km} \times 25 \text{ km}$), several studies (Kawamura et al. 2008; Qin et al. 2007, 2008; Qin and Kawamura 2009, 2010; Wirasatriya et al. 2015, 2016, 2017a, 2020) were able to identify high SST events (i.e., more than 30°C) in specific areas and at certain periods and define them as Hot Event (HE). In summary, they concluded that considerable daily heat gains characterize the production of HE under high solar radiation and low wind speed brought on by "remote convection" mechanisms.

Wirasatriya et al. (2015) elaborated on the climatology of HEs in the western equatorial Pacific using the SST dataset derived from satellite microwave sensors. Throughout nine years of observation (2003-2011), they discovered 71 HE cases in the western equatorial Pacific, with the majority centered on the Solomon Islands and New Guinea Island's northern coasts, which extend eastward up to 160°W . According to the climatology, the region is located in an area with solar radiation of more than 200 W/m^2 and wind speeds of less than 4 m/s . Low wind speeds heavily influence the mechanism for HE incidence in the western equatorial Pacific. Much of the equatorial region experiences sun radiation above 200 W/m^2 during the HE periods. Low wind speeds minimize latent heat loss, which results in high SSTs and HEs in specific locations. Wirasatriya et al. (2015) also emphasized that high solar radiation and low wind speed are much more common during the development stage and less common during the decay stage. This study also demonstrated that a rise in the long-term mean SST in the western equatorial Pacific is correlated with an increase in the frequency of HE events. HEs were responsible for 51.5% of the SSTs $>30^\circ\text{C}$ in the warm pool region bounded by the 29.5°C isotherms of the climatological SST. Thus, statistically, there is a relation between the occurrence of HE and the formation of the western Pacific warm pool. Moreover, Wirasatriya et al. (2020) demonstrated the role of HE in maintaining the warm mixed layer in the western Pacific warm pool. The frequent occurrence of HE transports heat from the surface layer to the deeper layer above the thermocline.

Within the Indonesian seas, the frequent appearance of high SST $> 30^\circ\text{C}$ has been reported by Tita et al. (2020) in Tomini Bay and by Swandiko et al. (2021) in the Malacca Strait. Their appearances also the high solar radiation and weak wind. The morphology of the semi-enclosed waters causes the high-frequency occurrence of low wind speed by less than 4 m/s in both areas.

In the present study, we demonstrate the constant high SST occurrence (more than 30°C) in Cendrawasih Bay. Cendrawasih Bay is located in northern Papua Island, part of the western Pacific warm pool. It is semi-enclosed water with a deep basin, surrounded by

mountain chains (Fig. 1). Cendrawasih Bay is the habitat for whale sharks (Ihsan et al. 2018). Since the area determines warm pool definition with SST climatology higher than 28°C, we call the high-frequency occurrence of SSTs higher than 30°C in the Cendrawasih Bay the Cendrawasih Hot Pool. Wirasatriya et al. (2015) found that the duration of HE occurrence in the western equatorial Pacific is no longer than two months since HE requires a typical condition of high solar radiation and low wind speed. Thus, it is interesting to understand the mechanisms of the high-frequency occurrence of high SST in the Cendrawasih Hot Pool.

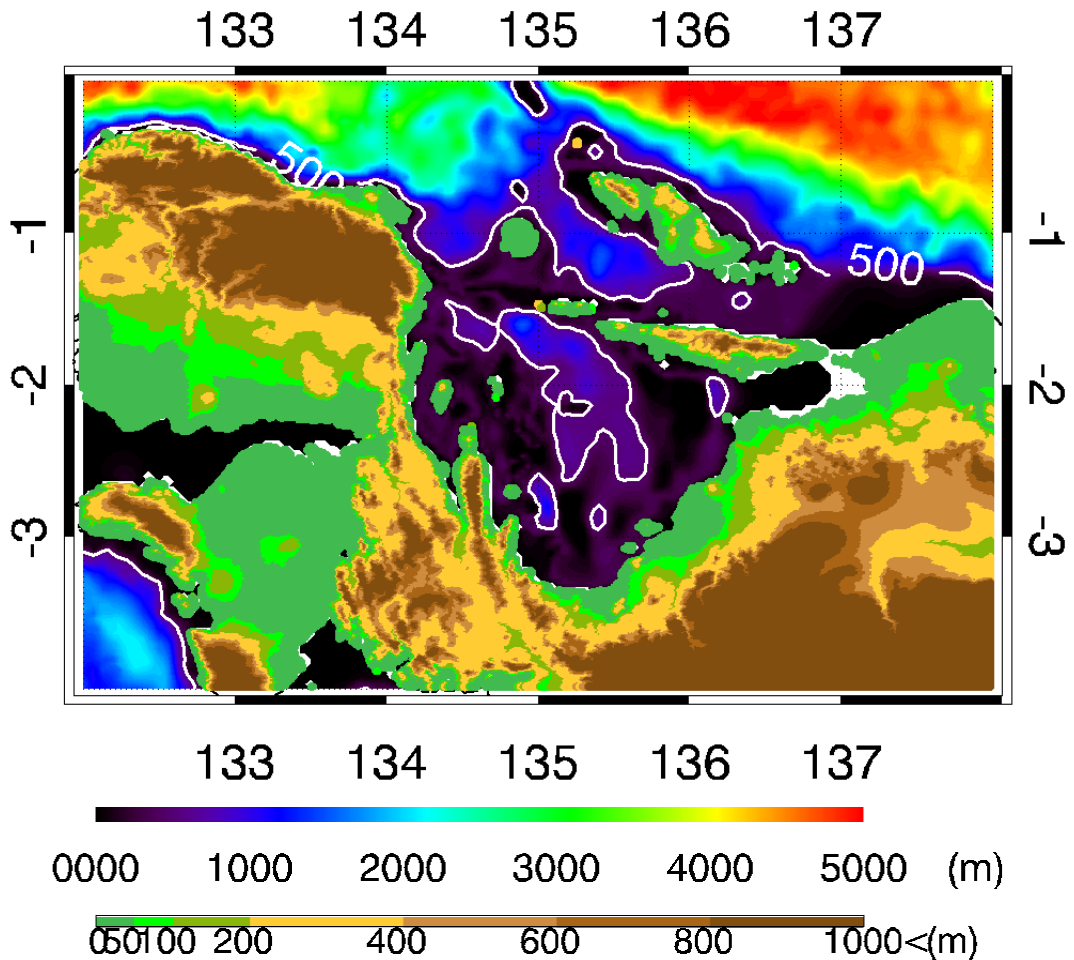


Fig. 1. Bathymetry and topography of Cendrawasih Bay

MATERIALS AND METHODS

We used the semi-daily 11 μ m SST products from Moderate Resolution Imaging Spectroradiometer (MODIS) Aqua Lv3 with a spatial resolution of 0.04°×0.04° (Esaias et al. 1998) and observation period from 2003 to 2015. The MODIS SST 11 μ m Multi-Channel SST algorithm generates SST by using brightness temperatures at 11 μ m and 12 μ m (Brown and Minnet 2009). The best accuracy was achieved by validating and testing this data against in-situ measurements (e.g., Lee et al. 2010; Qin et al. 2014; Ghanea et al. 2015).

We can obtain the Cross-Calibrated Multi-Platform (CCMP) gridded surface vector winds version 2.0 from the following website: www.remss.com, and use it to calculate surface winds. As a result, CCMP is regarded as a Level-3 ocean vector wind analysis product. They are created utilizing data from moored buoys, satellites, and model winds.

The surface wind data for spatial and temporal resolutions is $0.25^\circ \times 0.25^\circ$ respectively and quarter daily. This wind product has greater accuracy than other wind reanalysis data.

We used a 0.5° grid from the Objectively Analyzed Air-Sea Fluxes project to analyze daily latent heat fluxes and solar radiation (Yu and Weller, 2007). The period of observation for latent heat flux is from 2003 to 2015, but for solar radiation is only from 2003 to 2009 due to the availability of the data. Global 30 Arc-Second Elevation data was used to obtain topography. Global 30 Arc-Second Elevation (GTOPO30) data was used to obtain topography (https://www.usgs.gov/centers/eros/science/usgs-eros-archive-digital-elevation-global-30-arc-second-elevation-gtopo30?qt-science_center_objects=0#qt-science_center_objects). ETOPO, a one-arc-minute global relief model of Earth's surface that combines land topography and ocean bathymetry, provides bathymetry (<https://www.ngdc.noaa.gov/mgg/global/>). The analysis of the current study is based on a climatological feature. To create monthly and monthly climatology, all geophysical parameters are incorporated as follows: (Wirasatriya et al. 2017b)

$$\bar{X}(x, y) = \frac{1}{n} \sum_{i=1}^n x_i(x, y, t) \quad (1)$$

where $\bar{X}(x, y)$ is monthly mean value or monthly climatology value at position (x, y) , $x_i(x, y, t)$ is i^{th} value of the data at (x, y) position and time t . Next, n is the number of data in 1 month and the number of monthly data in 1 period of climatology (i.e., from 2003 to 2016 = 13 data) for monthly calculation and monthly climatology calculation, respectively. Moreover, if pixel x_i is a hollow pixel, it is not included in the calculation. To look at the relationships between the parameters, we also used correlation analysis

For oceanic parameters, we used the current reanalysis data from <http://marine.copernicus.eu/> to obtain the seasonal variation of the current pattern in the Cendrawasih Hot Pool.

RESULTS AND DISCUSSION

Cendrawasih Hot Pool

The definition of the Cendrawasih Hot Pool is described in Fig. 2, which shows the climatological mean of SST and the percentage of daily SSTs more than 30°C from 2003 to 2015 (13 years). In Cendrawasih Bay, the mean SST at the southern part of 2°S is more than 30°C , and the high SSTs are more than 50%. It means that the mean SST in Cendrawasih Bay is higher than the definition of the western Pacific warm pool which is 28°C (Wyrtki 1989), and more than half-year SSTs in Cendrawasih Bay can reach more than 30°C . At the southernmost of the bay, the mean SST reaches 30.5°C , with the percentage of high SSTs being more than 70%. It has become the hottest area of Cendrawasih Bay.

In contrast, in the area outside Cendrawasih Bay, i.e., the western Pacific warm pool's offshore seas, the percentage is less than 30%. Thus, this evidence supports the definition of Cendrawasih Bay as the Cendrawasih Hot Pool. The high-frequency occurrence of high SST in Cendrawasih Bay may correspond to whale sharks' appearance throughout the year. Ihsan et al. (2018) indicated the tolerance of whale sharks in high SST conditions.

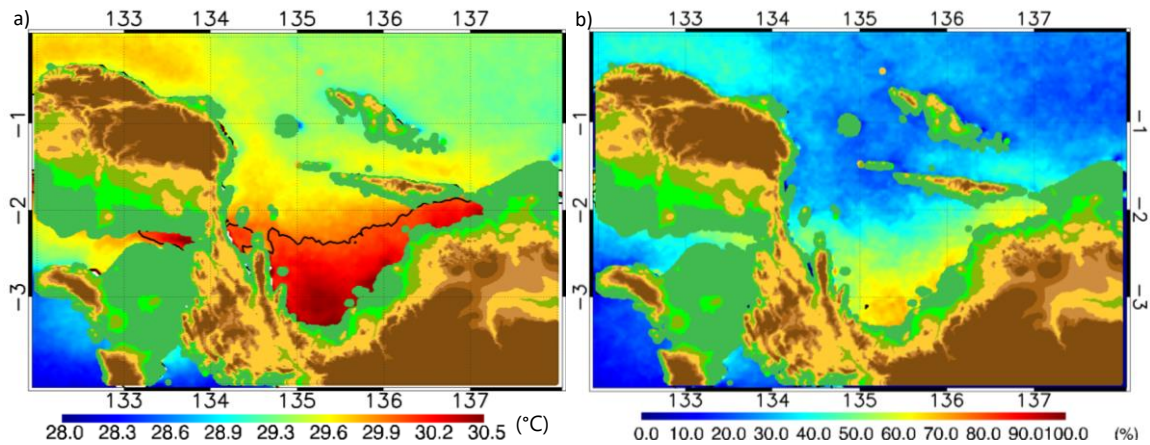


Fig. 2. Climatological mean of SST (a) and Percentage of daily SSTs > 30°C during thirteen years of observation (2003-2015) (b). The black contour is 30°C.

Atmospheric aspect of Cendrawasih Hot Pool

Since high SST of more than 30°C occurs during the condition of low wind speed (i.e., < 4 m/s) and high solar radiation (i.e., > 200 W/m²) (Wirasatriya et al. 2015), we calculated the percentage of low wind speed and high solar radiation in the Cendrawasih Bay. Furthermore, we also calculated the percentage of low latent heat release to explain how wind speed influences the variability of SST. The result is depicted in Fig. 3, which clearly shows that Cendrawasih Bay is dominated by the condition of low wind speed, i.e., 80% wind speed of fewer than four m/s occurred during 2003-2015. This causes low latent loss in Cendrawasih Bay. For solar radiation, the percentage of high solar radiation inside and outside Cendrawasih Bay is almost similar i.e., nearly 50%, indicating that high solar radiation is not dominant in both areas. This means that, although high solar radiation does not frequently occur, the absence of high wind speed may maintain the latent heat loss to keep the persistence of high SST in Cendrawasih Bay.

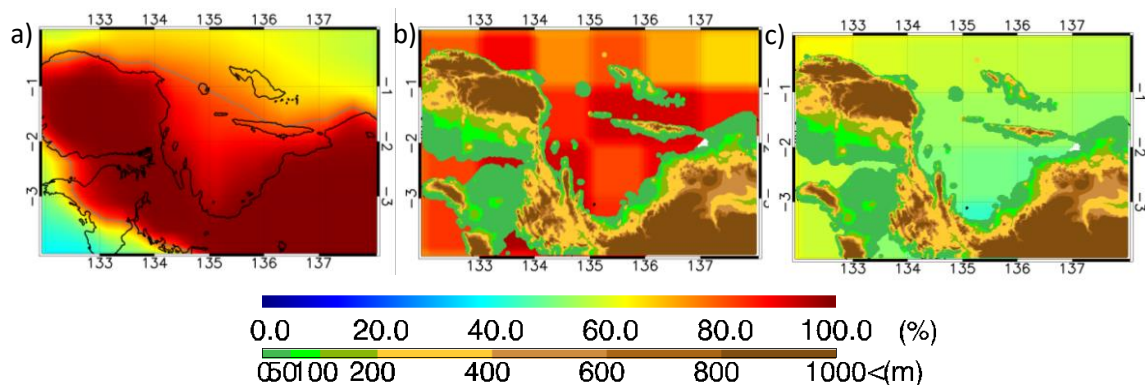


Fig. 3. The percentage of wind speed less than four m/s (a), latent heat release less than 120 W/m² (b), and solar radiation more than 200 W/m² (c).

To investigate how low wind speed occurs in the study area, we show the monthly climatology map of surface wind during summer and winter (Fig. 4). For both seasons, Cendrawasih Bay is protected from high wind speed due to high mountain chains in the western, southern, and eastern parts. In the northern part, small islands at the mouth of the bay prevent the strong wind speed from entering the bay. Thus, the role of topography is crucial for the occurrence of constant high SSTs in Cendrawasih Bay. The same tendencies

are also found at the Tomini Bay and Malacca Strait, as indicated by Tita et al. (2020) and Swandiko et al. (2021).

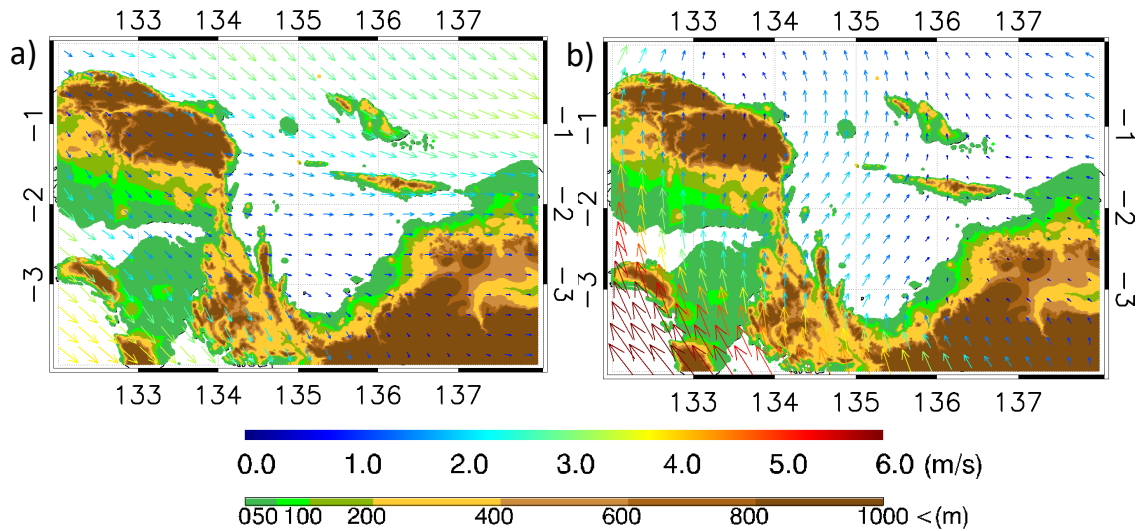


Fig. 4. Monthly climatology of surface wind speed during a) wet season (January) and b) dry season (August)

The oceanic aspect of Cendrawasih Hot Pool

To investigate the oceanic aspect of the Cendrawasih Hot Pool, we plotted the current patterns shown in Fig. 5. A northwestward current has been identified north of Cendrawasih Bay. During the wet season (January), the speed of this current ranges from 0.5 m/s to 0.7 m/s. During the dry season (August), when the northwesterly wind disappears (Fig. 4b), the speed of the northwestward current increases. This northwestward current is known as the New Guinea Coastal Under Current (NGCUC). NGCUC is a permanent current feature at 200 m depth regardless of the wind reversals (Tsuchiya et al. 1989). Inside Cendrawasih Bay, strong currents are absent in this area which prevents the mixing process from maintaining the high temperature in the surface layer. During the dry and wet seasons, Cendrawasih Bay is fully isolated since the easterly subsurface water flow does not enter Cendrawasih Bay. Thus, this isolated basin causes the Cendrawasih Bay area to be influenced mainly by the air-sea interaction process mentioned in the previous section creating Cendrawasih Hot Pool.

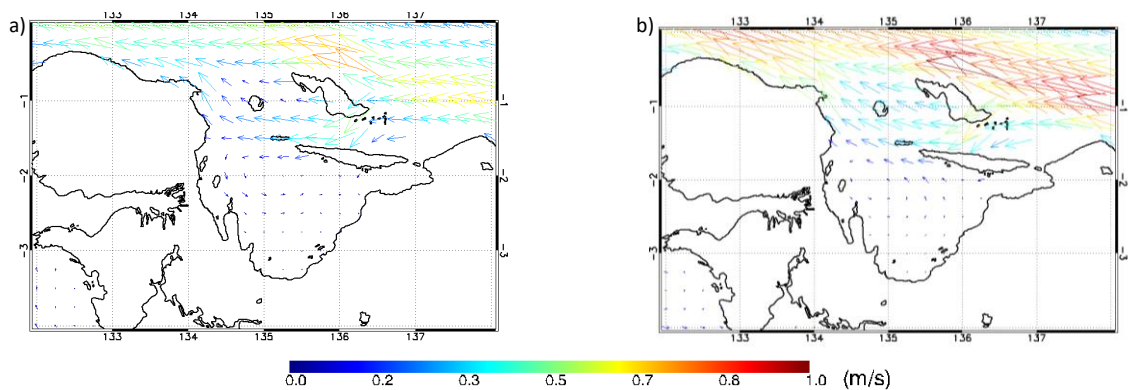


Fig. 5. Monthly climatology of current 100 m in January (a) and August (b)

CONCLUSIONS

The hot Pool is used to categorize high SST episodes ($>30^{\circ}\text{C}$) in particular areas and during specific periods (relatively long periods). Since the "warm pool" has been used to define an area with the annual average SST distributed below 30°C in dominant, therefore "hot pool" was taken to represent the area with SST dominated above 30°C .

Under high solar radiation and low wind speed, this event is characterized by considerable daily heat gains. The constant high SST occurrence in Cendrawasih Bay, which is more than half a year, SSTs can reach more than 30°C and are dominated by the condition of low wind speed, i.e., 80% wind speed of fewer than four m/s along the years, also solar radiation measured more than 200 W/m^2 , those are the solid evidence for define Cendrawasih Bay as the "Cendrawasih Hot Pool."

ACKNOWLEDGMENTS

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REFERENCES

- Arking A, Ziskin D (1994), Relationship between clouds and sea surface temperatures in the western tropical Pacific, *J. Climate*, 7, 988-1000.
- Brown OB, Minnett, PJ (2009), MODIS infrared sea surface temperature algorithm theoretical basis document", version 2.0., (http://modis.gsfc.nasa.gov/data/atbd/atbd_mod25.pdf).
- Chongyin L, Mingquan M, Guangqing Z (1999) The variation of warm pool in the equatorial western Pacific and its impacts on climate, *Advances in Atmos. Sci.*, 16(3): 378-394
- Clement A, Seager R (1999) Climate and the tropical oceans. *Am Meteorol Soc* 12: 3383-3401
- Clement AC, Seager R, Murtugudde R (2005) Why are there tropical warm pools? *J Climate* 18: 5294-5310
- Esaias, WE *et al.* (1998) An overview of MODIS capabilities for ocean science observations, *IEEE Transactions on Geoscience and Remote Sensing* 36(4): 1250-1265. doi: 10.1109/36.701076.
- Ghanea M, Moradi M, Kabiri K, Mehdinia A (2015) Investigation and validation of MODIS SST in the northern Persian Gulf, *Adv. Space Res.*, 57(1) <http://dx.doi.org/10.1016/j.asr.2015.10.040>
- Harweijer C, Seager R, Winton M, Clement A (2005) Why ocean heat transport warms the global mean climate. *Tellus* 57A: 662-675
- Hosoda K (2013) Empirical method of diurnal correction for estimating sea surface temperature at dawn and noon. *J Oceanogr* 69: 631-646 doi: 10.1007/s10872-013-0194-4
- Ihsan EN, Enita SY, Kunarso, Wirasatriya A (2018) Oceanographic Factors in Fishing Ground Location of Anchovy at Teluk Cenderawasih National Park, West Papua: Are These Factors Have an Effect of Whale Sharks Appearance Frequencies? *IOP Conf. Ser.: Earth Environ. Sci.*, 116(2018): 012017, doi :10.1088/1755-1315/116/1/012017
- Kawamura H, Qin H, Ando K (2008) In-situ diurnal sea surface temperature variations and near-surface thermal structure in the tropical Hot Event of the Indo-Pacific warm pool. *J Oceanogr* 64: 847-857
- Lee MA, Tzeng MT, Hosoda K, Sakaida F, Kawamura H, Shieh WJ, Yang Y, Chang Y (2010) Validation of JAXA/MODIS sea surface temperature in water around Taiwan using

the Terra and Aqua satellites, *Terr. Atmos. Ocean Sci.* 21(4): 727–736.

Pierrehumbert RT (2000) Climate change and the tropical Pacific: The sleeping dragon wakes. *PNAS* 97(4): 1355-1358

Qin H, Kawamura H, Kawai Y (2007) Detection of hot event in the equatorial Indo-Pacific warm pool using advanced satellite sea surface temperature, solar radiation, and wind speed. *J Geophys Res* 112: C07015, doi:10.1029/2006JC003969

Qin H, Kawamura H, Sakaida F, Ando K (2008) A case study of the tropical hot event in November 2006 (HE0611) using a geostationary meteorological satellite and the TAO/TRITON mooring array. *J Geophys Res* 113: C08045. doi:10.1029/2007JC004640

Qin H, Kawamura H (2009) Atmosphere response to a hot SST event in November 2006 as observed by AIRS instrument. *Adv Space Res* 44: 395-400. doi: 10.1016/j.asr.2009.03.003

Qin H, Kawamura H (2010) Air-sea interaction throughout the troposphere over a very high sea surface temperature. *Geophys Res Lett*, 37: 1-4. doi:10.1029/2009GL041685

Qin H, Chen G, Wang W, Wang D, Zeng L (2014) Validation and application of MODIS-derived SST in the South China Sea, *Int. J. Remote Sens.* 35(11–12): 4315–4328.

Ramanathan V, Collins W (1991) Thermodynamic regulation of ocean warming by cirrus clouds deduced from observations of the 1987 El Niño. *Nature* 351: 27–32

Swandiko M, Wirasatriya A, Marwoto J, Muslim, Indrayanti E, Subardjo P, Ismunarti DH (2021) Study of the persistence of high SST (>30°C) in the Malacca Strait. *Buletin Oseanografi Marina*, 10(2): 162-170, *in bahasa*

Thoron TG, Rosenthal Y, Bassinot F, Beaufort L (2005) Stable sea surface temperatures in the western Pacific warm pool over the past 1.75 million years. *Nature* 433: 294-298

Tita ADC, Wirasatriya A, Sugianto DN, Maslukah L, Handoyo G, Hariyadi, Helmi M, Avianto P (2020) Persistence of high sea surface temperature (> 30°C) in Tomini Bay. *IOP Conf. Ser.: Earth Environ. Sci.*, 530(2020): 012038, doi:10.1088/1755-1315/530/1/012038

Waliser DE, Graham NE (1993) Convective cloud systems and warm-pool sea-surface temperatures: Coupled interactions and self-regulation. *J Geophys Res* 98 (D7): 12881–12893

Wallace JM (1992) Effect of deep convection on the regulation of tropical sea surface temperature. *Nature* 357: 230–231

Wirasatriya A, Kawamura H, Shimada T, Hosoda K (2015) Climatology of hot events in the western equatorial Pacific. *Journal of Oceanography* 71: 77-90 doi: 10.1007/s10872-014-0263-3

Wirasatriya A, Kawamura H, Shimada T, Hosoda K (2016) Atmospheric structure favoring high sea surface temperatures in the western equatorial Pacific. *Journal of Geophysical Research* 121: 1-14 doi:10.1002/2016JD025268.

Wirasatriya A, Sugianto DN, Helmi M (2017) The Influence of Madden Julian Oscillation on the Formation of the Hot Event in the Western Equatorial Pacific, 2nd International Conference on Tropical and Coastal Region Eco Development 2016, IOP Conf. Series: Earth and Environmental Science 55: 012006 doi:10.1088/1755-1315/55/1/012006

Wirasatriya, A., R.Y. Setiawan, and P. Subardjo (2017) The effect of ENSO on the variability of chlorophyll-a and sea surface temperature in the Maluku Sea, *IEEE J. Sel. Topics Appl. Earth Observ. Remote Sens. (JSTARS)*, vol. 10, no. 12, pp. 5513-5518, 2017.

Wirasatriya A, Kawamura H, Helmi M, Sugianto DN, Shimada T, Hosoda K, Handoyo G, Putra YDG, Koch M (2020) Thermal structure of Hot Events and their possible role in maintaining the warm isothermal layer in the Western Pacific warm pool. *Ocean Dynamics*, 70:771–786.

Wyrтки K (1989) Some thoughts about the west Pacific warm pool. paper presented at Western Pacific International Meeting and Workshop on TOGA COARE, Nouméa, New Caledonia: 99-109

Yan, X-H, Ho CR, Zheng Q Klemas V (1992) Temperature and size variabilities of the western Pacific warm pool. *Science* 258: 1643-1645

Yu L, Weller RA (2007) Objectively Analyzed air-sea heat Fluxes for the global ocean-free oceans (1981–2005). *Bull Ameri Meteor Soc* 88: 527–539.



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22 Februari 2023 pukul 16.19

Dear Dr Abd. Rasyid Jalil!

We have reached a decision regarding your submission to GEOGRAPHY, ENVIRONMENT, SUSTAINABILITY, "CENDRAWASIH HOT POOL: STUDY OF THE HIGH-FREQUENCY OCCURRENCE OF HIGH SST PHENOMENA AT CENDRAWASIH BAY, PAPUA". Our decision is **revisions required**.

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A. Category and general nature of the manuscript:
Empirical (based on own research)

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Do you consider the content of the paper of high quality and originality?:
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Are the results clearly explained?:

Rather yes

Are the conclusions of the authors consistent with the results obtained during the study?:

Definitely yes

Are there any limitations of the study that should be noted?:

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Is the quality of the figures and tables satisfactory?:

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:

The paper has the potential to be published after revision.

Comments:

1. The title is not good, please revise it.
2. Please check through the grammatical error. The sentence structure should be appropriate.
3. Once defined, the acronyms for some words should be used along with the paper. Sometimes the acronyms are directly used, for example SST and so on.
4. Please improve colorbar in Fig 1.
5. It is recommended to introduce research question in the introduction section. Why the HE is important? What is the effect of the HE events on environment?
6. Introduction section needs to be revised and added more previous study to show the gap of knowledge.
7. It is obvious that there are differences between spatial resolution and temporal of datasets, however, are there any effect on results? Please describe it.
8. My major concern consists in the structure of the Methodology description. This would make it more understandable and let ease replication of the procedures.
9. Thirteen years of observation (2003-2015) able to represent the climatology of SST? Why not use data for longer period?
10. Fig. 5 should be improved.
11. It is recommended to add more analysis and results for supporting the conclusion.

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Author's Responses to Reviewer Comments and Suggestions

The paper has the potential to be published after revision.

General Response: Thanks for your suggestion, question, and criticism for improving our manuscript. In this revised version, we have explained the method in more detail, improved figures, and added more analysis of the temporal change of SST, wind speed, and solar radiation. The revised parts are highlighted in yellow.

1. The title is not good, please revise it.

Response: Thank you very much for the suggestion. We have revised the title to CENDRAWASIH HOT POOL: THE FREQUENT HIGH SST PHENOMENA AT CENDRAWASIH BAY, PAPUA.

2. Please check through the grammatical error. The sentence structure should be appropriate.

Response: Thank you very much for the comment. We have carefully checked the grammar. Hopefully, this revised version is better.

3. Once defined, the acronyms for some words should be used along with the paper. Sometimes the acronyms are directly used, for example SST and so on.

Response: Thank you very much for the correction. We have fixed it.

4. Please improve colorbar in Fig. 1

Response: The colorbar in Fig 1 has been improved.

5. It is recommended to introduce research question in the introduction section. Why the HE is important? What is the effect of the HE events on environment?

Response: Thank you very much for the recommendation. We have added more explanation to describe how HE has an important role in regulating global climate through its contribution to Pacific warm pool formation. L87-96

6. Introduction section need to needs to be revised and added more previous study to show the gap of knowledge.

Response: Thank you very much for the suggestion. The knowledge gap in this study is that there are few studies about the high SST phenomena within the Indonesian seas. (L97-101). We have emphasized the gap at L101-102.

7. It is obvious that there are differences between spatial resolution and temporal of dataset, however, are there any effect on results? Please describe it.

Response: Thank you very much for the question. The spatial resolution differs among datasets, but this is OK since the analysis is not based on the grid-by-grid analysis. Descriptively we compare the spatial distribution of all datasets and then take the mean values in 2 areas as a sample for investigating the temporal variation. Solar radiation differs from others for the observation period since the dataset was only available in 2009. But we think this is OK since the same method is applied by Wirasatriya et al. (2015, 2016, 2020), Tita et al. (2020), and Swandiko et al. (2021)

8. My major concern consists in the structure of the Methodology description. This would be more understandable and let ease replication of the procedures.

Response: Thank you very much for the suggestion. We have described the methodology clearer than before.

9. Thirteen years of observation (2003-2015) able to represent the climatology of SST? Why not use data for longer period?

Response: Thank you very much for the question. We follow the period of observation by Wirasatriya et al. (2015, 2016, 2020), Tita et al. (2020), and Swandiko et al. (2021)

10. Fig. 5 should be improved.

Response: Thank you very much for the suggestion. It is done.

11. It is recommended to add more analysis and results to support the conclusion

Response: Thank you very much for the suggestion. We have added the analysis of the temporal change of SST, wind speed, and solar radiation in the area inside and outside Cenderawasih Bay (Fig. 5, L214-237)

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22
23 **CENDERAWASIH HOT POOL: THE FREQUENT HIGH SST**
24 **PHENOMENA AT CENDERAWASIH BAY, PAPUA**

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26
27 **ABSTRACT.** The warm pool has been used to define a water body with the
28 characteristic of SST exceeding 28°C within a particular area and a relatively long period
29 in an annual circle. However, there are regions with an annual mean SST measured above
30 30°C, and we classified them as Hot Pools because of the conditions of intense solar
31 radiation and low wind speed. One of the Hot Pool spots was found in Indonesia, in
32 Cenderawasih Bay. ~~In the present study, the existence of the Cenderawasih Hot Pool is~~
33 ~~examined~~The present study examines the existence of the Cenderawasih Hot Pool using
34 long-term observation of satellite SST data. We also used surface wind data, surface heat
35 flux, and surface current to investigate their mechanisms. The results show that within 13
36 years of observation (2013-2015), SSTs in Cenderawasih Bay can reach more than 30°C
37 with 50% occurrence within the ~~period of observation~~observation period. Heat input comes
38 from strong solar radiation, i.e., 50% of solar radiation is more than 200 W/m². The location
39 is also dominated by low wind speed, i.e., 80% wind speed of lower than 4 m/s, and
40 ~~these~~which caused the low latent loss in CenderawasihCenderawasih Bay. ~~For surface~~
41 ~~current, Cenderawasih Bay is fully isolated~~Cenderawasih Bay is fully isolated for surface
42 ~~currents~~ currents during the dry and wet seasons, since the easterly subsurface water flow does not
43 enter the bay. The absence of strong currents prevents the mixing process, maintaining the
44 high temperature in the surface layer. Those processes are figured out and become strong
45 evidence to justify CenderawasihCenderawasih Bay as one of the Hot Pool areas within the
46 Indonesian seas.

47
48 **KEYWORDS:** Global climate; Sea Surface Temperature; Hot Pool Spot;

49 ~~Cenderawasih~~Cenderawasih Bay.

51 INTRODUCTION

52 The western equatorial Pacific significantly influences the global climate. The warm
53 pool, a region with average sea surface temperatures (SST) above 28°C, has a specific
54 impact on the Earth's circulation (e.g., Wyrski 1989; Yan et al. 1992; Clement and Seager
55 1999; Chongyin et al. 1999; Pierrehumbert 2000; Clement et al. 2005; Thoron et al. 2005;
56 Herweijer et al. 2005).

57 On the other hand, high SST in tropical regions has attracted researchers to investigate
58 the mechanisms since the formation of high SST requires a particular atmospheric process
59 (e.g., Ramanathan and Collins 1991; Wallace 1992; Arking and Ziskin 1994). This process
60 is depicted by Waliser and Graham (1993), which shows the relation between SSTs and
61 deep convection. They utilized highly reflective cloud information from an arbitrary
62 examination of monthly SST data with a grid spacing of 2° (produced from combined
63 satellite observation and in situ data) and daily visible and infrared satellite image data.
64 The highly reflective cloud increases along with the SST as it rises from 26°C to 29.5°C.
65 In contrast, the highly reflective cloud diminishes with increased SST in the temperature
66 from 29.5°C to 32°C. As a result, the analysis proved that several atmospheric processes
67 impacted SSTs below and above 29.5°C.

68 By taking advantage of high temporal and spatial resolution SST products derived
69 from satellite observations (i.e., daily and ≤ 25 km×25 km), several studies (Kawamura et
70 al. 2008; Qin et al. 2007, 2008; Qin and Kawamura 2009, 2010; Wirasatriya et al. 2015,
71 2016, 2017a, 2020) were able to identify high SST events (i.e., more than 30°C) in specific
72 areas and at certain periods and define them as Hot Event (HE). In summary, they
73 concluded that considerable daily heat gains characterize the production of HE under high
74 solar radiation and low wind speed brought on by "remote convection" mechanisms.

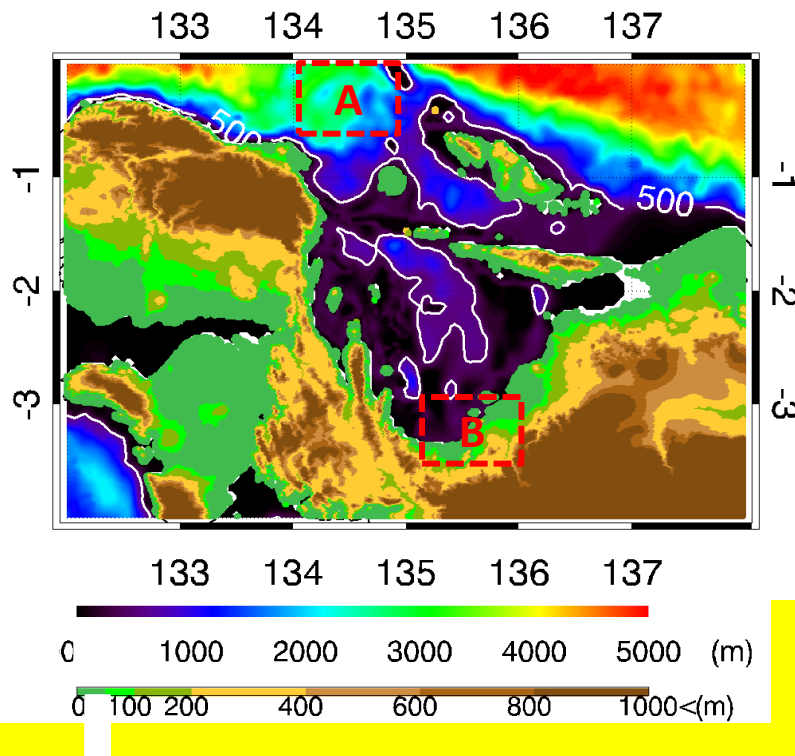
75 Wirasatriya et al. (2015) elaborated on the climatology of HEs in the western equatorial
76 Pacific using the SST dataset derived from satellite microwave sensors. Throughout nine
77 years of observation (2003-2011), they discovered 71 HE cases in the western equatorial
78 Pacific, with the majority centered on the Solomon Islands and New Guinea Island's
79 northern coasts, which extend eastward up to 160°W. According to the climatology, the
80 region ~~is located in an area with~~has solar radiation of more than 200 W/m² and wind speeds
81 of less than 4 m/s. Low wind speeds heavily influence the mechanism for HE incidence in
82 the western equatorial Pacific. Much of the equatorial region experiences sun radiation
83 above 200 W/m² during the HE periods. Low wind speeds minimize latent heat loss, which
84 results in high SSTs and HEs in specific locations. Wirasatriya et al. (2015) also
85 emphasized that high solar radiation and low wind speed are much more common during
86 the development stage and less common during the decay stage. This study also
87 demonstrated that a rise in the long-term mean SST in the western equatorial Pacific is
88 correlated with an increase in the frequency of HE events. HEs were responsible for 51.5%
89 of the SSTs >30°C in the warm pool region bounded by the 29.5 °C isotherms of the
90 climatological SST. Thus, statistically, there is a relation between the occurrence of HE and
91 the formation of the western Pacific warm pool. Moreover, Wirasatriya et al. (2020)
92 demonstrated the role of HE in maintaining the warm mixed layer in the western Pacific
93 warm pool. The frequent occurrence of HE transports heat from the surface layer to the
94 deeper layer above the thermocline to maintain the warm mixed layer in the warm pool.
95 Since the Pacific warm pool influences the ~~variability of the global climate through the~~
96 ~~process of global climate's variability through~~ coupled ocean-atmosphere dynamics and
97 thermodynamics (e.g., Clement and Seager 1999; Herweijer et al. 2005), this becomes an
98 example of the contribution of high SST phenomena to regulate the global climate.

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99 Within the Indonesian seas, the frequent appearance of high SST > 30°C has been
 100 reported by Tita et al. (2020) in Tomini Bay and by Swandiko et al. (2021) in the Malacca
 101 Strait. Their appearances also the high solar radiation and weak wind. The morphology of
 102 the semi-enclosed waters causes the high-frequency occurrence of low wind speed by less
 103 than 4 m/s in both areas. Besides these two areas, there is no more study on the high SST
 104 phenomena within the Indonesian Seas.

105 In the present study, we demonstrate The present study demonstrates the frequent high
 106 SST occurrence (more than 30°C) in Cenderawasih Bay. Cenderawasih Bay is located
 107 in northern Papua Island, part of the western Pacific warm pool. It is semi-enclosed water
 108 with a deep basin, surrounded by mountain chains (Fig. 1). Cenderawasih Bay is the habitat
 109 for whale sharks (Ihsan et al. 2018). Since the area determines warm pool definition with
 110 SST climatology higher than 28°C, we call the high-frequency occurrence of SSTs higher
 111 than 30°C in the Cenderawasih Bay the Cenderawasih Hot Pool. Wirasatriya et al.
 112 (2015) found that the duration of HE occurrence in the western equatorial Pacific is no
 113 longer than two months since HE requires a typical condition of high solar radiation and
 114 low wind speed. Thus, it is interesting to understand the mechanisms of the high-frequency
 115 occurrence of high SST in the Cenderawasih Hot Pool.



118
 119 **Fig. 1. Bathymetry and topography of Cenderawasih Bay. The dashed**
 120 **boxes A, and B represent the sampling area for outside and inside**
 121 **Cenderawasih Bay, respectively, as shown in Fig. 5.**

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MATERIALS AND METHODS

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We used the semi-daily 11 μm SST products from Moderate Resolution Imaging Spectroradiometer (MODIS) Aqua Lv3 with a spatial resolution of $0.04^\circ \times 0.04^\circ$ (Esaias et al. 1998) and observation period from 2003 to 2015. The MODIS SST 11 μm Multi-Channel SST algorithm generates m by using brightness temperatures at 11 μm and 12 μm (Brown and Minnet 2009). The best accuracy was achieved by validating and testing this data against in-situ measurements (e.g., Lee et al. 2010; Qin et al. 2014; Ghanea et al. 2015).

We can obtain the Cross-Calibrated Multi-Platform (CCMP) gridded surface vector winds version 2.0 from the following website: www.remss.com, and use it to calculate surface winds. As a result, CCMP is regarded as a Level-3 ocean vector wind analysis product. They are created utilizing data from moored buoys, satellites, and model winds. The surface wind data for spatial and temporal resolutions is $0.25^\circ \times 0.25^\circ$ respectively and quarter daily. This wind product has greater accuracy than other wind reanalysis data.

We used a 0.5° grid from the Objectively Analyzed Air-Sea Fluxes project to analyze daily latent heat fluxes and solar radiation (Yu and Weller, 2007). The period of observation for latent heat flux is from 2003 to 2015, but for solar radiation is only from 2003 to 2009 due to the availability of the data. Global 30 Arc-Second Elevation data was used to obtain topography. Global 30 Arc-Second Elevation (GTOPO30) data was used to obtain topography (https://www.usgs.gov/centers/eros/science/usgs-eros-archive-digital-elevation-global-30-arc-second-elevation-gtopo30?qt-science_center_objects=0#qt-science_center_objects). ETOPO, a one-arc-minute global relief model of Earth's surface that combines land topography and ocean bathymetry, provides bathymetry (<https://www.ngdc.noaa.gov/mgg/global/>). For oceanic parameters, we used the current reanalysis data from <http://marine.copernicus.eu/> to obtain the seasonal variation of the current pattern at 100 m depth in the Cenderawasih Bay Hot Pool.

To calculate the frequency occurrence of SST within 13 years of observation, we calculate the percentage of SST $> 30^\circ\text{C}$ occurrence in each grid. Since high SST of more than 30°C occurs during the condition of low wind speed (i.e., < 4 m/s) and high solar radiation (i.e., > 200 W/m²) (Wirasatriya et al. 2015), we calculated the percentage of low wind speed and high solar radiation in the Cenderawasih Bay. Furthermore, we also calculated the percentage of low latent heat release (< 120 W/m²) to explain how wind speed influences the variability of SST. The equation for calculating the percentage is as follows:

$$\%(x, y) = \frac{1}{n} \sum_{i=1}^n p_i(x, y) \times 100\% \quad (1)$$

where $\%(x, y)$ is percentage the percentage of high SST $> 30^\circ\text{C}$ or percentage of weak wind speed < 4 m/s or percentage of solar radiation > 200 W/m² or percentage of latent heat flux < 120 W/m² at position (x,y); p_i is the amount of SST data $> 30^\circ\text{C}$ or the amount of weak wind speed < 4 m/sec or the amount solar radiation > 200 W/m² or the amount of latent heat flux < 120 W/m² at position (x,y). Furthermore, we also analyze the climatological mean of each parameter. To create monthly and monthly climatology, all geophysical parameters are incorporated as follows: (Wirasatriya et al. 2017b)

$$\bar{X}(x, y) = \frac{1}{n} \sum_{i=1}^n x_i(x, y, t) \quad (2)$$

where $\bar{X}(x, y)$ is monthly mean value or monthly climatology value at position (x,y),

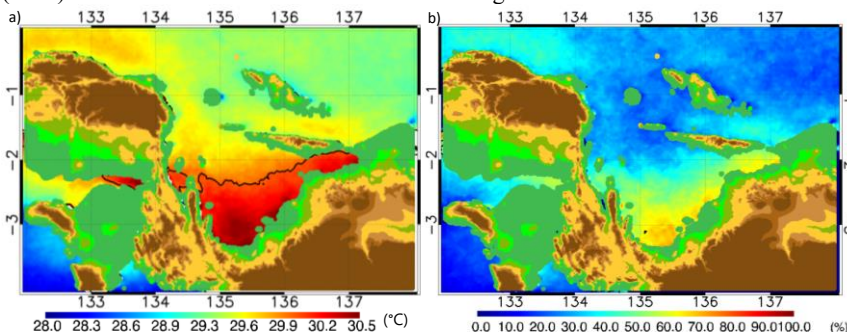
169 $x_i(x, y, t)$ is i^{th} value of the data at (x, y) position and time t . Next, n is the number of data
 170 in 1 month and the number of monthly data in 1 period of climatology (i.e., from 2003 to
 171 2016 = 13 data) for monthly calculation and monthly climatology calculation, respectively.
 172 Moreover, if pixel x_i is a hollow pixel, it is not included in the calculation.

173 RESULTS AND DISCUSSION

174 Cenderawasih Cenderawasih Hot Pool

175 The definition of the Cenderawasih Cenderawasih Hot Pool is described in Fig. 2, which
 176 shows the climatological mean of SST and the percentage of daily SSTs more than 30°C
 177 from 2003 to 2015 (13 years). In Cenderawasih Cenderawasih Bay, the mean SST at the
 178 southern part of 2°S is more than 30°C, and the high SSTs are more than 50%. It means
 179 that the mean SST in Cenderawasih Cenderawasih Bay is higher than the definition of the
 180 western Pacific warm pool which is 28°C (Wyrtki 1989), and ~~totally~~ more than 3.5 years
 181 within 2003-2015, SSTs in Cenderawasih Cenderawasih Bay can reach more than 30°C. At
 182 the southernmost of the bay, the mean SST reaches 30.5°C, with the percentage of high
 183 SSTs being more than 70%. It has become the hottest area of Cenderawasih Cenderawasih
 184 Bay.

185 In contrast, in the area outside Cenderawasih Cenderawasih Bay, i.e., the western Pacific
 186 warm pool's offshore seas, the percentage is less than 30%. Thus, this evidence supports
 187 the definition of Cenderawasih Cenderawasih Bay as the Cenderawasih Cenderawasih Hot
 188 Pool. The high-frequency occurrence of high SST in Cenderawasih Cenderawasih Bay may
 189 correspond to whale sharks' ~~appearance throughout the year~~ yearly appearance. Ihsan et al.
 190 (2018) indicated the tolerance of whale sharks in high SST conditions.

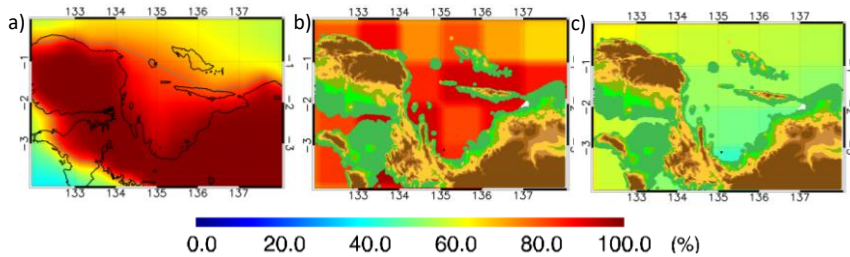


191
 192 **Fig. 2. Climatological mean of SST (a) and Percentage-percentage of daily SSTs >**
 193 **30°C during thirteen years of observation (2003-2015) (b). The black contour is**
 194 **30°C.**

196 Atmospheric aspect of Cenderawasih Cenderawasih Hot Pool

197 Fig. 3 clearly shows that Cenderawasih low wind speed dominates Cenderawasih Bay
 198 Bay is dominated by the condition of low wind speed, i.e., 80% wind speed < 4 m/s
 199 occurred during 2003-2015. This causes low latent loss in Cenderawasih Cenderawasih Bay.
 200 For solar radiation, the percentage of high solar radiation inside and outside
 201 Cenderawasih Cenderawasih Bay is almost similar i.e., nearly 50%, indicating that high solar
 202 radiation is not dominant in both areas. This means that, although high solar radiation does
 203 not frequently occur, the absence of strong wind speed may maintain the latent heat loss to
 204 keep the SST in Cenderawasih Cenderawasih Bay higher than 30°C.
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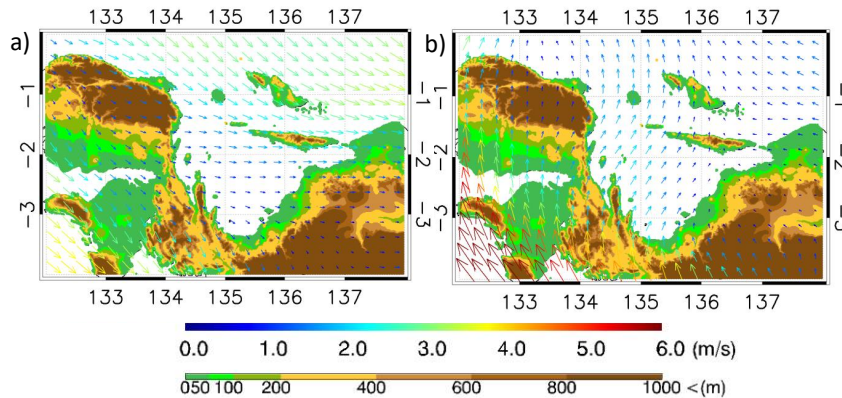


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208 **Fig. 3. The percentage of wind speed less than four m/s (a), latent heat release less**
 209 **than 120 W/m² (b), and solar radiation more than 200 W/m² (c).**

210

211 To investigate how low wind speed occurs in the study area, we show the monthly
 212 climatology map of surface wind during summer and winter (Fig. 4). For both seasons,
 213 Cenderawasih Bay is protected from high wind speed due to high mountain
 214 chains in the western, southern, and eastern parts. In the northern part, small islands at the
 215 mouth of the bay prevent the strong wind speed from entering the bay. Thus, the role of
 216 topography is crucial for the occurrence of constant high SSTs in
 217 Cenderawasih Bay. The same tendencies are also found at the Tomini Bay and
 218 Malacca Strait, as indicated by Tita et al. (2020) and Swandiko et al. (2021).



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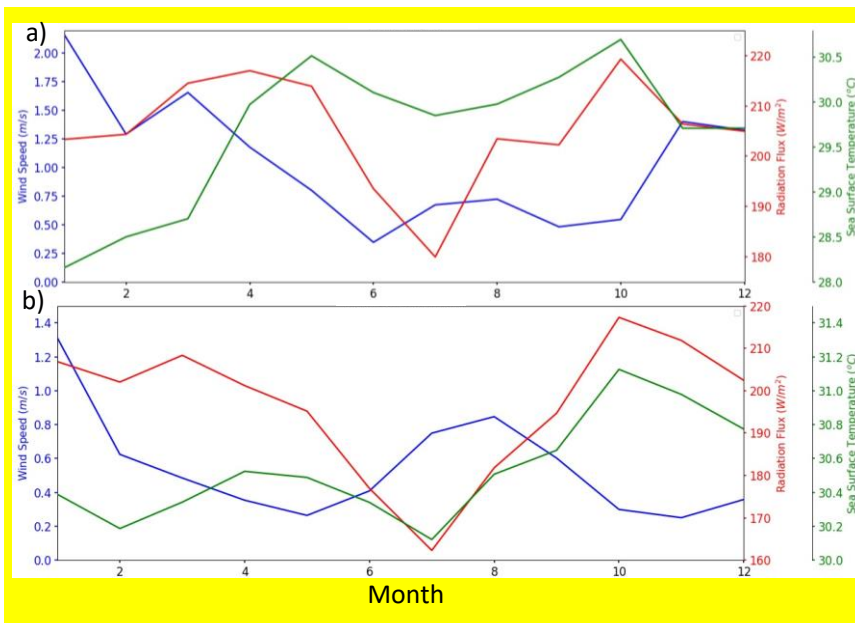
220 **Fig. 4. Monthly climatology of surface wind speed during a) wet season (January)**
 221 **and b) dry season (August)**

222 The influence of wind speed and solar radiation on the SST variation is further
 223 examined by plotting the monthly climatology of wind speed, solar radiation, and SST at
 224 the area A, and B, areas A and B (Fig. 1) that represent the areas outside and inside
 225 Cenderawasih Bay. Seasonal variations of SSTs are observed in both areas. ~~t~~
 226 The minimum SSTs occur during the wet and dry seasons, while the maximum SSTs during
 227 the transition season. This seasonal SST variability is similar ~~with to~~ the other areas in the
 228 Indonesian seas, such as the Java Sea (Wirasatriya et al. 2018), Maluku Sea (Wirasatriya
 229 et al. 2019), Halmahera Sea (Setiawan et al. 2019), etc.

230 For the area outside Cenderawasih Bay, SST ranges from 28.2°C to
 231 30.6°C. It is ~~clearly~~ seen that the variability of SST is controlled by wind speed and solar

232 radiation. From January to June, when wind speed decreases from 2.2 m/s to 0.3 m/s and
 233 solar radiation is more than 200 W/m², SST increases from 28.2°C to 30.5°C. From June
 234 to October, wind speed is lower than 1 m/s. In the absence of strong wind, the low SST in
 235 July is caused by the decrease of solar radiation into a minimum.

236 For the area inside Cendrawasih Bay, persistent high SST is observed as
 237 the SST ranges from 30.2°C to 31.1°C. It may be due to the low wind speed of
 238 due to the low wind speed that is less than 1 m/s almost throughout fewer than 1 m/s for
 239 years. Thus, the variability of SST is clearly seen as ruled by solar radiation. The minimum
 240 SST in July is caused by the minimum solar radiation solar radiation causes the minimum
 241 SST in July. However, the absence of strong wind prevents latent heat loss that maintains
 242 the high SST inside the Cendrawasih Bay.
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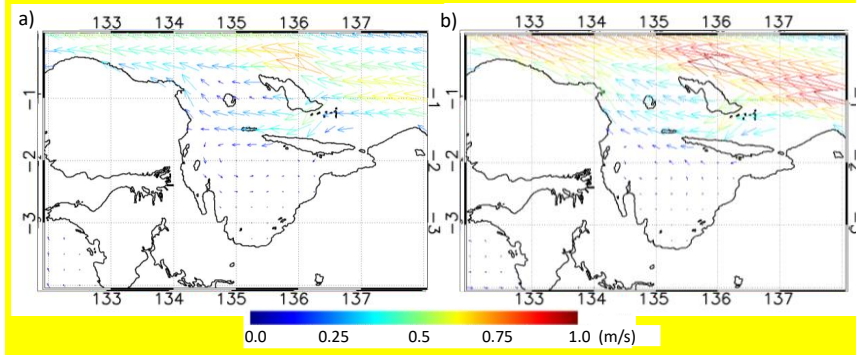


244 **Fig. 5. Monthly climatology of SST, wind speed, and solar radiation at a) area**
 245 **outside and b) area inside Cendrawasih Bay, as shown in Fig. 1.**
 246
 247

248 **The oceanic aspect of Cendrawasih Hot Pool**

249 To investigate the oceanic aspect of the Cendrawasih Hot Pool, we
 250 plotted the current patterns at 100 m depth shown in Fig. 6. A northwestward current has
 251 been identified north of Cendrawasih Bay. During the wet season (January),
 252 the speed of this current ranges from 0.5 m/s to 0.7 m/s. During the dry season (August),
 253 when the northwesterly wind disappears (Fig. 4b), the speed of the northwestward current
 254 increases. This northwestward current is known as the New Guinea Coastal Under Current
 255 (NGCUC). NGCUC is a permanent current feature at 200 m depth regardless of the wind
 256 reversals (Tsuchiya et al. 1989). Inside Cendrawasih Bay, strong currents are
 257 absent in this area which prevents, preventing the mixing process from maintaining the
 258 high temperature in the surface layer. During the dry and wet seasons,
 259 Cendrawasih Bay is fully isolated since the westward subsurface water flow
 260 does not enter Cendrawasih Bay. Thus, this isolated basin causes the

261 Cenderawasih Bay area to be influenced mainly by the air-sea interaction
262 process mentioned in the previous section creating Cenderawasih Hot Pool.



263
264 **Fig. 6. Monthly climatology of current at 100 m depth in January (a) and August (b)**

265

266 CONCLUSIONS

267 The hot Pool is used to categorize high SST episodes ($>30^{\circ}\text{C}$) in particular areas and
268 during specific periods (relatively long periods). Since the "warm pool" has been used to
269 define an area with the annual average SST distributed below 30°C in dominant, therefore
270 "hot pool" was taken to represent the area with SST dominated above 30°C .

271 Under high solar radiation and low wind speed, this event is characterized by
272 considerable daily heat gains. The constant high SST occurrence in
273 Cenderawasih Bay, which is more than 50% of 13 years, SSTs can reach more
274 than 30°C and are dominated by the condition of low wind speed, i.e., 80% wind speed is
275 less than m/s along the years, also 50% solar radiation measured is more than 200 W/m^2 .
276 The current pattern at 100 m depth also shows that NGCUC does not enter
277 Cenderawasih Bay. Thus, this indicates that Cenderawasih bay
278 is an isolated water. Those are the solid evidence for defining Cenderawasih
279 Bay as the "Cenderawasih Hot Pool."

280

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285

286 REFERENCES

287 Arking A., and Ziskin D. (1994). Relationship between clouds and sea surface
288 temperatures in the western tropical Pacific. *Journal of Climate*, 7(6), 988-1000.
289 [https://doi.org/10.1175/1520-0442\(1994\)007%3C0988:RBCASS%3E2.0.CO;2](https://doi.org/10.1175/1520-0442(1994)007%3C0988:RBCASS%3E2.0.CO;2)

290

291 Brown O.B., and Minnett, P.J. (2009). MODIS infrared sea surface temperature
292 algorithm theoretical basis document, version 2.0. [online]. Available at:
293 https://modis.gsfc.nasa.gov/data/atbd/atbd_mod25.pdf. (http://modis.gsfc.nasa.gov/data/atbd/atbd_mod25.pdf) [Accessed 25 October 2022].

294

295
296 Chongyin L., Mingquan M., and Guangqing Z. (1999). The variation of warm pool in
297 the equatorial western Pacific and its impacts on climate. *Advances in Atmospheric*

348 Pierrehumbert R.T. (2000). Climate change and the tropical Pacific: The sleeping
349 dragon wakes. Proceedings of the National Academy of Sciences, PNAS 97(4), 1355-1358.
350 <https://doi.org/10.1073/pnas.97.4.1355>

351
352 Qin H., Kawamura H., ~~and~~ Kawai Y. (2007). Detection of ~~hot~~ Hot event-Event in the
353 equatorial Indo-Pacific warm pool using advanced satellite sea surface temperature, solar
354 radiation, and wind speed. Journal of Geophysical Research: Oceans, J Geophys Res 112:
355 (C07)015. <https://doi.org/10.1029/2006JC003969>, doi:10.1029/2006JC003969

356
357 Qin H., Kawamura H., Sakaida F., ~~and~~ Ando K. (2008). A case study of the tropical
358 hot event in November 2006 (HE0611) using a geostationary meteorological satellite and
359 the TAO/TRITON mooring array. Journal of Geophysical Research: Oceans, 113(C8), J
360 Geophys Res 113: C08045. <https://doi.org/10.1029/2007JC004640>
361 -doi:10.1029/2007JC004640

362 Qin H. ~~and~~ Kawamura H. (2009). Atmosphere response to a hot SST event in
363 November 2006 as observed by AIRS instrument. Advances in space research, 44(3), 395-
364 400. Adv Space Res 44: 395-400. <https://doi.org/10.1016/j.asr.2009.03.003>
365 doi:10.1016/j.asr.2009.03.003

366 Qin H. ~~and~~ Kawamura H. (2010). Air-sea interaction throughout the troposphere over
367 a very high sea surface temperature. Geophysical research letters, 37(1) Geophys Res Let,
368 37, 1-4. <https://doi.org/10.1029/2009GL041685>
369 -doi:10.1029/2009GL041685

370 Qin H., Chen G., Wang W., Wang D., ~~and~~ Zeng L. (2014). Validation and application
371 of MODIS-derived SST in the South China Sea. International journal of remote
372 sensing, 35(11-12), 4315-4328. <https://doi.org/10.1080/01431161.2014.916439>
373 Int. J. Remote Sens. 35(11-12): 4315-4328.

374 Ramanathan V. ~~and~~ Collins W. (1991). Thermodynamic regulation of ocean warming
375 by cirrus clouds deduced from observations of the 1987 El Niño. Nature 351 (6321), 27-
376 32. <https://doi.org/10.1038/351027a0>

377
378 Setiawan, R. Y., A. W Wirasatriya A., U. Hernawan U., S. L Leung S., and I. Iskandar I.
379 (2020). Spatio-temporal variability of surface chlorophyll-*a* in the Halmahera Sea and its
380 relation to ENSO and the Indian Ocean Dipole. International Journal of Remote
381 Sensing, 41(1), 284-299. Int. J. Remote Sens. 41(1):284-299,
382 <https://doi.org/10.1080/01431161.2019.1641244>
383 <https://doi.org/10.1080/01431161.2019.1641244>

384 Swandiko M., Wirasatriya A., Marwoto J., Muslim, Indrayanti E., Subardjo P.,
385 Ismunarti D.H. (2021). Study of the persistence of high SST Studi persistensi suhu
386 permukaan laut tinggi (>30°C) in the Malacca Strait di perairan Selat Malaka. Buletin
387 Oceanografi Marina, 10(2), 162-170. [online]. Available at:
388 <https://ejournal.undip.ac.id/index.php/buloma/article/download/31554/19352>. [Accessed
389 25 October 2022] (in Indonesian). *-in-bahasa*

390
391 ~~De Garidel-Thoron T., Rosenthal Y., Bassinot F., and Beaufort, L. Thoron TG,~~
392 ~~Rosenthal Y, Bassinot F, Beaufort L~~ (2005). Stable sea surface temperatures in the western
393 Pacific warm pool over the past 1.75 million years. Nature, 433(7023), 294-298,
394 <https://doi.org/10.1038/nature03189>

395
396 Tita A.D.C., Wirasatriya A., Sugianto D.N., Maslukah L., Handoyo G., Helmi M., and
397 Avianto, P. Tita ADC, Wirasatriya A, Sugianto DN, Maslukah L, Handoyo G, Hariyadi,
398 Helmi M, Avianto P (2020). Persistence of high sea surface temperature (> 30°C) in Tomini

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450 COARE, Nouméa, New Caledonia: 99-109 [online]. Available at:
451 https://horizon.documentation.ird.fr/exl-doc/pleins_textes/doc34-08/30195.pdf#page=114,
452 (Accessed 26 October 2022).

453
454 Yan, X.-H., Ho C.R., Zheng Q., and Klemas V. (1992). Temperature and size
455 variabilities of the western Pacific warm pool. Science, 258 (5088), 1643-1645.
456 <https://doi.org/10.1126/science.258.5088.1643>

457
458 Yu L., and Weller R.A. (2007). Objectively Analyzed air-sea heat Fluxes-fluxes for the
459 global ice-free oceans (1981–2005). Bulletin of the American Meteorological
460 Society, 88(4), 527-540. <https://doi.org/10.1175/BAMS-88-4-527>,
461 Bull Ameri Meteor Soc 88: 527–539.

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Abdul Rasyid <abdulrasyid.fayufi@gmail.com>

GES Editor Decision

Иванов Владимир Владимирович <vladimir.ivanov@aari.ru>
Kepada: abdulrasyid fayufi <abdulrasyid.fayufi@gmail.com>

11 Maret 2023 pukul 19.31

Dear Dr Abd. Rasyid Jalil ,

I have received revised version of your manuscript and I will sent it to the reviewer shortly.

Kind regards,
Vladimir Ivanov

От: "abdulrasyid fayufi" <abdulrasyid.fayufi@gmail.com>
Кому: "Иванов Владимир Владимирович" <vladimir.ivanov@aari.ru>
Копия: "GES Journal" <ges-journal@geogr.msu.ru>
Отправленные: Суббота, 11 Март 2023 г 0:16:48
Тема: Re: GES Editor Decision

[Kutipan teks disembunyikan]



Abdul Rasyid <abdulrasyid.fayufi@gmail.com>

GES Editor Decision

Abdul Rasyid <abdulrasyid.fayufi@gmail.com>

12 April 2023 pukul 00.10

Kepada: Иванов Владимир Владимирович <vladimir.ivanov@aari.ru>

Dear Vladimir Ivanov
Editor of GES journal

We just want to know the progress of the manuscript entitled "CENDRAWASIH HOT POOL: STUDY OF THE HIGH-FREQUENCY OCCURRENCE OF HIGH SST PHENOMENA AT CENDRAWASIH BAY, PAPUA" after we sent the revised version of the manuscript.
Thank you.

Best regards,
Abd. Rasyid Jalil
(Corresponding Author)

[Kutipan teks disembunyikan]



Abdul Rasyid <abdulrasyid.fayufi@gmail.com>

GES Editor Decision

Иванов Владимир Владимирович <vladimir.ivanov@aari.ru>
Kepada: abdulrasyid fayufi <abdulrasyid.fayufi@gmail.com>

12 April 2023 pukul 00.39

Dear Dr Abd. Rasyid Jalil,

I have received the review on your revised manuscript several days ago. You will get the decision in few days.

Kind regards,
Vladimir Ivanov

От: "abdulrasyid fayufi" <abdulrasyid.fayufi@gmail.com>
Кому: "Иванов Владимир Владимирович" <vladimir.ivanov@aari.ru>
Отправленные: Вторник, 11 Апрель 2023 г 19:10:37

[Kutipan teks disembunyikan]

[Kutipan teks disembunyikan]



Abdul Rasyid <abdulrasyid.fayufi@gmail.com>

GES Editor Decision

3 pesan

Dr. Vladimir V. Ivanov <no-reply@subs.elpub.ru>
Balas Ke: "Dr. Vladimir V. Ivanov" <vladimir.ivanov@aari.ru>
Kepada: "Abd. Rasyid Jalil" <abdulrasyid.fayufi@gmail.com>

15 April 2023 pukul 01.37

Dear Abd. Rasyid Jalil:

We have reached a decision regarding your submission to GEOGRAPHY, ENVIRONMENT, SUSTAINABILITY, "CENDRAWASIH HOT POOL: STUDY OF THE HIGH-FREQUENCY OCCURRENCE OF HIGH SST PHENOMENA AT CENDRAWASIH BAY, PAPUA".

Our decision is to: accept your manuscript for publication. Congratulations and thank you very much for sending the results of your best research to GES journal!

Dr. Vladimir V. Ivanov
AARI, MSU
vladimir.ivanov@aari.ru

Editorial Office

<http://ges.rgo.ru/>

[Отказаться от рассылки](#)

Dr. Vladimir V. Ivanov <no-reply@subs.elpub.ru>
Balas Ke: "Dr. Vladimir V. Ivanov" <vladimir.ivanov@aari.ru>
Kepada: "Abd. Rasyid Jalil" <abdulrasyid.fayufi@gmail.com>

15 April 2023 pukul 01.38

Dear Abd. Rasyid Jalil:

We have reached a decision regarding your submission to GEOGRAPHY, ENVIRONMENT, SUSTAINABILITY, "CENDRAWASIH HOT POOL: STUDY OF THE HIGH-FREQUENCY OCCURRENCE OF HIGH SST PHENOMENA AT CENDRAWASIH BAY, PAPUA".

Our decision is to:

Dr. Vladimir V. Ivanov
AARI, MSU
vladimir.ivanov@aari.ru

[Kutipan teks disembunyikan]

Abdul Rasyid <abdulrasyid.fayufi@gmail.com>
Kepada: "Dr. Vladimir V. Ivanov" <vladimir.ivanov@aari.ru>

27 April 2023 pukul 00.56

Dear GES editor

Thank you very much for accepting our paper for publishing in your journal. We are waiting for the next process.

Best regards,
Abd. Rasyid Jalil
(Corresponding Author)
[Kutipan teks disembunyikan]



Abdul Rasyid <abdulrasyid.fayufi@gmail.com>

GES Editor Decision

4 pesan

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25 April 2023 pukul 15.07

Kepada: "abdulrasyid.fayufi@gmail.com" <abdulrasyid.fayufi@gmail.com>

Dear Dr. Abd. Rasyid Jalil,

We are pleased to announce that your manuscript has been approved for the Editorial Board consideration at the nearest meeting. Please, send us the final version of your manuscript without track changes along with the figures in high-resolution raster format (png or tiff) by the end of this week.

--

Best regards,
Nina Komova
Assistant editor
Geography, Environment, Sustainability
Editorial Office.
Website: <http://ges.rgo.ru/jour>

GES Journal <ges-journal@geogr.msu.ru>

2 Mei 2023 pukul 18.47

Kepada: "abdulrasyid.fayufi@gmail.com" <abdulrasyid.fayufi@gmail.com>

Cc: Abdul Malik <abdulmalik@unm.ac.id>, "aninosi@yahoo.co.id" <aninosi@yahoo.co.id>

Dear Dr. Abd. Rasyid Jalil,

Hope you are doing well. We asked you a few days ago about your manuscript files and haven't heard back. Please let us know if there's anything we can do to help.

--

Best regards,
Nina Komova
Assistant editor
Geography, Environment, Sustainability
Editorial Office.
Website: <http://ges.rgo.ru/jour>

25.04.2023, 10:07, "GES Journal" <ges-journal@geogr.msu.ru>:

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Abdul Rasyid <abdulrasyid.fayufi@gmail.com>

3 Mei 2023 pukul 01.28

Kepada: GES Journal <ges-journal@geogr.msu.ru>

Dear Nina Komova
Assistant Editor of GES

We apologize for the delay in sending the final version of manuscript (without track changes) and the figures in high resolution, because we just found your email in a spam folder.
Here is the final version of manuscript and the figures (attached)

Best regards,
Abd. Rasyid Jalil
(Corresponding Author)

[Kutipan teks disembunyikan]

--
 Dr. Abd. Rasyid Jalil
 Department of Marine Science
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7 lampiran

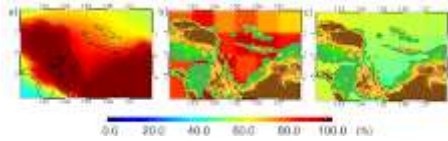


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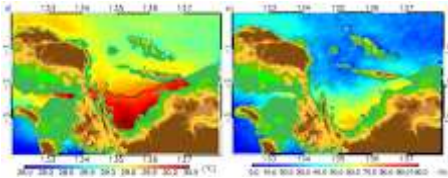


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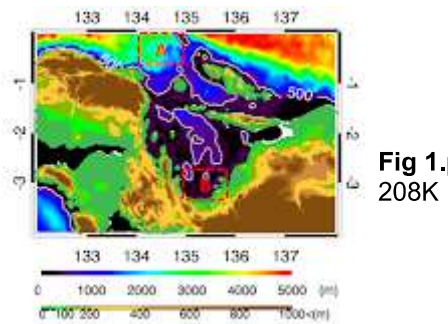


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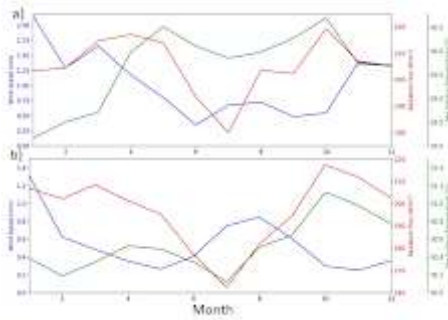


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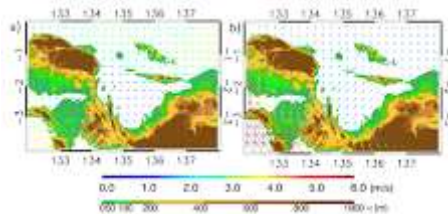


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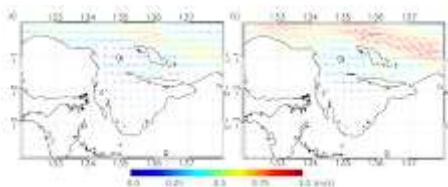


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Kepada: Abdul Rasyid <abdulrasyid.fayufi@gmail.com>

3 Mei 2023 pukul 17.22

Dear Dr. Abd. Rasyid Jalil,
Thank you for your files. Unfortunately, it is a common problem that emails end up in the spam folder.

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Best regards,
Nina Komova
Assistant editor
Geography, Environment, Sustainability
Editorial Office.
Website: <http://ges.rgo.ru/jour>

02.05.2023, 20:29, "Abdul Rasyid" <abdulrasyid.fayufi@gmail.com>:

[Kutipan teks disembunyikan]

Abd. Rasyid Jalil^{1*}, Anindya Wirasatriya^{2,3}, Fatwa Ramdani^{4,5}, Abdul Malik⁶, Puji Rahmadi⁷, Gentio Harsono⁸, Riza Yuliratno Setiawan⁹

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CENDERAWASIH HOT POOL: THE FREQUENT HIGH SST PHENOMENA AT CENDERAWASIH BAY, PAPUA

ABSTRACT. The warm pool has been used to define a water body with the characteristic of SST exceeding 28°C within a particular area and a relatively long period in an annual circle. However, there are regions with an annual mean SST measured above 30°C, and we classified them as Hot Pools because of the conditions of intense solar radiation and low wind speed. One of the Hot Pool spots was found in Indonesia, in Cenderawasih Bay. The present study examines the existence of the Cenderawasih Hot Pool using long-term observation of satellite SST data. We also used surface wind data, surface heat flux, and surface current to investigate their mechanisms. The results show that within 13 years of observation (2013-2015), SSTs in Cenderawasih Bay can reach more than 30°C with 50% occurrence within the observation period. Heat input comes from strong solar radiation, i.e., 50% of solar radiation is more than 200 W/m². The location is also dominated by low wind speed, i.e., 80% wind speed of lower than 4 m/s, which caused the low latent loss in Cenderawasih Bay. Cenderawasih Bay is fully isolated for surface currents during the dry and wet seasons since the easterly subsurface water flow does not enter the bay. The absence of strong currents prevents the mixing process, maintaining the high temperature in the surface layer. Those processes are figured out and become strong evidence to justify Cenderawasih Bay as one of the Hot Pool areas within the Indonesian seas.

KEYWORDS: Global climate; Sea Surface Temperature; Hot Pool Spot; Cenderawasih Bay.

INTRODUCTION

The western equatorial Pacific significantly influences the global climate. The warm pool, a region with average sea surface temperatures (SST) above 28°C, has a specific impact on the Earth's circulation (e.g., Wyrтки 1989; Yan et al. 1992; Clement and Seager 1999; Chongyin et al. 1999; Pierrehumbert 2000; Clement et al. 2005; Thoron et al. 2005; Herweijer et al. 2005).

On the other hand, high SST in tropical regions has attracted researchers to investigate the mechanisms since the formation of high SST requires a particular atmospheric process (e.g., Ramanathan and Collins 1991; Wallace 1992; Arking and Ziskin 1994). This process is depicted by Waliser and Graham (1993), which shows the relation between SSTs and deep convection. They utilized highly reflective cloud information from an arbitrary examination of monthly SST data with a grid spacing of 2° (produced from combined satellite observation and in situ data) and daily visible and infrared satellite image data. The highly reflective cloud increases along with the SST as it rises from 26°C to 29.5°C. In contrast, the highly reflective cloud diminishes with increased SST in the temperature from 29.5°C to 32°C. As a result, the analysis proved that several atmospheric processes impacted SSTs below and above 29.5°C.

By taking advantage of high temporal and spatial resolution SST products derived from satellite observations (i.e., daily and $\leq 25 \text{ km} \times 25 \text{ km}$), several studies (Kawamura et al. 2008; Qin et al. 2007, 2008; Qin and Kawamura 2009, 2010; Wirasatriya et al. 2015, 2016, 2017a, 2020) were able to identify high SST events (i.e., more than 30°C) in specific areas and at certain periods and define them as Hot Event (HE). In summary, they concluded that considerable daily heat gains characterize the production of HE under high solar radiation and low wind speed brought on by "remote convection" mechanisms.

Wirasatriya et al. (2015) elaborated on the climatology of HEs in the western equatorial Pacific using the SST dataset derived from satellite microwave sensors. Throughout nine years of observation (2003-2011), they discovered 71 HE cases in the western equatorial Pacific, with the majority centered on the Solomon Islands and New Guinea Island's northern coasts, which extend eastward up to 160°W. According to the climatology, the region has solar radiation of more than 200 W/m² and wind speeds of less than 4 m/s. Low wind speeds heavily influence the mechanism for HE incidence in the western equatorial Pacific. Much of the equatorial region experiences sun radiation above 200 W/m² during the HE periods. Low wind speeds minimize latent heat loss, which results in high SSTs and HEs in specific locations. Wirasatriya et al. (2015) also emphasized that high solar radiation and low wind speed are much more common during the development stage and less common during the decay stage. This study also demonstrated that a rise in the long-term mean SST in the western equatorial Pacific is correlated with an increase in the frequency of HE events. HEs were responsible for 51.5% of the SSTs >30°C in the warm pool region bounded by the 29.5 °C isotherms of the climatological SST. Thus, statistically, there is a relation between the occurrence of HE and the formation of the western Pacific warm pool. Moreover, Wirasatriya et al. (2020) demonstrated the role of HE in maintaining the warm mixed layer in the western Pacific warm pool. The frequent occurrence of HE transports heat from the surface layer to the deeper layer above the thermocline to maintain the warm mixed layer in the warm pool. Since the Pacific warm pool influences the global climate's variability through coupled ocean-atmosphere dynamics and thermodynamics (e.g., Clement and Seager 1999; Herweijer et al. 2005), this becomes an example of the contribution of high SST phenomena to regulate the global climate.

Within the Indonesian seas, the frequent appearance of high SST > 30°C has been reported by Tita et al. (2020) in Tomini Bay and by Swandiko et al. (2021) in the Malacca

Strait. Their appearances also the high solar radiation and weak wind. The morphology of the semi-enclosed waters causes the high-frequency occurrence of low wind speed by less than 4 m/s in both areas. Besides these two areas, there is no more study on the high SST phenomena within the Indonesian Seas.

The present study demonstrates the frequent high SST occurrence (more than 30°C) in Cenderawasih Bay. Cenderawasih Bay is in northern Papua Island, part of the western Pacific warm pool. It is semi-enclosed water with a deep basin, surrounded by mountain chains (Fig. 1). Cenderawasih Bay is the habitat for whale sharks (Ihsan et al. 2018). Since the area determines warm pool definition with SST climatology higher than 28°C, we call the high-frequency occurrence of SSTs higher than 30°C in the Cenderawasih Bay the Cenderawasih Hot Pool. Wirasatriya et al. (2015) found that the duration of HE occurrence in the western equatorial Pacific is no longer than two months since HE requires a typical condition of high solar radiation and low wind speed. Thus, it is interesting to understand the mechanisms of the high-frequency occurrence of high SST in the Cenderawasih Hot Pool.

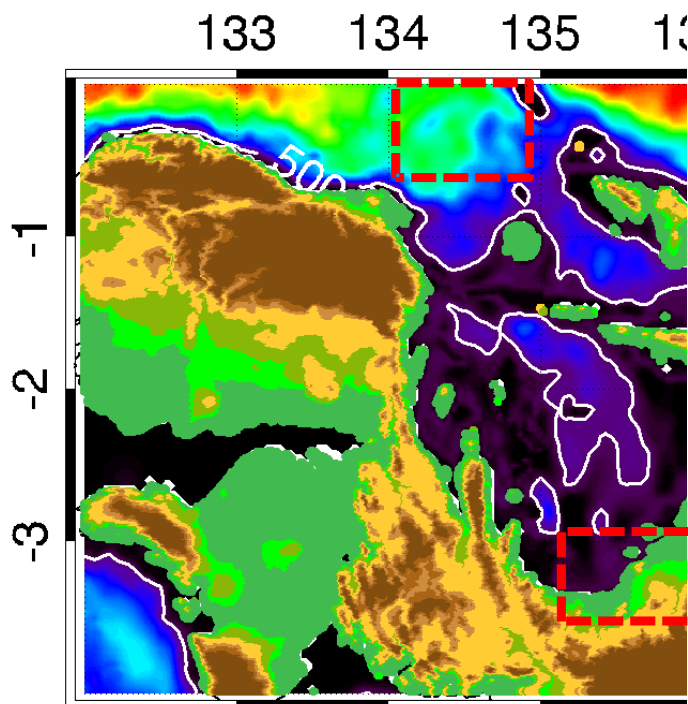


Fig. 1. Bathymetry and topography of Cenderawasih Bay. The dashed boxes A and B represent the sampling area for outside and inside Cenderawasih Bay, respectively, as shown in Fig. 5.

MATERIALS AND METHODS

We used the semi-daily 11 μm SST products from Moderate Resolution Imaging

Spectroradiometer (MODIS) Aqua Lv3 with a spatial resolution of $0.04^\circ \times 0.04^\circ$ (Esaias et al. 1998) and observation period from 2003 to 2015. The MODIS SST 11 μ m Multi-Channel SST algorithm generates SST by using brightness temperatures at 11 μ m and 12 μ m (Brown and Minnet 2009). The best accuracy was achieved by validating and testing this data against in-situ measurements (e.g., Lee et al. 2010; Qin et al. 2014; Ghanea et al. 2015).

We can obtain the Cross-Calibrated Multi-Platform (CCMP) gridded surface vector winds version 2.0 from the following website: www.remss.com, and use it to calculate surface winds. As a result, CCMP is regarded as a Level-3 ocean vector wind analysis product. They are created utilizing data from moored buoys, satellites, and model winds. The surface wind data for spatial and temporal resolutions is $0.25^\circ \times 0.25^\circ$ respectively and quarter daily. This wind product has greater accuracy than other wind reanalysis data.

We used a 0.5° grid from the Objectively Analyzed Air-Sea Fluxes project to analyze daily latent heat fluxes and solar radiation (Yu and Weller, 2007). The period of observation for latent heat flux is from 2003 to 2015, but for solar radiation is only from 2003 to 2009 due to the availability of the data. Global 30 Arc-Second Elevation data was used to obtain topography. Global 30 Arc-Second Elevation (GTOPO30) data was used to obtain topography (https://www.usgs.gov/centers/eros/science/usgs-eros-archive-digital-elevation-global-30-arc-second-elevation-gtopo30?qt-science_center_objects=0#qt-science_center_objects). ETOPO, a one-arc-minute global relief model of Earth's surface that combines land topography and ocean bathymetry, provides bathymetry (<https://www.ngdc.noaa.gov/mgg/global/>). For oceanic parameters, we used the current reanalysis data from <http://marine.copernicus.eu/> to obtain the seasonal variation of the current pattern at 100 m depth in the Cenderawasih Hot Pool.

To calculate the frequency occurrence of SST within 13 years of observation, we calculate the percentage of SST $> 30^\circ\text{C}$ occurrence in each grid. Since high SST of more than 30°C occurs during the condition of low wind speed (i.e., < 4 m/s) and high solar radiation (i.e., > 200 W/m²) (Wirasatriya et al. 2015), we calculated the percentage of low wind speed and high solar radiation in the Cenderawasih Bay. Furthermore, we also calculated the percentage of low latent heat release (< 120 W/m²) to explain how wind speed influences the variability of SST. The equation for calculating the percentage is as follows:

$$(1)$$

where P_{HSS} is percentage the percentage of high SST $> 30^\circ\text{C}$ or percentage of weak wind speed < 4 m/s or percentage of solar radiation > 200 W/m² or percentage of latent heat flux < 120 W/m² at position (x,y); p_i is the amount of SST data $> 30^\circ\text{C}$ or the amount of weak wind speed < 4 m/sec or the amount solar radiation > 200 W/m² or the amount of latent heat flux < 120 W/m² at position (x,y). Furthermore, we also analyze the climatological mean of each parameter. To create monthly and monthly climatology, all geophysical parameters are incorporated as follows: (Wirasatriya et al. 2017b)

$$(2)$$

where \bar{p} is monthly mean value or monthly climatology value at position (x,y), p_i is i^{th} value of the data at (x,y) position and time t. Next, n is the number of data in 1 month and the number of monthly data in 1 period of climatology (i.e., from 2003 to 2016 = 13 data) for monthly calculation and monthly climatology calculation, respectively. Moreover, if pixel xi is a hollow pixel, it is not included in the calculation.

RESULTS AND DISCUSSION

Cenderawasih Hot Pool

The definition of the Cenderawasih Hot Pool is described in Fig. 2, which shows the climatological mean of SST and the percentage of daily SSTs more than 30°C from 2003 to 2015 (13 years). In Cenderawasih Bay, the mean SST at the southern part of 2°S is more than 30°C, and the high SSTs are more than 50%. It means that the mean SST in Cenderawasih Bay is higher than the definition of the western Pacific warm pool which is 28°C (Wyrtki 1989), and more than 3.5 years within 2003-2015, SSTs in Cenderawasih Bay can reach more than 30°C. At the southernmost of the bay, the mean SST reaches 30.5°C, with the percentage of high SSTs being more than 70%. It has become the hottest area of Cenderawasih Bay.

In contrast, in the area outside Cenderawasih Bay, i.e., the western Pacific warm pool's offshore seas, the percentage is less than 30%. Thus, this evidence supports the definition of Cenderawasih Bay as the Cenderawasih Hot Pool. The high-frequency occurrence of high SST in Cenderawasih Bay may correspond to whale sharks' yearly appearance. Ihsan et al. (2018) indicated the tolerance of whale sharks in high SST conditions.

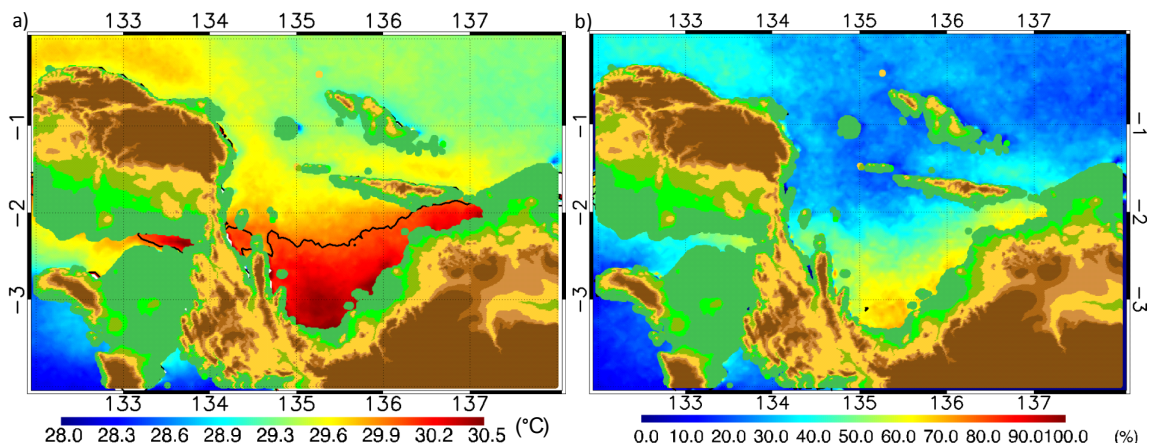


Fig. 2. Climatological mean of SST (a) and percentage of daily SSTs > 30°C during thirteen years of observation (2003-2015) (b). The black contour is 30°C.

Atmospheric Aspect of Cenderawasih Hot Pool

Fig. 3 clearly shows that low wind speed dominates Cenderawasih Bay, i.e., 80% wind speed < 4 m/s occurred during 2003-2015. This causes low latent loss in Cenderawasih Bay. For solar radiation, the percentage of high solar radiation inside and outside Cenderawasih Bay is almost similar i.e., nearly 50%, indicating that high solar radiation is not dominant in both areas. This means that, although high solar radiation does not frequently occur, the absence of strong wind speed may maintain the latent heat loss to keep the SST in Cenderawasih Bay higher than 30°C.

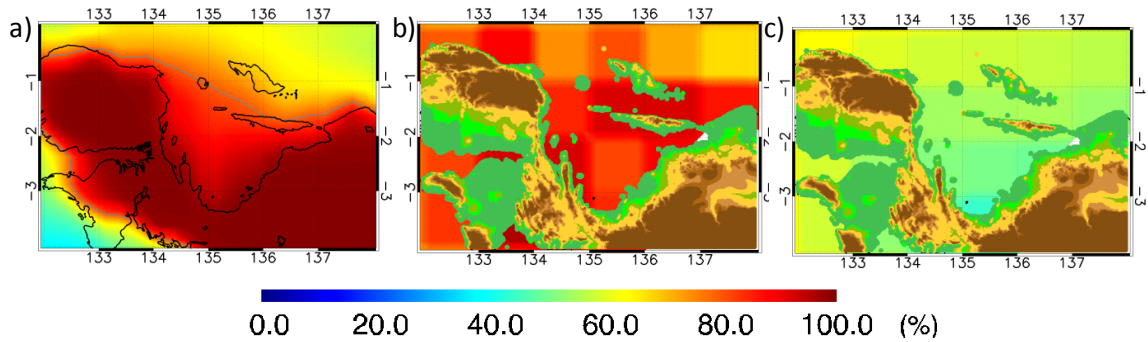


Fig. 3. The percentage of wind speed less than four m/s (a), latent heat release less than 120 W/m² (b), and solar radiation more than 200 W/m² (c).

To investigate how low wind speed occurs in the study area, we show the monthly climatology map of surface wind during summer and winter (Fig. 4). For both seasons, Cenderawasih Bay is protected from high wind speed due to high mountain chains in the western, southern, and eastern parts. In the northern part, small islands at the mouth of the bay prevent the strong wind speed from entering the bay. Thus, the role of topography is crucial for the occurrence of constant high SSTs in Cenderawasih Bay. The same tendencies are also found at the Tomini Bay and Malacca Strait, as indicated by Tita et al. (2020) and Swandiko et al. (2021).

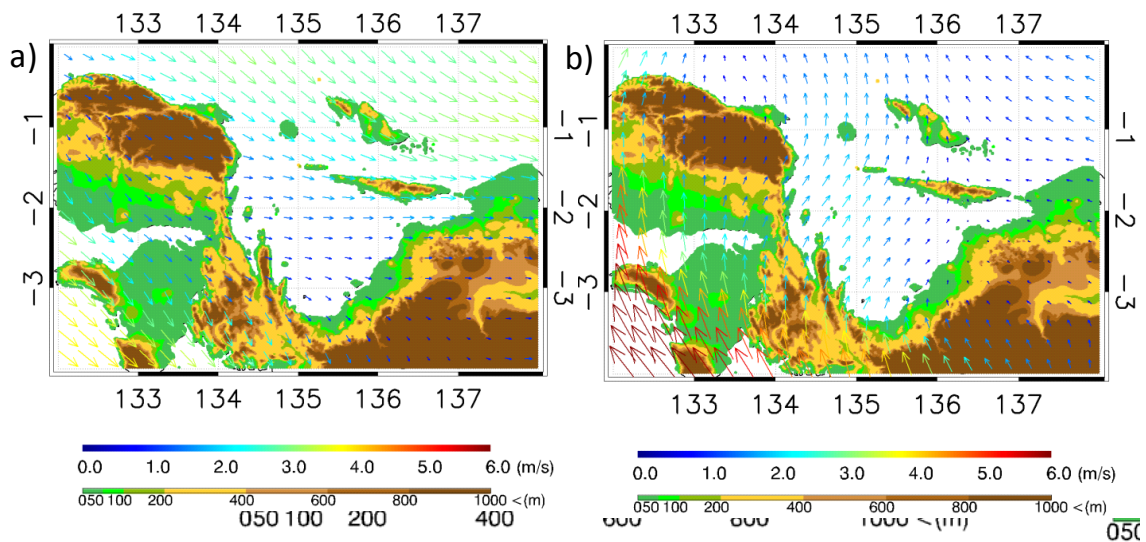


Fig. 4. Monthly climatology of surface wind speed during a) wet season (January) and b) dry season (August)

The influence of wind speed and solar radiation on the SST variation is further examined by plotting the monthly climatology of wind speed, solar radiation, and SST at areas A and B (Fig. 1), representing the areas outside and inside Cenderawasih Bay. Seasonal variations of SSTs are observed in both areas. The minimum SSTs occur during the wet and dry seasons, while the maximum SSTs during the transition season. This seasonal SST variability is similar to the other areas in the Indonesian seas, such as the Java Sea (Wirasatriya et al. 2018), Maluku Sea (Wirasatriya et al. 2019), Halmahera Sea (Setiawan et al. 2019), etc.

For the area outside Cenderawasih Bay, SST ranges from 28.2°C to 30.6°C. It is seen that the variability of SST is controlled by wind speed and solar radiation. From January to June, when wind speed decreases from 2.2 m/s to 0.3 m/s and solar radiation is more

than 200 W/m^2 , SST increases from 28.2°C to 30.5°C . From June to October, wind speed is lower than 1 m/s . In the absence of strong wind, the low SST in July is caused by the decrease of solar radiation to a minimum.

For the area inside Cenderawasih Bay, persistent high SST is observed as the SST ranges from 30.2°C to 31.1°C . It may be due to the low wind speed of fewer than 1 m/s for years. Thus, the variability of SST is seen as ruled by solar radiation. The minimum solar radiation causes the minimum SST in July. However, the absence of strong wind prevents latent heat loss that maintains the high SST inside Cenderawasih Bay.

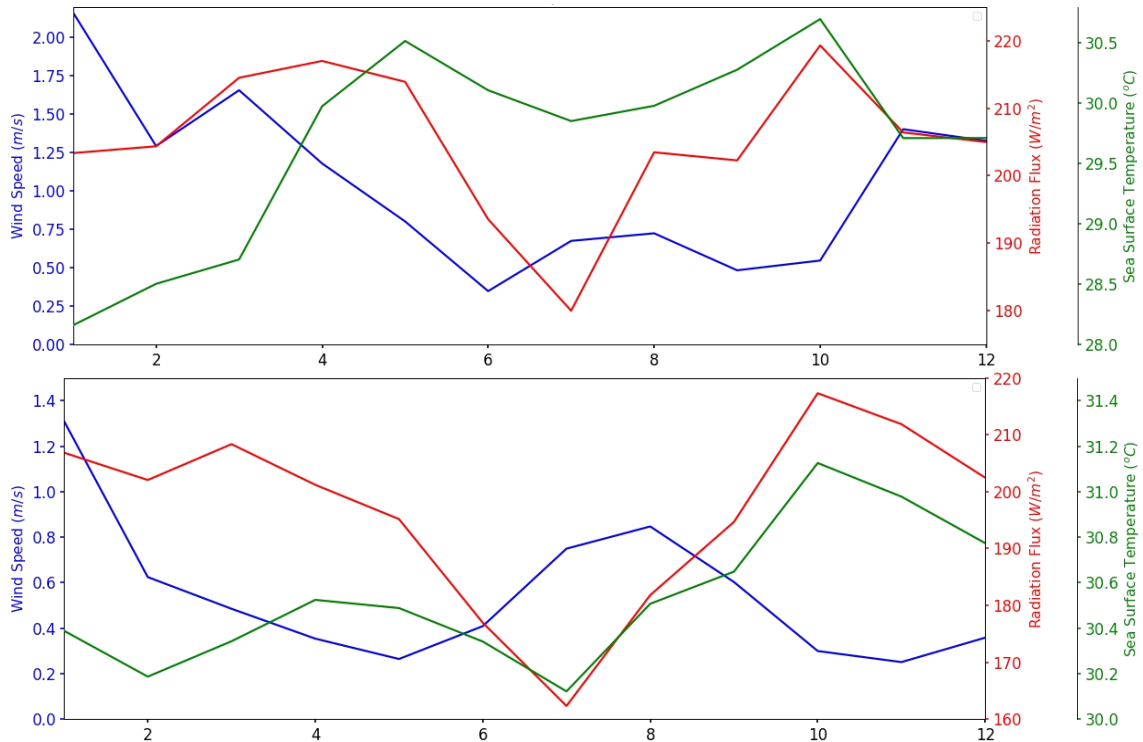


Fig. 5. Monthly climatology of SST, wind speed, and solar radiation at a) area outside and b) area inside Cenderawasih Bay, as shown in Fig. 1.

The oceanic aspect of Cenderawasih Hot Pool

To investigate the oceanic aspect of the Cenderawasih Hot Pool, we plotted the current patterns at 100 m depth shown in Fig. 6. A northwestward current has been identified north of Cenderawasih Bay. During the wet season (January), the speed of this current ranges from 0.5 m/s to 0.7 m/s . During the dry season (August), when the northwesterly wind disappears (Fig. 4b), the speed of the northwestward current increases. This northwestward current is known as the New Guinea Coastal Under Current (NGCUC). NGCUC is a permanent current feature at 200 m depth regardless of the wind reversals (Tsuchiya et al. 1989). Inside Cenderawasih Bay, strong currents are absent in this area, preventing the mixing process from maintaining the high temperature in the surface layer. During the dry and wet seasons, Cenderawasih Bay is fully isolated since the westward subsurface water flow does not enter Cenderawasih Bay. Thus, this isolated basin causes the Cenderawasih Bay area to be influenced mainly by the air-sea interaction process mentioned in the previous section creating Cenderawasih Hot Pool.

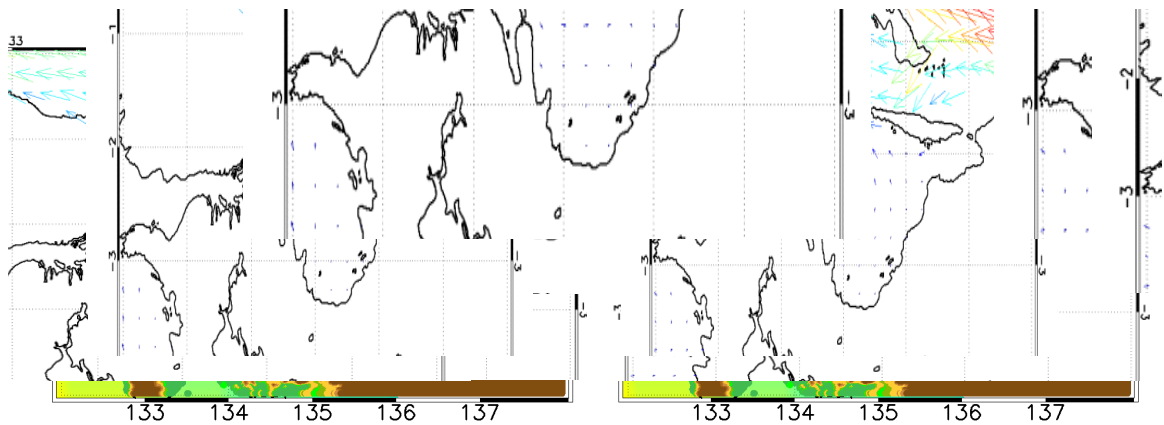


Fig. 6. Monthly climatology of current at 100 m depth in January (a) and August (b)

CONCLUSIONS

The hot Pool is used to categorize high SST episodes ($>30^{\circ}\text{C}$) in particular areas and during specific periods (relatively long periods). Since the "warm pool" has been used to define an area with the annual average SST distributed below 30°C in dominant, therefore "hot pool" was taken to represent the area with SST dominated above 30°C .

Under high solar radiation and low wind speed, this event is characterized by considerable daily heat gains. The constant high SST occurrence in Cenderawasih Bay, which is more than 50% of 13 years, SSTs can reach more than 30°C and are dominated by the condition of low wind speed, i.e., 80% wind speed is less than m/s along the years, also 50% solar radiation measured is more than 200 W/m^2 . The current pattern at 100 m depth also shows that NGCUC does not enter Cenderawasih Bay. Thus, this indicates that Cenderawasih Bay is an isolated water. Those are the solid evidence for defining Cenderawasih Bay as the "Cenderawasih Hot Pool."

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REFERENCES

- Arking A. and Ziskin D. (1994). Relationship between clouds and sea surface temperatures in the western tropical Pacific. *Journal of Climate*, 7(6), 988-1000. [https://doi.org/10.1175/1520-0442\(1994\)007%3C0988:RBCASS%3E2.0.CO;2](https://doi.org/10.1175/1520-0442(1994)007%3C0988:RBCASS%3E2.0.CO;2)
- Brown O.B. and Minnett P.J. (2009). MODIS infrared sea surface temperature algorithm theoretical basis document, ver 2.0. [online]. Available at: https://modis.gsfc.nasa.gov/data/atbd/atbd_mod25.pdf. [Accessed 25 October 2022].
- Chongyin L., Mingquan M., and Guangqing Z. (1999). The variation of warm pool in the equatorial western Pacific and its impacts on climate. *Advances in Atmospheric Sciences*, 16(3), 378-394. <https://doi.org/10.1007/s00376-999-0017-0>
- Clement A. and Seager R. (1999). Climate and the tropical oceans. *Journal of Climate*, 12(12), 3383-3401. [https://doi.org/10.1175/1520-0442\(1999\)012%3C3383:CATTO%3E2.0.CO;2](https://doi.org/10.1175/1520-0442(1999)012%3C3383:CATTO%3E2.0.CO;2)
- Clement A.C., Seager R., and Murtugudde R. (2005). Why are there tropical warm pools? *Journal of Climate*, 18(24), 5294-5311. <https://doi.org/10.1175/JCLI3582.1>

- Esaias, W.E., Abbott M.R., Barton I., Brown O.B., Campbell J.W., Carder K.L., ... and Minnett P.J. (1998). An overview of MODIS capabilities for ocean science observations, *IEEE Transactions on Geoscience and Remote Sensing*, 36(4), 1250-1265. <https://doi.org/10.1109/36.701076>
- Ghanea M., Moradi M., Kabiri K., and Mehdinia A. (2015). Investigation and validation of MODIS SST in the northern Persian Gulf, *Advances in Space Research*, 57(1), 127-136. <https://doi.org/10.1016/j.asr.2015.10.040>
- Harweijer C., Seager R., Winton M., and Clement A.M.Y. (2005) Why ocean heat transport warms the global mean climate. *Tellus*, 57(4), 662-675. <https://doi.org/10.3402/tellusa.v57i4.14708>
- Hosoda K. (2013). Empirical method of diurnal correction for estimating sea surface temperature at dawn and noon. *Journal of Oceanography*, 69, 631-646. <https://doi.org/10.1007/s10872-013-0194-4>
- Ihsan E.N., Enita S.Y., and Wirasatriya, A. (2018). Oceanographic factors in fishing ground location of Anchovy at Teluk Cenderawasih National Park, West Papua: Are these factors have an effect of Whale Sharks appearance frequencies? In IOP Conference Series: Earth and Environmental Science, 116(1), p. 012017. <https://doi.org/10.1088/1755-1315/116/1/012017>
- Kawamura H., Qin H., and Ando, K. (2008). In-situ diurnal sea surface temperature variations and near-surface thermal structure in the tropical hot event of the Indo-Pacific warm pool. *Journal of Oceanography*, 64, 847-857. <https://doi.org/10.1007/s10872-008-0070-9>
- Ming-An, L., Tzeng, M. T., Hosoda, K., Sakaida, F., Kawamura, H., Shieh, W. J., Yang Y., and Chang, Y. (2010). Validation of JAXA/MODIS sea surface temperature in water around Taiwan using the Terra and Aqua satellites. *TAO: Terrestrial, Atmospheric and Oceanic Sciences*, 21(4), 727-736. [https://doi.org/10.3319/TAO.2009.09.07.01\(Oc\)](https://doi.org/10.3319/TAO.2009.09.07.01(Oc))
- Pierrehumbert R.T. (2000). Climate change and the tropical Pacific: The sleeping dragon wakes. *Proceedings of the National Academy of Sciences*, 97(4), 1355-1358. <https://doi.org/10.1073/pnas.97.4.1355>
- Qin H., Kawamura H., and Kawai Y. (2007). Detection of Hot Event in the equatorial Indo-Pacific warm pool using advanced satellite sea surface temperature, solar radiation, and wind speed. *Journal of Geophysical Research: Oceans*, 112 (C7). <https://doi.org/10.1029/2006JC003969>
- Qin H., Kawamura H., Sakaida F., and Ando K. (2008). A case study of the tropical hot event in November 2006 (HE0611) using a geostationary meteorological satellite and the TAO/TRITON mooring array. *Journal of Geophysical Research: Oceans*, 113(C8). <https://doi.org/10.1029/2007JC004640>
- Qin H. and Kawamura H. (2009). Atmosphere response to a hot SST event in November 2006 as observed by AIRS instrument. *Advances in space research*, 44(3), 395-400. <https://doi.org/10.1016/j.asr.2009.03.003>
- Qin H. and Kawamura H. (2010). Air-sea interaction throughout the troposphere over a very high sea surface temperature. *Geophysical research letters*, 37(1),1-4. <https://doi.org/10.1029/2009GL041685>
- Qin H., Chen G., Wang W., Wang D., and Zeng L. (2014). Validation and application of MODIS-derived SST in the South China Sea. *International journal of remote sensing*, 35(11-12), 4315-4328. <https://doi.org/10.1080/01431161.2014.916439>
- Ramanathan V. and Collins W. (1991). Thermodynamic regulation of ocean warming by cirrus clouds deduced from observations of the 1987 El Niño. *Nature* 351 (6321), 27-32. <https://doi.org/10.1038/351027a0>
- Setiawan R.Y., Wirasatriya A., Hernawan U., Leung S., and Iskandar I. (2020).

Spatio-temporal variability of surface chlorophyll-*a* in the Halmahera Sea and its relation to ENSO and the Indian Ocean Dipole. *International Journal of Remote Sensing*, 41(1), 284-299. <https://doi.org/10.1080/01431161.2019.1641244>

Swandiko M., Wirasatriya A., Marwoto J., Muslim, Indrayanti E., Subardjo P., Ismunarti D.H. (2021). Studi persistensi suhu permukaan laut tinggi (>30°C) di perairan Selat Malaka. *Buletin Oseanografi Marina*, 10(2), 162-170. [online]. Available at: <https://ejournal.undip.ac.id/index.php/buloma/article/download/31554/19352>. [Accessed 25 October 2022] (in Indonesian).

De Garidel-Thoron T., Rosenthal Y., Bassinot F., and Beaufort, L. (2005). Stable sea surface temperatures in the western Pacific warm pool over the past 1.75 million years. *Nature*, 433(7023), 294-298. <https://doi.org/10.1038/nature03189>

Tita A.D.C., Wirasatriya A., Sugianto D.N., Maslukah L., Handoyo G., Helmi M., and Avianto, P. (2020). Persistence of high sea surface temperature (> 30°C) in Tomini Bay. In *IOP Conference Series: Earth and Environmental Science*, 530(1), p. 012038. <https://doi.org/10.1088/1755-1315/530/1/012038>

Waliser D.E. and Graham N.E. (1993). Convective cloud systems and warm-pool sea-surface temperatures: Coupled interactions and self-regulation. *Journal of Geophysical Research: Atmospheres*, 98(D7), 12881-12893. <https://doi.org/10.1029/93JD00872>

Wallace J.M. (1992). Effect of deep convection on the regulation of tropical sea surface temperature. *Nature*, 357(6375), 230-231. <https://doi.org/10.1038/357230a0>

Wirasatriya A., Kawamura H., Shimada T., and Hosoda K. (2015). Climatology of hot events in the western equatorial Pacific. *Journal of Oceanography*, 71, 77-90. <https://doi.org/10.1007/s10872-014-0263-3>

Wirasatriya A., Kawamura H., Shimada T., Hosoda K. (2016). Atmospheric structure favoring high sea surface temperatures in the western equatorial Pacific. *Journal of Geophysical Research*, 121(19), 11-368. <https://doi.org/10.1002/2016JD025268>

Wirasatriya A., Sugianto D.N., and Helmi M. (2017). The Influence of Madden Julian Oscillation on the Formation of the hot event in the western equatorial Pacific. In *IOP Conference Series: Earth and Environmental Science*, 55(1), p. 012006. <https://doi.org/10.1088/1755-1315/55/1/012006>

Wirasatriya, A., Setiawan R.Y., and Subardjo P. (2017). The effect of ENSO on the variability of chlorophyll-*a* and sea surface temperature in the Maluku Sea, *IEEE Journal of Selected Topics in Applied Earth Observations and Remote Sensing*, 10(12), 5513-5518. <https://doi.org/10.1109/JSTARS.2017.2745207>

Wirasatriya, A., Prasetyawan I.B., Triyono C.D., Muslim, and Maslukah L. (2018). Effect of ENSO on the variability of SST and chlorophyll-*a* in Java Sea. In *IOP Conference Series: Earth and Environmental Science*, 116(1), p 012063. <https://doi.org/10.1088/1755-1315/116/1/012063>

Wirasatriya A., Sugianto D.N., Helmi M., Setiawan R.Y., and Koch M. (2019). Distinct characteristics of SST variabilities in the Sulawesi Sea and the northern part of the Maluku Sea during the southeast monsoon. *IEEE Journal of Selected Topics in Applied Earth Observations and Remote Sensing*, 12(6), 1763-1770. <https://doi.org/10.1109/JSTARS.2019.2913739>

Wirasatriya A., Kawamura H., Helmi M., Sugianto D.N., Shimada T., Hosoda K., Handoyo G., Putra Y.D.G., and Koch M. (2020). Thermal structure of Hot Events and their possible role in maintaining the warm isothermal layer in the Western Pacific warm pool. *Ocean Dynamics*, 70, 771-786. <https://doi.org/10.1007/s10236-020-01362-8>

Wyrtki K. (1989). Some thoughts about the West Pacific warm pool. In *Proceedings of the Western Pacific International Meeting and Workshop on TOGA COARE*, Nouméa,

New Caledonia: 99-109 [online]. Available at: https://horizon.documentation.ird.fr/exl-doc/pleins_textes/doc34-08/30195.pdf#page=114 (Accessed 26 October 2022)

Yan X.H., Ho C.R., Zheng Q., and Klemas V. (1992). Temperature and size variabilities of the western Pacific warm pool. *Science*, 258 (5088), 1643-1645. <https://doi.org/10.1126/science.258.5088.1643>

Yu L. and Weller R.A. (2007). Objectively Analyzed air-sea heat fluxes for the global ice-free oceans (1981–2005). *Bulletin of the American Meteorological Society*, 88(4), 527-540. <https://doi.org/10.1175/BAMS-88-4-527>



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CENDERAWASIH HOT POOL: THE FREQUENT HIGH SEA SURFACE TEMPERATURE PHENOMENA AT CENDERAWASIH BAY, PAPUA

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CENDERAWASIH HOT POOL: THE FREQUENT HIGH SEA SURFACE TEMPERATURE PHENOMENA AT CENDERAWASIH BAY, PAPUA

ABSTRACT. The term “warm pool” ~~warm pool has been used~~ refers to define a water body of water with the characteristic of SST exceeding 28°C within a particular area and a relatively long period in an annual circle. However, there are regions with an annual mean SST measured above 30°C, and we classified them as ~~h~~Hot ~~p~~Pools because of the conditions of intense solar radiation and low wind speed. One of the Hot Pool spots was found in Indonesia, in Cenderawasih Bay. The present study examines the existence of the Cenderawasih Hot Pool using long-term observation of satellite SST data. ~~In order to learn more about their mechanisms, we also analyzed surface wind, surface heat flux, and surface current data. We also used surface wind data, surface heat flux, and surface current to investigate their mechanisms.~~ The results show that ~~within 13 years of observation (2013–2015), SSTs in Cenderawasih Bay can reach more than 30°C with~~ have a 50% chance of exceeding 30°C ~~occurrence within the observation period~~ within the 13 years of study (2013–2015). Heat input comes from strong solar radiation, i.e., 50% of solar radiation is more than 200 W/m². The location is also dominated by low wind speed, i.e., 80% wind speed of lower than 4 m/s, which caused the low latent loss in Cenderawasih Bay. Cenderawasih Bay is fully ~~isolated for~~ separated from surface currents during the dry and wet seasons since the easterly subsurface water flow does not enter the bay. The absence of strong currents prevents the mixing process, maintaining the high temperature in the

surface layer. Those processes are ~~figured out~~ discovered and ~~they serve as compelling become strong~~ evidence to ~~justify support~~ Cenderawasih Bay as one of the Hot Pool areas within the Indonesian seas.

KEYWORDS: ~~g~~Global climate; ~~s~~sea ~~s~~Surface ~~t~~Temperature; ~~H~~Hot ~~P~~Pool ~~s~~Spot; Cenderawasih Bay.–

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INTRODUCTION

The western equatorial Pacific ~~has a big impact on significantly influences~~ the global climate. The ~~warm pool, a region with average sea surface temperatures (SST) above 28°C, has a specific impact on the Earth's circulation is specifically affected by the warm pool, an area with average sea surface temperatures (SST) exceeding 28°C~~ (e.g., Wyrki 1989; Yan et al. 1992; Clement and Seager 1999; Chongyin et al. 1999; Pierrehumbert 2000; Clement et al. 2005; Thoron et al. 2005; Herweijer et al. 2005).–

On the other hand, high SST in tropical regions has attracted researchers to investigate the mechanisms since the formation of high SST requires a particular atmospheric process (e.g., Ramanathan and Collins 1991; Wallace 1992; Arking and Ziskin 1994). This process is depicted by (Waliser and Graham (1993), which shows the relation between SSTs and deep convection. –They utilized highly reflective cloud information from an arbitrary examination of monthly SST data with a grid spacing of 2° –(produced from combined satellite observation and in situ data) and daily visible and infrared satellite image data. The highly reflective cloud increases along with the SST as it rises from 26°C to 29.5°C. In contrast, the highly reflective cloud diminishes with increased SST in the temperature from 29.5°C to 32°C. As a result, the analysis proved that several atmospheric processes impacted SSTs below and above 29.5°C.

By taking advantage of high temporal and spatial resolution SST products derived from satellite observations (i.e., daily and $\leq 25 \text{ km} \times 25 \text{ km}$), several studies (Kawamura et al. 2008; Qin et al. 2007, 2008; Qin and Kawamura 2009, 2010; Wirasatriya et al. 2015, 2016, 2017a, 2020) were able to identify high SST events (i.e., more than 30°C) in specific areas and at certain periods and define them as Hot Event (HE). In summary, they concluded that considerable daily heat gains characterize the production of HE under high solar radiation and low wind speed brought on by "remote convection" mechanisms.

(Wirasatriya et al. (2015) elaborated on the climatology of HEs in the western equatorial Pacific using the SST dataset derived from satellite microwave sensors. Throughout nine years of observation (2003-2011), they discovered 71 HE cases in the western equatorial Pacific, with the majority centered on the Solomon Islands and New Guinea Island's northern coasts, which extend eastward up to 160°W. According to the climatology, the region has solar radiation of more than 200 W/m² and wind speeds of less than 4 m/s. Low wind speeds heavily influence the mechanism for HE incidence in the western equatorial Pacific. Much of the equatorial region experiences sun radiation above 200 W/m² during the HE periods. Low wind speeds minimize latent heat loss, which results in high SSTs and HEs in specific locations. (Wirasatriya et al. (2015) also emphasized that high solar radiation and low wind speed are much more common during the development stage and less common during the decay stage. This study also demonstrated that a rise in

the long-term mean SST in the western equatorial Pacific is correlated with an increase in the frequency of HE events. HEs were responsible for 51.5% of the SSTs $>30^{\circ}\text{C}$ in the warm pool region bounded by the 29.5°C isotherms of the climatological SST. Thus, statistically, there is a relation between the occurrence of HE and the formation of the western Pacific warm pool. Moreover, (Wirasatriya et al. (2020) demonstrated the role of HE in maintaining the warm mixed layer in the western Pacific warm pool. The frequent occurrence of HE transports heat from the surface layer to the deeper layer above the thermocline to maintain the warm mixed layer in the warm pool. Since the Pacific warm pool influences the global climate's variability through coupled ocean-atmosphere dynamics and thermodynamics (e.g., Clement and Seager 1999; Herweijer et al. 2005), this becomes an example of the contribution of high SST phenomena to regulate the global climate.

Within the Indonesian seas, the frequent appearance of high SST $> 30^{\circ}\text{C}$ has been reported by (Tita et al. (2020) in Tomini Bay and by (Swandiko et al. (2021) in the Malacca Strait. Their appearances also the high solar radiation and weak wind. The morphology of the semi-enclosed waters causes the high-frequency occurrence of low wind speed by less than 4 m/s in both areas. Besides these two areas, there is no more study on the high SST phenomena within the Indonesian Seas.

The present study demonstrates the frequent high SST occurrence (more than 30°C) in Cenderawasih Bay. Cenderawasih Bay is in northern Papua Island, part of the western Pacific warm pool. It is semi-enclosed water with a deep basin, surrounded by mountain chains (Fig. 1). Cenderawasih Bay is the habitat for whale sharks (Ihsan et al. 2018). Since the area determines warm pool definition with SST climatology higher than 28°C , we call the high-frequency occurrence of SSTs higher than 30°C in the Cenderawasih Bay the Cenderawasih Hot Pool. (Wirasatriya et al. (2015) found that the duration of HE occurrence in the western equatorial Pacific is no longer than two months since HE requires a typical condition of high solar radiation and low wind speed. Thus, it is interesting to understand the mechanisms of the high-frequency occurrence of high SST in the Cenderawasih Hot Pool.

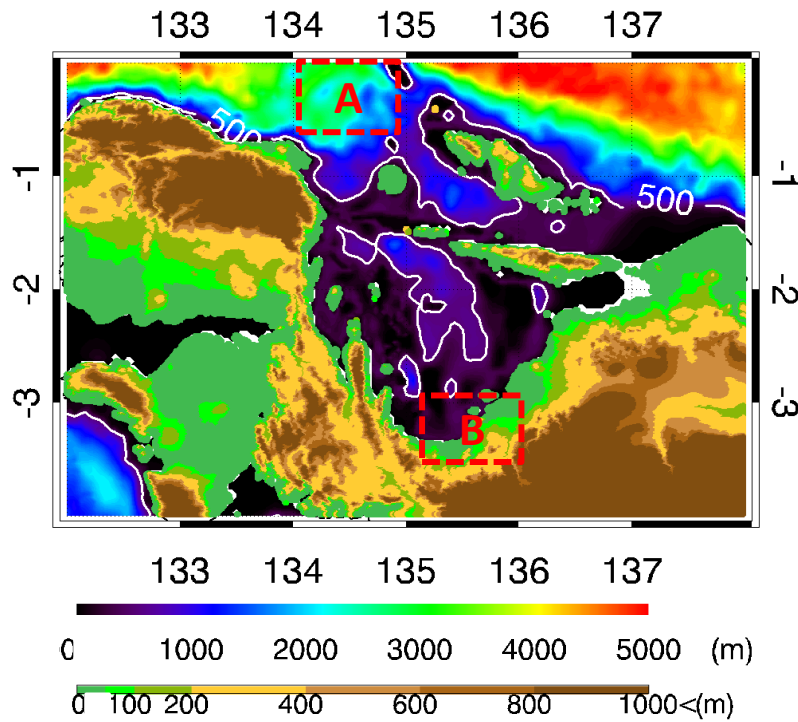


Fig. 1. Bathymetry and topography of Cenderawasih Bay. The dashed boxes A and B represent the sampling area for outside and inside Cenderawasih Bay, respectively, as shown in Fig. 5.

MATERIALS AND METHODS

We used the semi-daily 11 μm SST products from Moderate Resolution Imaging Spectroradiometer (MODIS) Aqua Lv3 with a spatial resolution of $0.04^\circ \times 0.04^\circ$ (Esaias et al., 1998) and observation period from 2003 to 2015. The MODIS SST 11 μm the Multi-Channel SST algorithm generates SST by using brightness temperatures at 11 μm and 12 μm (Brown and Minnet 2009). The best accuracy was achieved by validating and testing this data against in-situ measurements (e.g., Lee et al. 2010; Qin et al. 2014; Ghanea et al. 2015).

We can obtain the Cross-Calibrated Multi-Platform (CCMP) gridded surface vector winds version 2.0 from the following website: www.remss.com, and use it to calculate surface winds. As a result, CCMP is regarded as a Level-3 ocean vector wind analysis product. They are created utilizing data from moored buoys, satellites, and model winds. The surface wind data for spatial and temporal resolutions is $0.25^\circ \times 0.25^\circ$ respectively and quarter daily. This wind product has greater accuracy than other wind reanalysis data.

We used a 0.5° grid from the Objectively Analyzed Air-Sea Fluxes project to analyze daily latent heat fluxes and solar radiation (Yu and Weller, 2007). The period of observation for latent heat flux is from 2003 to 2015, but for solar radiation is only from 2003 to 2009

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due to the availability of the data. Global 30 Arc-Second Elevation data was used to obtain topography. Global 30 Arc-Second Elevation (GTOPO30) data was used to obtain topography ¹. (https://www.usgs.gov/centers/eros/science/usgs-eros-archive-digital-elevation-global-30-arc-second-elevation-gtopo30?qt-science_center_objects=0#qt-science_center_objects). ETOPO, a one-arc-minute global relief model of Earth's surface that combines land topography and ocean bathymetry, provides bathymetry ². (<https://www.ngdc.noaa.gov/mgg/global/>). For oceanic parameters, we used the current reanalysis data from <http://marine.copernicus.eu/> to obtain the seasonal variation of the current pattern at 100 m depth in the Cenderawasih Hot Pool.

To calculate the frequency occurrence of SST within 13 years of observation, we calculate the percentage of SST > 30°C occurrence in each grid. Since high SST of more than 30°C occurs during the condition of low wind speed (i.e., < 4 m/s) and high solar radiation (i.e., > 200 W/m²) (Wirasatriya et al. 2015), we calculated the percentage of low wind speed and high solar radiation in the Cenderawasih Bay. Furthermore, we also calculated the percentage of low latent heat release (<120 W/m²) to explain how wind speed influences the variability of SST. The equation for calculating the percentage is as follows:

$$\%(x, y) = \frac{1}{n} \sum_{i=1}^n p_i(x, y) \times 100\% \quad (1)$$

where $\%(x, y)$ is percentage the percentage of high SST > 30°C or percentage of weak wind speed < 4 m/s or percentage of solar radiation > 200 W/m² or percentage of latent heat flux < 120W/m² at position (x,y); p_i is the amount of SST data > 30°C or the amount of weak wind speed < 4 m/sec or the amount solar radiation > 200 W/m² or the amount of latent heat flux < 120W/m² at position (x,y). Furthermore, we also analyze the climatological mean of each parameter. To create monthly and monthly climatology, all geophysical parameters are incorporated as follows: (Wirasatriya et al. 2017b):

$$\bar{X}(x, y) = \frac{1}{n} \sum_{i=1}^n x_i(x, y, t) \quad (2)$$

where $\bar{X}(x, y)$ is monthly mean value or monthly climatology value at position (x,y), $x_i(x, y, t)$ is i^{th} value of the data at (x,y) position and time t. Next, n is the number of data in 1 month and the number of monthly data in 1 period of climatology (i.e., from 2003 to 2016 = 13 data) for monthly calculation and monthly climatology calculation, respectively. Moreover, if pixel x_i is a hollow pixel, it is not included in the calculation.

RESULTS AND DISCUSSION

Cenderawasih Hot Pool

The definition of the Cenderawasih Hot Pool is described in Fig. 2, which shows the climatological mean of SST and the percentage of daily SSTs more than 30°C from 2003 to 2015 (13 years). In Cenderawasih Bay, the mean SST at the southern part of 2°S is more than 30°C, and the high SSTs are more than 50%. It means that the mean SST in Cenderawasih Bay is higher than the definition of the western Pacific warm pool which is 28°C (Wyrtki 1989), and more than 3.5 years within 2003-2015, SSTs in Cenderawasih Bay can reach more than 30°C. At the southernmost of the bay, the mean SST reaches 30.5°C, with the percentage of high SSTs being more than 70%. It has become the hottest area of Cenderawasih Bay.

In contrast, in the area outside Cenderawasih Bay, i.e., the western Pacific warm pool's

¹ https://www.usgs.gov/centers/eros/science/usgs-eros-archive-digital-elevation-global-30-arc-second-elevation-gtopo30?qt-science_center_objects=0#qt-science_center_objects

² <https://www.ngdc.noaa.gov/mgg/global/>

offshore seas, the percentage is less than 30%. Thus, this evidence supports the definition of Cenderawasih Bay as the Cenderawasih Hot Pool. The high-frequency occurrence of high SST in Cenderawasih Bay may correspond to whale sharks' yearly appearance. (Ihsan et al. (2018) indicated the tolerance of whale sharks in high SST conditions.

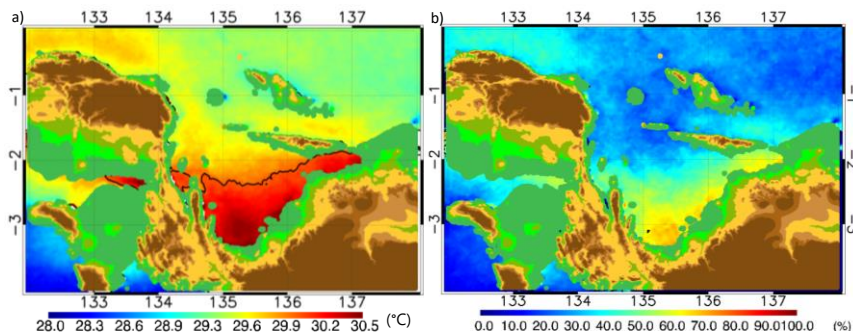


Fig. 2. Climatological mean of SST (a) and percentage of daily SSTs > 30°C during thirteen years of observation (2003-2015) (b). The black contour is 30°C.

Atmospheric Aspect of Cenderawasih Hot Pool

Fig. 3 clearly shows that low wind speed dominates Cenderawasih Bay, i.e., 80% wind speed < 4 m/s occurred during 2003-2015. This causes low latent loss in Cenderawasih Bay. For solar radiation, the percentage of high solar radiation inside and outside Cenderawasih Bay is almost similar i.e., nearly 50%, indicating that high solar radiation is not dominant in both areas. This means that, although high solar radiation does not frequently occur, the absence of strong wind speed may maintain the latent heat loss to keep the SST in Cenderawasih Bay higher than 30°C.

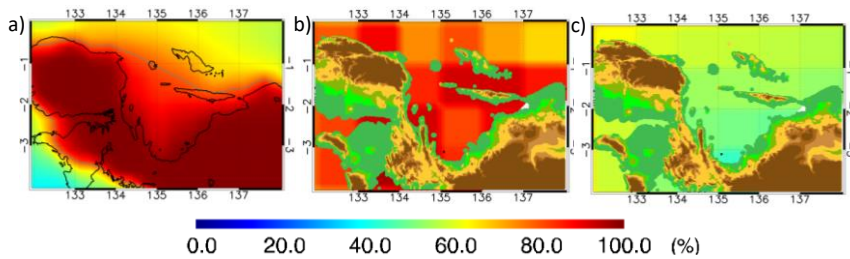


Fig. 3. The percentage of wind speed less than four m/s (a), latent heat release less than 120 W/m² (b), and solar radiation more than 200 W/m² (c)-

To investigate how low wind speed occurs in the study area, we show the monthly climatology map of surface wind during summer and winter (Fig. 4). For both seasons, Cenderawasih Bay is protected from high wind speed due to high mountain chains in the western, southern, and eastern parts. In the northern part, small islands at the mouth of the bay prevent the strong wind speed from entering the bay. Thus, the role of topography is crucial for the occurrence of constant high SSTs in Cenderawasih Bay. The same tendencies are also found at the Tomini Bay and Malacca Strait, as indicated by (Tita et al. (2020) and (Swandiko et al. (2021).

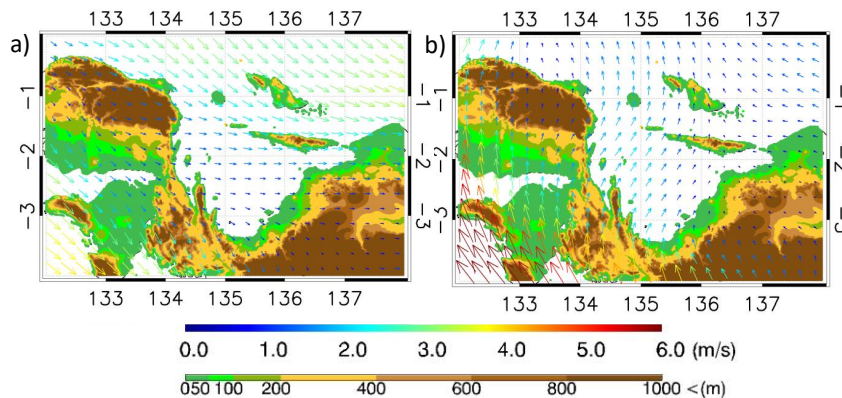


Fig. 4. Monthly climatology of surface wind speed during a) wet season (January) and b) dry season (August)

The influence of wind speed and solar radiation on the SST variation is further examined by plotting the monthly climatology of wind speed, solar radiation, and SST at areas A and B (Fig. 1), representing the areas outside and inside Cenderawasih Bay. Seasonal variations of SSTs are observed in both areas. The minimum SSTs occur during the wet and dry seasons, while the maximum SSTs during the transition season. This seasonal SST variability is similar to the other areas in the Indonesian seas, such as the Java Sea (Wirasatriya et al. 2018), Maluku Sea (Wirasatriya et al. 2019), Halmahera Sea (Setiawan et al. 2019), etc.

For the area outside Cenderawasih Bay, SST ranges from 28.2°C to 30.6°C. It is seen that the variability of SST is controlled by wind speed and solar radiation. From January to

June, when wind speed decreases from 2.2 m/s to 0.3 m/s and solar radiation is more than 200 W/m², SST increases from 28.2°C to 30.5°C. From June to October, wind speed is lower than 1 m/s. In the absence of strong wind, the low SST in July is caused by the decrease of solar radiation to a minimum.

For the area inside Cenderawasih Bay, persistent high SST is observed as the SST ranges from 30.2°C to 31.1°C. It may be due to the low wind speed of fewer than 1 m/s for years. Thus, the variability of SST is seen as ruled by solar radiation. The minimum solar radiation causes the minimum SST in July. However, the absence of strong wind prevents latent heat loss that maintains the high SST inside Cenderawasih Bay.

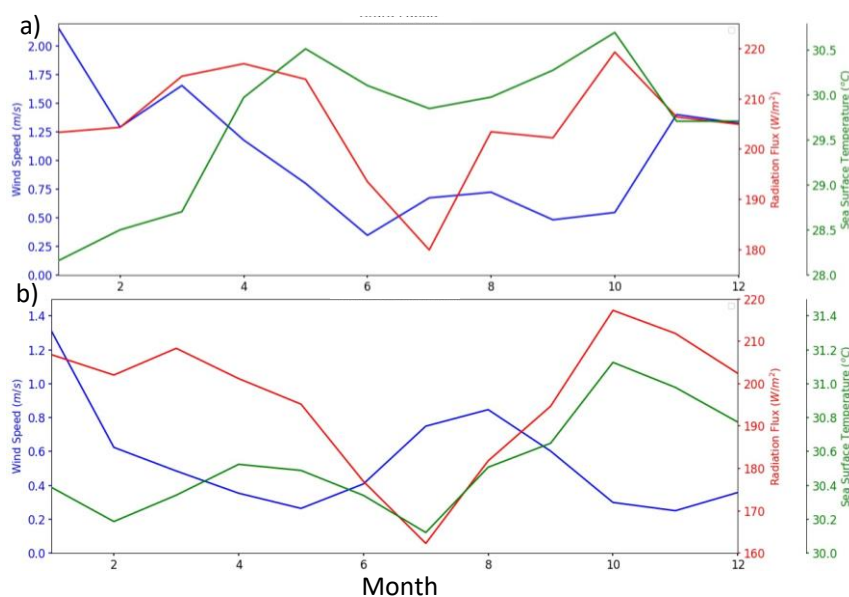


Fig. 5. Monthly climatology of SST, wind speed, and solar radiation at a) area outside and b) area inside Cenderawasih Bay, as shown in Fig. 1.

The oceanic aspect of Cenderawasih Hot Pool

To investigate the oceanic aspect of the Cenderawasih Hot Pool, we plotted the current patterns at 100 m depth shown in Fig. 6. A northwestward current has been identified north of Cenderawasih Bay. During the wet season (January), the speed of this current ranges from 0.5 m/s to 0.7 m/s. During the dry season (August), when the northwesterly wind disappears (Fig. 4b), the speed of the northwestward current increases. This northwestward current is known as the New Guinea Coastal Under Current (NGCUC). NGCUC is a permanent current feature at 200 m depth regardless of the wind reversals (Tsuchiya et al. 1989). Inside Cenderawasih Bay, strong currents are absent in this area, preventing the mixing process from maintaining the high temperature in the surface layer. During the dry and wet seasons, Cenderawasih Bay is fully isolated since the westward subsurface water flow does not enter Cenderawasih Bay. Thus, this isolated basin causes the Cenderawasih Bay area to be influenced mainly by the air-sea interaction process mentioned in the previous section creating Cenderawasih Hot Pool.

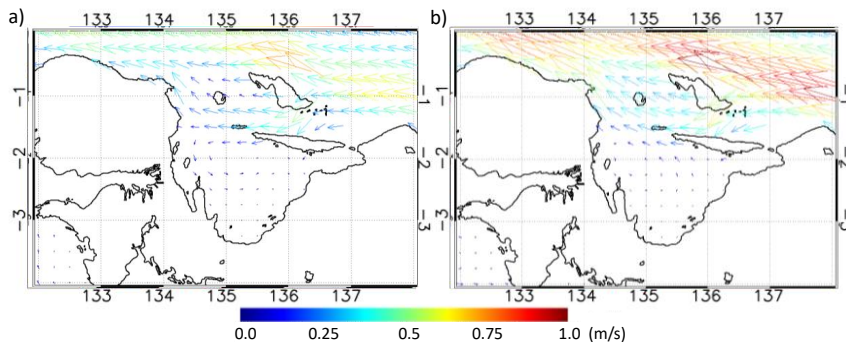


Fig. 6. Monthly climatology of current at 100 m depth in January (a) and August (b)

CONCLUSIONS

The hot Pool is used to categorize high SST episodes ($>30^{\circ}\text{C}$) in particular areas and during specific periods (relatively long periods). Since the term "warm pool" has been used to define an area with the annual average SST distributed below 30°C in dominant, therefore "hot pool" was taken to represent the area with SST dominated above 30°C .

Under high solar radiation and low wind speed, this event is characterized by considerable daily heat gains. The constant high SST occurrence in Cenderawasih Bay, which is more than 50% of 13 years, SSTs can reach more than 30°C and are dominated by the condition of low wind speed, i.e., 80% wind speed is less than m/s along the years, also 50% solar radiation measured is more than 200 W/m^2 . The current pattern at 100 m depth also shows that NGCUC does not enter Cenderawasih Bay. Thus, this indicates that Cenderawasih Bay is an isolated water. Those are the solid evidence for defining Cenderawasih Bay as the "Cenderawasih Hot Pool."

ACKNOWLEDGMENTS

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REFERENCES

- Arking A. and Ziskin D. (1994). Relationship between clouds and sea surface temperatures in the western tropical Pacific. *Journal of Climate*, 7(6), 988-1000. DOI: [https://doi.org/10.1175/1520-0442\(1994\)007%3C0988:RBCASS%3E2.0.CO;2](https://doi.org/10.1175/1520-0442(1994)007%3C0988:RBCASS%3E2.0.CO;2).
- Brown O.B. and Minnett P.J. (2009). MODIS infrared sea surface temperature algorithm theoretical basis document, ver 2.0. [online]. Available at: https://modis.gsfc.nasa.gov/data/atbd/atbd_mod25.pdf; https://modis.gsfc.nasa.gov/data/atbd/atbd_mod25.pdf. [Accessed 25 October 2022].
- Chongyin L., Mingquan M., and Guangqing Z. (1999). The variation of warm pool in the equatorial western Pacific and its impacts on climate. *Advances in Atmospheric Sciences*, 16(3), 378-394. DOI: <https://doi.org/10.1007/s00376-999-0017-0>.
- Clement A. and Seager R. (1999). Climate and the tropical oceans. *Journal of Climate*, 12(12), 3383-3401. DOI: [https://doi.org/10.1175/1520-0442\(1999\)012%3C3383:CATTO%3E2.0.CO;2](https://doi.org/10.1175/1520-0442(1999)012%3C3383:CATTO%3E2.0.CO;2).

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[DOI: 10.1175/1520-0442\(1999\)012%3C3383:CATTO%3E2.0.CO;2](https://doi.org/10.1175/1520-0442(1999)012%3C3383:CATTO%3E2.0.CO;2)

Clement A.C., Seager R., and Murtugudde R. (2005). Why are there tropical warm pools? *Journal of Climate*, 18(24), 5294-5311. [DOI: https://doi.org/10.1175/JCLI3582](https://doi.org/10.1175/JCLI3582)

Esaias, W.E., Abbott M.R., Barton I., Brown O.B., Campbell J.W., Carder K.L., ... and Minnett P.J. (1998). An overview of MODIS capabilities for ocean science observations, *IEEE Transactions on Geoscience and Remote Sensing*, 36(4), 1250-1265. [DOI: https://doi.org/10.1109/36.701076](https://doi.org/10.1109/36.701076)

Ghanea M., Moradi M., Kabiri K., and Mehdinia A. (2015). Investigation and validation of MODIS SST in the northern Persian Gulf, *Advances in Space Research*, 57(1), 127-136. [DOI: https://doi.org/10.1016/j.asr.2015.10.040](https://doi.org/10.1016/j.asr.2015.10.040)

Harweijer C., Seager R., Winton M., and Clement A.M.Y. (2005). Why ocean heat transport warms the global mean climate. *Tellus*, 57(4), 662-675. [DOI: https://doi.org/10.3402/tellusa.v57i4.14708](https://doi.org/10.3402/tellusa.v57i4.14708)

Hosoda K. (2013). Empirical method of diurnal correction for estimating sea surface temperature at dawn and noon. *Journal of Oceanography*, 69, 631-646. [DOI: https://doi.org/10.1007/s10872-013-0194-4](https://doi.org/10.1007/s10872-013-0194-4)

Ihsan E.N., Enita S.Y., and Wirasatriya, A. (2018). Oceanographic factors in fishing ground location of Anchovy at Teluk Cenderawasih National Park, West Papua: Are these factors have an effect of Whale Sharks appearance frequencies? In *IOP Conference Series: Earth and Environmental Science*, 116(1), p. 012-017. [DOI: https://doi.org/10.1088/1755-1315/116/1/012017](https://doi.org/10.1088/1755-1315/116/1/012017)

Kawamura H., Qin H., and Ando, K. (2008). In-situ diurnal sea surface temperature variations and near-surface thermal structure in the tropical hot event of the Indo-Pacific warm pool. *Journal of Oceanography*, 64, 847-857. [DOI: https://doi.org/10.1007/s10872-008-0070-9](https://doi.org/10.1007/s10872-008-0070-9)

Ming-An, L., Tzeng, M. T., Hosoda, K., Sakaida, F., Kawamura, H., Shieh, W. J., Yang Y., and Chang, Y. (2010). Validation of JAXA/MODIS sea surface temperature in water around Taiwan using the Terra and Aqua satellites. *TAO: Terrestrial, Atmospheric and Oceanic Sciences*, 21(4), 727-736. [DOI: https://doi.org/10.3319/TAO.2009.09.07.01\(Oc\)](https://doi.org/10.3319/TAO.2009.09.07.01(Oc))

Pierrehumbert R.T. (2000). Climate change and the tropical Pacific: The sleeping dragon wakes. *Proceedings of the National Academy of Sciences*, 97(4), 1355-1358. [DOI: https://doi.org/10.1073/pnas.97.4.1355](https://doi.org/10.1073/pnas.97.4.1355)

Qin H., Kawamura H., and Kawai Y. (2007). Detection of Hot Event in the equatorial Indo-Pacific warm pool using advanced satellite sea surface temperature, solar radiation, and wind speed. *Journal of Geophysical Research: Oceans*, 112 (C7). [DOI: https://doi.org/10.1029/2006JC003969](https://doi.org/10.1029/2006JC003969)

Qin H., Kawamura H., Sakaida F., and Ando K. (2008). A case study of the tropical hot event in November 2006 (HE0611) using a geostationary meteorological satellite and the TAO/TRITON mooring array. *Journal of Geophysical Research: Oceans*, 113(C8). [DOI: https://doi.org/10.1029/2007JC004640](https://doi.org/10.1029/2007JC004640)

Qin H. and Kawamura H. (2009). Atmosphere response to a hot SST event in November 2006 as observed by AIRS instrument. *Advances in space research*, 44(3), 395-400. [DOI: https://doi.org/10.1016/j.asr.2009.03.003](https://doi.org/10.1016/j.asr.2009.03.003)

Qin H. and Kawamura H. (2010). Air-sea interaction throughout the troposphere over a very high sea surface temperature. *Geophysical research letters*, 37(1), 1-4. [DOI: https://doi.org/10.1029/2009GL041685](https://doi.org/10.1029/2009GL041685)

Qin H., Chen G., Wang W., Wang D., and Zeng L. (2014). Validation and application of MODIS-derived SST in the South China Sea. *International journal of remote*

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sensing, 35(11-12), 4315-4328. <https://doi.org/10.1080/01431161.2014.916439>, DOI: [10.1080/01431161.2014.916439](https://doi.org/10.1080/01431161.2014.916439).

Formatted

Ramanathan V. and Collins W. (1991). Thermodynamic regulation of ocean warming by cirrus clouds deduced from observations of the 1987 El Niño. *Nature* 351 (6321), 27-32. <https://doi.org/10.1038/351027a0>, DOI: [10.1038/351027a0](https://doi.org/10.1038/351027a0).

Formatted

Setiawan R.Y., Wirasatriya A., Hernawan U., Leung S., and Iskandar I. (2020). Spatio-temporal variability of surface chlorophyll-*a* in the Halmahera Sea and its relation to ENSO and the Indian Ocean Dipole. *International Journal of Remote Sensing*, 41(1), 284-299. DOI: <https://doi.org/10.1080/01431161.2019.1641244>.

Formatted: Font color: Auto

Swandiko M., Wirasatriya A., Marwoto J., Muslim, Indrayanti E., Subardjo P., Ismunarti D.H. (2021). Studi persistensi suhu permukaan laut tinggi (>30°C) di perairan Selat Malaka. *Buletin Oseanografi Marina*, 10(2), 162-170. [online]. Available at: <https://ejournal.undip.ac.id/index.php/buloma/article/download/31554/19352> <https://ejournal.undip.ac.id/index.php/buloma/article/download/31554/19352>. [Accessed 25 October 2022] (in Indonesian).

Formatted

De Garidel-Thoron T., Rosenthal Y., Bassinot F., and Beaufort, L. (2005). Stable sea surface temperatures in the western Pacific warm pool over the past 1.75 million years. *Nature*, 433(7023), 294-298. <https://doi.org/10.1038/nature03189>, DOI: [10.1038/nature03189](https://doi.org/10.1038/nature03189).

Formatted: Font color: Auto

Tita A.D.C., Wirasatriya A., Sugianto D.N., Maslukah L., Handoyo G., Helmi M., and Avianto, P. (2020). Persistence of high sea surface temperature (> 30°C) in Tomini Bay. In *IOP Conference Series: Earth and Environmental Science*, 530(1), p.012-038, DOI: <https://doi.org/10.1088/1755-1315/530/1/012038>.

Formatted: Font color: Auto

Waliser D.E. and Graham N.E. (1993). Convective cloud systems and warm-pool sea-surface temperatures: Coupled interactions and self-regulation. *Journal of Geophysical Research: Atmospheres*, 98(D7), 12881-12893. <https://doi.org/10.1029/93JD00872>, DOI: [10.1029/93JD00872](https://doi.org/10.1029/93JD00872).

Formatted

Wallace J.M. (1992). Effect of deep convection on the regulation of tropical sea surface temperature. *Nature*, 357(6375), 230-231. <https://doi.org/10.1038/357230a0>, DOI: [10.1038/357230a0](https://doi.org/10.1038/357230a0).

Formatted

Wirasatriya A., Kawamura H., Shimada T., and Hosoda K. (2015). Climatology of hot events in the western equatorial Pacific. *Journal of Oceanography*, 71, 77-90. <https://doi.org/10.1007/s10872-014-0263-3>, DOI: [10.1007/s10872-014-0263-3](https://doi.org/10.1007/s10872-014-0263-3).

Formatted

Wirasatriya A., Kawamura H., Shimada T., Hosoda K. (2016). Atmospheric structure favoring high sea surface temperatures in the western equatorial Pacific. *Journal of Geophysical Research*, 121(19), 11-368. <https://doi.org/10.1002/2016JD025268>, DOI: [10.1002/2016JD025268](https://doi.org/10.1002/2016JD025268).

Formatted

Wirasatriya A., Sugianto D.N., and Helmi M. (2017). The Influence of Madden Julian Oscillation on the Formation of the hot event in the western equatorial Pacific. In *IOP Conference Series: Earth and Environmental Science*, 55(1), p.012-006. <https://doi.org/10.1088/1755-1315/55/1/012006>, DOI: [10.1088/1755-1315/55/1/012006](https://doi.org/10.1088/1755-1315/55/1/012006).

Formatted: Font color: Auto

Wirasatriya A., Setiawan R.Y., and Subardjo P. (2017). The effect of ENSO on the variability of chlorophyll-*a* and sea surface temperature in the Maluku Sea, *IEEE Journal of Selected Topics in Applied Earth Observations and Remote Sensing*, 10(12), 5513-5518. <https://doi.org/10.1109/JSTARS.2017.2745207>, DOI: [10.1109/JSTARS.2017.2745207](https://doi.org/10.1109/JSTARS.2017.2745207).

Formatted

Wirasatriya A., Prasetyawan I.B., Triyono C.D., Muslim, and Maslukah L. (2018). Effect of ENSO on the variability of SST and chlorophyll-*a* in Java Sea. In *IOP Conference Series: Earth and Environmental Science*, 116(1), p.012-063. <https://doi.org/10.1088/1755-1315/116/1/012063>, DOI: [10.1088/1755-1315/116/1/012063](https://doi.org/10.1088/1755-1315/116/1/012063).

Formatted: Font color: Auto

Wirasatriya A., Sugianto D.N., Helmi M., Setiawan R.Y., and Koch M. (2019). Distinct characteristics of SST variabilities in the Sulawesi Sea and the northern part of the Maluku Sea during the southeast monsoon. *IEEE Journal of Selected Topics in Applied Earth*

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Observations and Remote Sensing, 12(6), 1763-1770. <https://doi.org/10.1109/JSTARS.2019.2913739>, DOI: 10.1109/JSTARS.2019.2913739.

Wirasatriya A., Kawamura H., Helmi M., Sugianto D.N., Shimada T., Hosoda K., Handoyo G., Putra Y.D.G., and Koch M. (2020). Thermal structure of Hot Events and their possible role in maintaining the warm isothermal layer in the Western Pacific warm pool. *Ocean Dynamics*, 70, 771–786. <https://doi.org/10.1007/s10236-020-01362-8>, DOI: 10.1007/s10236-020-01362-8.

Wyrski K. (1989). Some thoughts about the West Pacific warm pool. In *Proceedings of the Western Pacific International Meeting and Workshop on TOGA COARE*, Nouméa, New Caledonia: 99-109 [online]. Available at: https://horizon.documentation.ird.fr/exl-doe/pleins_textes/doc34-08/30195.pdf#page=114 https://horizon.documentation.ird.fr/exl-doe/pleins_textes/doc34-08/30195.pdf#page=114 [(Accessed 26 October 2022).]

Yan X.H., Ho C.R., Zheng Q., and Klemas V. (1992). Temperature and size variabilities of the western Pacific warm pool. *Science*, 258 (5088), 1643-1645. <https://doi.org/10.1126/science.258.5088.1643>, DOI: 10.1126/science.258.5088.1643.

Yu L. and Weller R.A. (2007). Objectively Analyzed air-sea heat fluxes for the global ice-free oceans (1981–2005). *Bulletin of the American Meteorological Society*, 88(4), 527-540. <https://doi.org/10.1175/BAMS-88-4-527>, DOI: 10.1175/BAMS-88-4-527.

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CENDERAWASIH HOT POOL: THE FREQUENT HIGH SEA SURFACE TEMPERATURE PHENOMENA AT CENDERAWASIH BAY, PAPUA

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ABSTRACT. The term “warm pool” refers to a body of water with the characteristic of SST exceeding 28°C within a particular area and a relatively long period in an annual circle. However, there are regions with an annual mean SST measured above 30°C, and we classified them as hot pools because of the conditions of intense solar radiation and low wind speed. One of the Hot Pool spots was found in Indonesia, in Cenderawasih Bay. The present study examines the existence of the Cenderawasih Hot Pool using long-term observation of satellite SST data. In order to learn more about their mechanisms, we also analyzed surface wind, surface heat flux, and surface current data. The results show that SSTs in Cenderawasih Bay have a 50% chance of exceeding 30°C within the 13 years of study (2013-2015). Heat input comes from strong solar radiation, i.e., 50% of solar radiation is more than 200 W/m². The location is also dominated by low wind speed, i.e., 80% wind speed of lower than 4 m/s, which caused the low latent loss in Cenderawasih Bay. Cenderawasih Bay is fully separated from surface currents during the dry and wet seasons since the easterly subsurface water flow does not enter the bay. The absence of strong currents prevents the mixing process, maintaining the high temperature in the surface layer. Those processes are discovered and they serve as compelling evidence to support Cenderawasih Bay as one of the Hot Pool areas within the Indonesian seas.

KEYWORDS: global climate; sea surface temperature; Hot Pool spot; Cenderawasih Bay

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INTRODUCTION

The western equatorial Pacific has a big impact on the climate. The Earth's circulation is specifically affected by the warm pool, an area with average sea surface temperatures (SST) exceeding 28°C (e.g., Wyrski 1989; Yan et al. 1992; Clement and Seager 1999; Chongyin et al. 1999; Pierrehumbert 2000; Clement et al. 2005; Thoron et al. 2005; Herweijer et al. 2005).

On the other hand, high SST in tropical regions has attracted researchers to investigate the mechanisms since the formation of high SST requires a particular atmospheric process (e.g., Ramanathan and Collins 1991; Wallace 1992; Arking and Ziskin 1994). This process is depicted by (Waliser and Graham 1993), which shows the relation between SSTs and deep convection. They utilized highly reflective cloud information from an arbitrary examination of monthly SST data with a grid spacing of 2° (produced from combined satellite observation and in situ data) and daily visible and infrared satellite image data. The highly reflective cloud increases along with the SST as it rises from 26°C to 29.5°C. In contrast, the highly reflective cloud diminishes with increased SST in the temperature from 29.5°C to 32°C. As a result, the analysis proved that several atmospheric processes impacted SSTs below and above 29.5°C.

By taking advantage of high temporal and spatial resolution SST products derived from satellite observations (i.e., daily and $\leq 25 \text{ km} \times 25 \text{ km}$), several studies (Kawamura et al. 2008; Qin et al. 2007, 2008; Qin and Kawamura 2009, 2010; Wirasatriya et al. 2015, 2016, 2017a, 2020) were able to identify high SST events (i.e., more than 30°C) in specific areas and at certain periods and define them as Hot Event (HE). In summary, they concluded that considerable daily heat gains characterize the production of HE under high solar radiation and low wind speed brought on by "remote convection" mechanisms.

(Wirasatriya et al. 2015) elaborated on the climatology of HEs in the western equatorial Pacific using the SST dataset derived from satellite microwave sensors. Throughout nine years of observation (2003-2011), they discovered 71 HE cases in the western equatorial Pacific, with the majority centered on the Solomon Islands and New Guinea Island's northern coasts, which extend eastward up to 160°W. According to the climatology, the region has solar radiation of more than 200 W/m² and wind speeds of less than 4 m/s. Low wind speeds heavily influence the mechanism for HE incidence in the western equatorial Pacific. Much of the equatorial region experiences sun radiation above 200 W/m² during the HE periods. Low wind speeds minimize latent heat loss, which results in high SSTs and HEs in specific locations. (Wirasatriya et al. 2015) also emphasized that high solar radiation and low wind speed are much more common during the development stage and less common during the decay stage. This study also demonstrated that a rise in the long-term mean SST in the western equatorial Pacific is correlated with an increase in the frequency of HE events. HEs were responsible for 51.5% of the SSTs >30°C in the warm pool region bounded by the 29.5 °C isotherms of the climatological SST. Thus, statistically, there is a relation between the occurrence of HE and the formation of the western Pacific warm pool. Moreover, (Wirasatriya et al. 2020) demonstrated the role of HE in maintaining the warm mixed layer in the western Pacific warm pool. The frequent occurrence of HE transports heat from the surface layer to the deeper layer above the thermocline to maintain the warm mixed layer in the warm pool. Since the Pacific warm pool influences the global climate's variability through coupled ocean-atmosphere dynamics and thermodynamics (e.g.,

Clement and Seager 1999; Herweijer et al. 2005), this becomes an example of the contribution of high SST phenomena to regulate the global climate.

Within the Indonesian seas, the frequent appearance of high SST $> 30^{\circ}\text{C}$ has been reported by (Tita et al. 2020) in Tomini Bay and by (Swandiko et al. 2021) in the Malacca Strait. Their appearances also the high solar radiation and weak wind. The morphology of the semi-enclosed waters causes the high-frequency occurrence of low wind speed by less than 4 m/s in both areas. Besides these two areas, there is no more study on the high SST phenomena within the Indonesian Seas.

The present study demonstrates the frequent high SST occurrence (more than 30°C) in Cenderawasih Bay. Cenderawasih Bay is in northern Papua Island, part of the western Pacific warm pool. It is semi-enclosed water with a deep basin, surrounded by mountain chains (Fig. 1). Cenderawasih Bay is the habitat for whale sharks (Ihsan et al. 2018). Since the area determines warm pool definition with SST climatology higher than 28°C , we call the high-frequency occurrence of SSTs higher than 30°C in the Cenderawasih Bay the Cenderawasih Hot Pool. (Wirasatriya et al. 2015) found that the duration of HE occurrence in the western equatorial Pacific is no longer than two months since HE requires a typical condition of high solar radiation and low wind speed. Thus, it is interesting to understand the mechanisms of the high-frequency occurrence of high SST in the Cenderawasih Hot Pool.

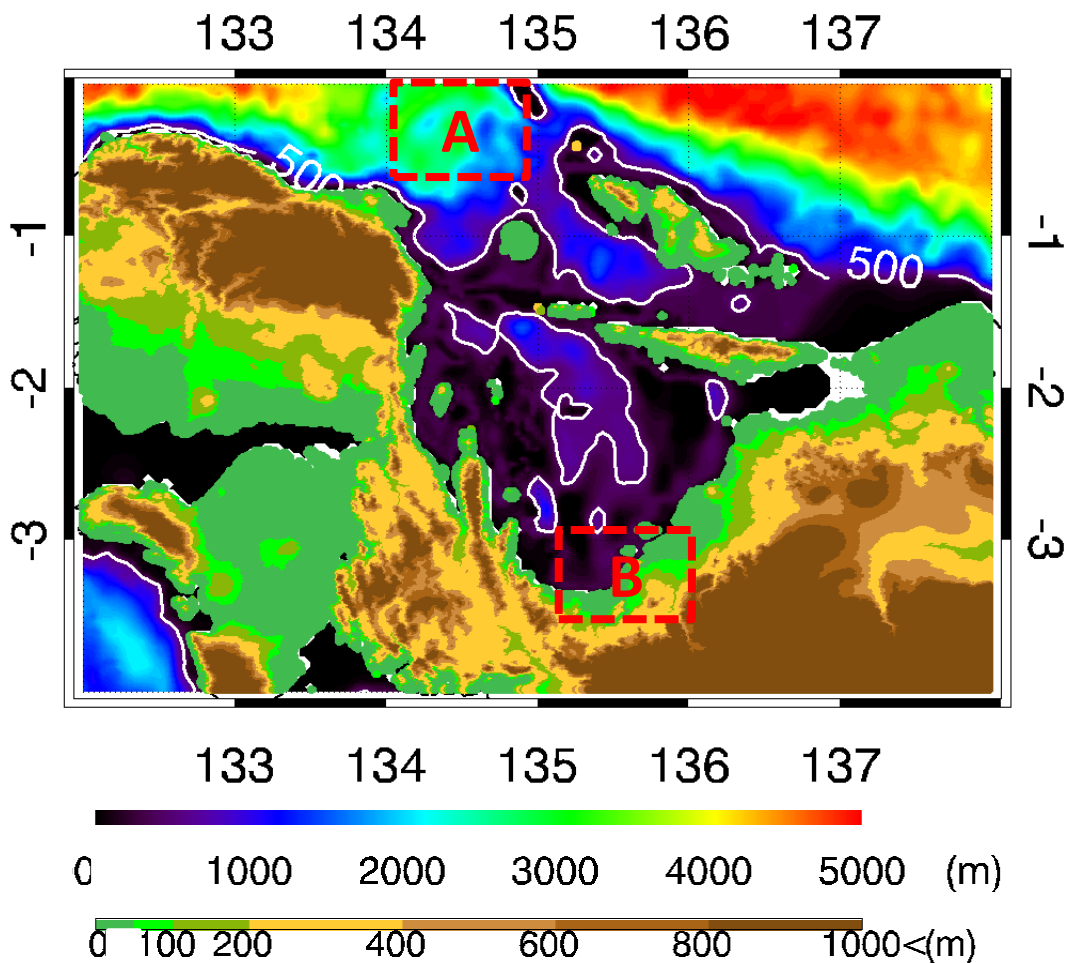


Fig. 1. Bathymetry and topography of Cenderawasih Bay. The dashed boxes A and B represent the sampling area for outside and inside Cenderawasih Bay,

respectively, as shown in Fig. 5

MATERIALS AND METHODS

We used the semi-daily 11 μm SST products from Moderate Resolution Imaging Spectroradiometer (MODIS) Aqua Lv3 with a spatial resolution of $0.04^\circ \times 0.04^\circ$ (Esaias et al. 1998) and observation period from 2003 to 2015. The MODIS SST 11 μm the Multi-Channel SST algorithm generates m by using brightness temperatures at 11 μm and 12 μm (Brown and Minnet 2009). The best accuracy was achieved by validating and testing this data against in-situ measurements (e.g., Lee et al. 2010; Qin et al. 2014; Ghanea et al. 2015).

We can obtain the Cross-Calibrated Multi-Platform (CCMP) gridded surface vector winds version 2.0 from the following website: www.remss.com, and use it to calculate surface winds. As a result, CCMP is regarded as a Level-3 ocean vector wind analysis product. They are created utilizing data from moored buoys, satellites, and model winds. The surface wind data for spatial and temporal resolutions is $0.25^\circ \times 0.25^\circ$ respectively and quarter daily. This wind product has greater accuracy than other wind reanalysis data.

We used a 0.5° grid from the Objectively Analyzed Air-Sea Fluxes project to analyze daily latent heat fluxes and solar radiation (Yu and Weller 2007). The period of observation for latent heat flux is from 2003 to 2015, but for solar radiation is only from 2003 to 2009 due to the availability of the data. Global 30 Arc-Second Elevation data was used to obtain topography. Global 30 Arc-Second Elevation (GTOPO30) data was used to obtain topography¹. ETOPO, a one-arc-minute global relief model of Earth's surface that combines land topography and ocean bathymetry, provides bathymetry². For oceanic parameters, we used the current reanalysis data from <http://marine.copernicus.eu/> to obtain the seasonal variation of the current pattern at 100 m depth in the Cenderawasih Hot Pool.

To calculate the frequency occurrence of SST within 13 years of observation, we calculate the percentage of SST $> 30^\circ\text{C}$ occurrence in each grid. Since high SST of more than 30°C occurs during the condition of low wind speed (i.e., < 4 m/s) and high solar radiation (i.e., > 200 W/m²) (Wirasatriya et al. 2015), we calculated the percentage of low wind speed and high solar radiation in the Cenderawasih Bay. Furthermore, we also calculated the percentage of low latent heat release (< 120 W/m²) to explain how wind speed influences the variability of SST. The equation for calculating the percentage is as follows:

$$\%(x, y) = \frac{1}{n} \sum_{i=1}^n p_i(x, y) \times 100\% \quad (1)$$

where $\%(x, y)$ is percentage the percentage of high SST $> 30^\circ\text{C}$ or percentage of weak wind speed < 4 m/s or percentage of solar radiation > 200 W/m² or percentage of latent heat flux < 120 W/m² at position (x,y); p_i is the amount of SST data $> 30^\circ\text{C}$ or the amount of weak wind speed < 4 m/sec or the amount solar radiation > 200 W/m² or the amount of latent heat flux < 120 W/m² at position (x,y). Furthermore, we also analyze the climatological mean of each parameter. To create monthly and monthly climatology, all geophysical parameters are incorporated as follows (Wirasatriya et al. 2017b):

$$\bar{X}(x, y) = \frac{1}{n} \sum_{i=1}^n x_i(x, y, t) \quad (2)$$

where $\bar{X}(x, y)$ is monthly mean value or monthly climatology value at position (x,y), $x_i(x, y, t)$ is i^{th} value of the data at (x,y) position and time t. Next, n is the number of data

¹ https://www.usgs.gov/centers/eros/science/usgs-eros-archive-digital-elevation-global-30-arc-second-elevation-gtopo30?qt-science_center_objects=0#qt-science_center_objects

² <https://www.ngdc.noaa.gov/mgg/global/>

in 1 month and the number of monthly data in 1 period of climatology (i.e., from 2003 to 2016 = 13 data) for monthly calculation and monthly climatology calculation, respectively. Moreover, if pixel xi is a hollow pixel, it is not included in the calculation.

RESULTS AND DISCUSSION

Cenderawasih Hot Pool

The definition of the Cenderawasih Hot Pool is described in Fig. 2, which shows the climatological mean of SST and the percentage of daily SSTs more than 30°C from 2003 to 2015 (13 years). In Cenderawasih Bay, the mean SST at the southern part of 2°S is more than 30°C, and the high SSTs are more than 50%. It means that the mean SST in Cenderawasih Bay is higher than the definition of the western Pacific warm pool which is 28°C (Wyrtki 1989), and more than 3.5 years within 2003-2015, SSTs in Cenderawasih Bay can reach more than 30°C. At the southernmost of the bay, the mean SST reaches 30.5°C, with the percentage of high SSTs being more than 70%. It has become the hottest area of Cenderawasih Bay.

In contrast, in the area outside Cenderawasih Bay, i.e., the western Pacific warm pool's offshore seas, the percentage is less than 30%. Thus, this evidence supports the definition of Cenderawasih Bay as the Cenderawasih Hot Pool. The high-frequency occurrence of high SST in Cenderawasih Bay may correspond to whale sharks' yearly appearance. (Ihsan et al. 2018) indicated the tolerance of whale sharks in high SST conditions.

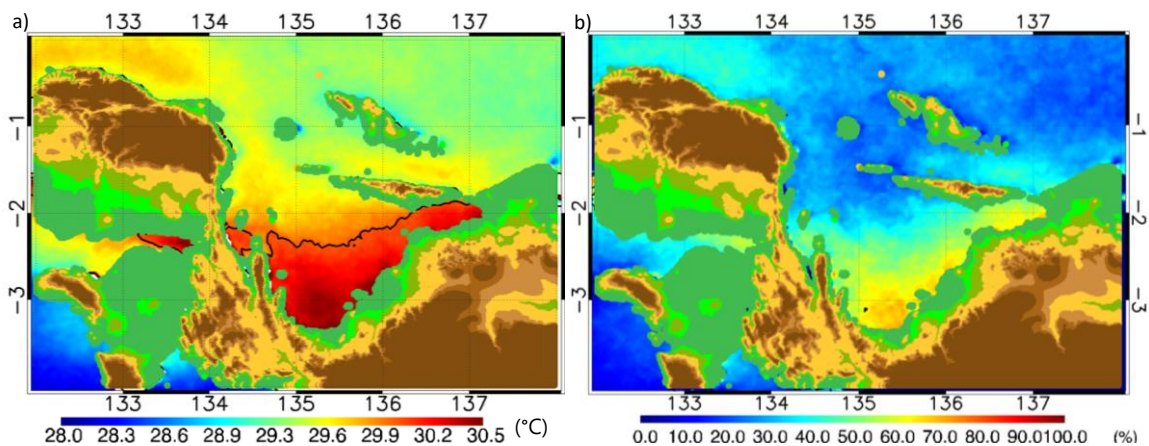


Fig. 2. Climatological mean of SST (a) and percentage of daily SSTs > 30°C during thirteen years of observation (2003-2015) (b). The black contour is 30°C

Atmospheric Aspect of Cenderawasih Hot Pool

Fig. 3 clearly shows that low wind speed dominates Cenderawasih Bay, i.e., 80% wind speed < 4 m/s occurred during 2003-2015. This causes low latent loss in Cenderawasih Bay. For solar radiation, the percentage of high solar radiation inside and outside Cenderawasih Bay is almost similar i.e., nearly 50%, indicating that high solar radiation is not dominant in both areas. This means that, although high solar radiation does not frequently occur, the absence of strong wind speed may maintain the latent heat loss to keep the SST in Cenderawasih Bay higher than 30°C.

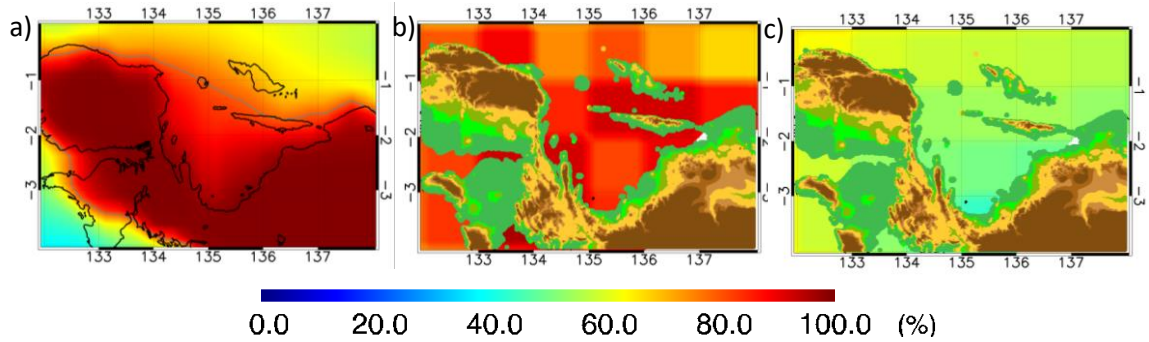


Fig. 3. The percentage of wind speed less than four m/s (a), latent heat release less than 120 W/m² (b), and solar radiation more than 200 W/m² (c)

To investigate how low wind speed occurs in the study area, we show the monthly climatology map of surface wind during summer and winter (Fig. 4). For both seasons, Cenderawasih Bay is protected from high wind speed due to high mountain chains in the western, southern, and eastern parts. In the northern part, small islands at the mouth of the bay prevent the strong wind speed from entering the bay. Thus, the role of topography is crucial for the occurrence of constant high SSTs in Cenderawasih Bay. The same tendencies are also found at the Tomini Bay and Malacca Strait, as indicated by (Tita et al. 2020) and (Swandiko et al. 2021).

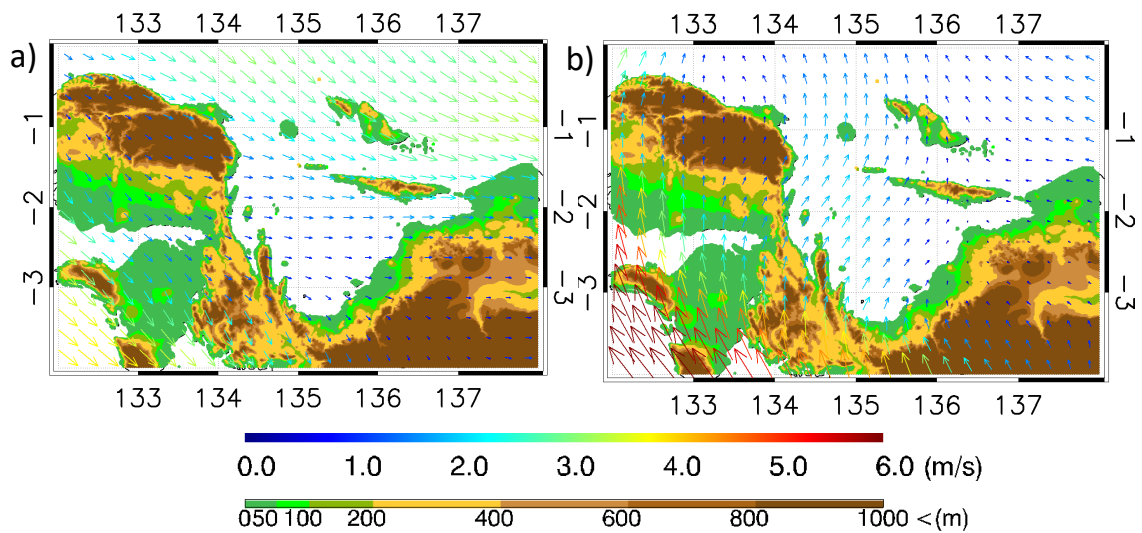


Fig. 4. Monthly climatology of surface wind speed during a) wet season (January) and b) dry season (August)

The influence of wind speed and solar radiation on the SST variation is further examined by plotting the monthly climatology of wind speed, solar radiation, and SST at areas A and B (Fig. 1), representing the areas outside and inside Cenderawasih Bay. Seasonal variations of SSTs are observed in both areas. The minimum SSTs occur during the wet and dry seasons, while the maximum SSTs during the transition season. This seasonal SST variability is similar to the other areas in the Indonesian seas, such as the Java Sea (Wirasatriya et al. 2018), Maluku Sea (Wirasatriya et al. 2019), Halmahera Sea (Setiawan et al. 2019), etc.

For the area outside Cenderawasih Bay, SST ranges from 28.2°C to 30.6°C. It is seen that the variability of SST is controlled by wind speed and solar radiation. From January to

June, when wind speed decreases from 2.2 m/s to 0.3 m/s and solar radiation is more than 200 W/m², SST increases from 28.2°C to 30.5°C. From June to October, wind speed is lower than 1 m/s. In the absence of strong wind, the low SST in July is caused by the decrease of solar radiation to a minimum.

For the area inside Cenderawasih Bay, persistent high SST is observed as the SST ranges from 30.2°C to 31.1°C. It may be due to the low wind speed of fewer than 1 m/s for years. Thus, the variability of SST is seen as ruled by solar radiation. The minimum solar radiation causes the minimum SST in July. However, the absence of strong wind prevents latent heat loss that maintains the high SST inside Cenderawasih Bay.

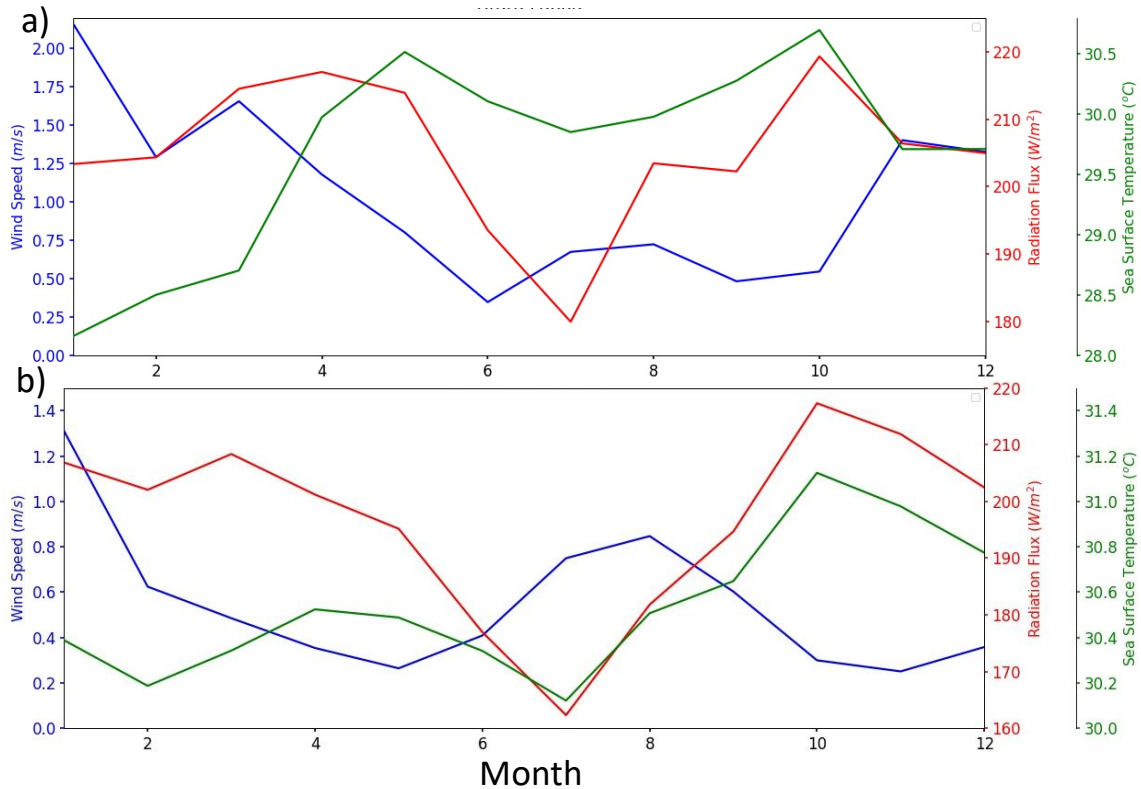


Fig. 5. Monthly climatology of SST, wind speed, and solar radiation at a) area outside and b) area inside Cenderawasih Bay, as shown in Fig. 1

The oceanic aspect of Cenderawasih Hot Pool

To investigate the oceanic aspect of the Cenderawasih Hot Pool, we plotted the current patterns at 100 m depth shown in Fig. 6. A northwestward current has been identified north of Cenderawasih Bay. During the wet season (January), the speed of this current ranges from 0.5 m/s to 0.7 m/s. During the dry season (August), when the northwesterly wind disappears (Fig. 4b), the speed of the northwestward current increases. This northwestward current is known as the New Guinea Coastal Under Current (NGCUC). NGCUC is a permanent current feature at 200 m depth regardless of the wind reversals (Tsuchiya et al. 1989). Inside Cenderawasih Bay, strong currents are absent in this area, preventing the mixing process from maintaining the high temperature in the surface layer. During the dry and wet seasons, Cenderawasih Bay is fully isolated since the westward subsurface water flow does not enter Cenderawasih Bay. Thus, this isolated basin causes the Cenderawasih Bay area to be influenced mainly by the air-sea interaction process mentioned in the previous section creating Cenderawasih Hot Pool.

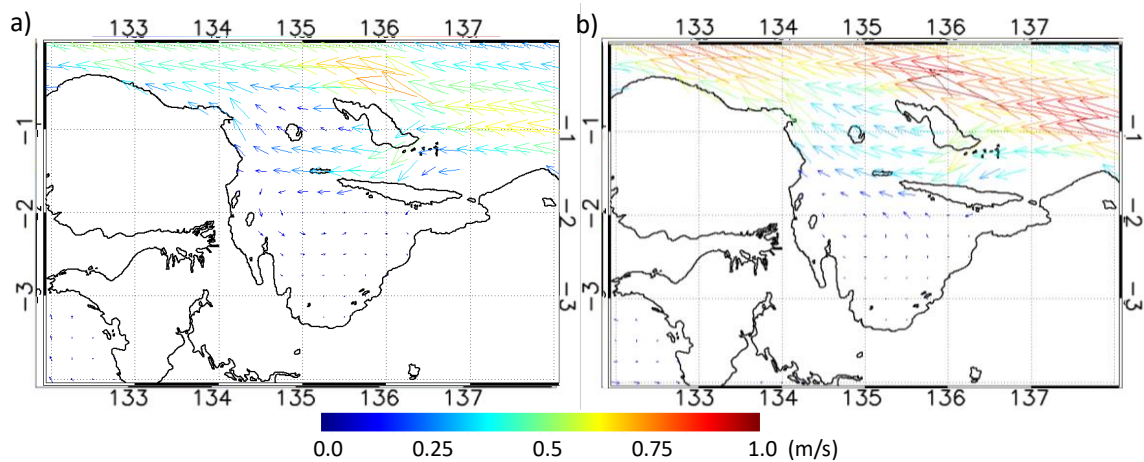


Fig. 6. Monthly climatology of current at 100 m depth in January (a) and August (b)

CONCLUSIONS

The hot Pool is used to categorize high SST episodes ($>30^{\circ}\text{C}$) in particular areas and during specific periods (relatively long periods). Since the term "warm pool" has been used to define an area with the annual average SST distributed below 30°C in dominant, therefore "hot pool" was taken to represent the area with SST dominated above 30°C .

Under high solar radiation and low wind speed, this event is characterized by considerable daily heat gains. The constant high SST occurrence in Cenderawasih Bay, which is more than 50% of 13 years, SSTs can reach more than 30°C and are dominated by the condition of low wind speed, i.e., 80% wind speed is less than m/s along the years, also 50% solar radiation measured is more than 200 W/m^2 . The current pattern at 100 m depth also shows that NGCUC does not enter Cenderawasih Bay. Thus, this indicates that Cenderawasih Bay is an isolated water. Those are the solid evidence for defining Cenderawasih Bay as the "Cenderawasih Hot Pool."

REFERENCES

- Arking A. and Ziskin D. (1994). Relationship between clouds and sea surface temperatures in the western tropical Pacific. *Journal of Climate*, 7(6), 988-1000, DOI: 10.1175/1520-0442(1994)007%3C0988:RBCASS%3E2.0.CO;2.
- Brown O.B. and Minnett P.J. (2009). MODIS infrared sea surface temperature algorithm theoretical basis document, ver 2.0. [online]. Available at: https://modis.gsfc.nasa.gov/data/atbd/atbd_mod25.pdf [Accessed 25 October 2022].
- Chongyin L., Mingquan M., and Guangqing Z. (1999). The variation of warm pool in the equatorial western Pacific and its impacts on climate. *Advances in Atmospheric Sciences*, 16(3), 378-394, DOI: 10.1007/s00376-999-0017-0.
- Clement A. and Seager R. (1999). Climate and the tropical oceans. *Journal of Climate*, 12(12), 3383-3401, DOI: 10.1175/1520-0442(1999)012%3C3383:CATTO%3E2.0.CO;2
- Clement A.C., Seager R., and Murtugudde R. (2005). Why are there tropical warm pools? *Journal of Climate*, 18(24), 5294-5311, DOI: 10.1175/JCLI3582.1.
- Esaias W.E., Abbott M.R., Barton I., Brown O.B., Campbell J.W., Carder K.L., ... and Minnett P.J. (1998). An overview of MODIS capabilities for ocean science observations, *IEEE Transactions on Geoscience and Remote Sensing*, 36(4), 1250-1265, DOI: 10.1109/36.701076.
- Ghanea M., Moradi M., Kabiri K., and Mehdinia A. (2015). Investigation and validation of MODIS SST in the northern Persian Gulf, *Advances in Space Research*, 57(1),

127-136, DOI: 10.1016/j.asr.2015.10.040.

Harweijer C., Seager R., Winton M., and Clement A.M.Y. (2005). Why ocean heat transport warms the global mean climate. *Tellus*, 57(4), 662-675, DOI: 10.3402/tellusa.v57i4.14708.

Hosoda K. (2013). Empirical method of diurnal correction for estimating sea surface temperature at dawn and noon. *Journal of Oceanography*, 69, 631-646, DOI: 10.1007/s10872-013-0194-4.

Ihsan E.N., Enita S.Y., and Wirasatriya, A. (2018). Oceanographic factors in fishing ground location of Anchovy at Teluk Cenderawasih National Park, West Papua: Are these factors have an effect of Whale Sharks appearance frequencies? In *IOP Conference Series: Earth and Environmental Science*, 116(1), 012-017, DOI: 10.1088/1755-1315/116/1/012017.

Kawamura H., Qin H., and Ando, K. (2008). In-situ diurnal sea surface temperature variations and near-surface thermal structure in the tropical hot event of the Indo-Pacific warm pool. *Journal of Oceanography*, 64, 847-857, DOI: 10.1007/s10872-008-0070-9.

Ming-An L., Tzeng M. T., Hosoda K., Sakaida F., Kawamura H., Shieh W. J., Yang Y., and Chang, Y. (2010). Validation of JAXA/MODIS sea surface temperature in water around Taiwan using the Terra and Aqua satellites. *TAO: Terrestrial, Atmospheric and Oceanic Sciences*, 21(4), 727-736, DOI: 10.3319/TAO.2009.09.07.01(Oc).

Pierrehumbert R.T. (2000). Climate change and the tropical Pacific: The sleeping dragon wakes. *Proceedings of the National Academy of Sciences*, 97(4), 1355-1358, DOI: 10.1073/pnas.97.4.1355.

Qin H., Kawamura H., and Kawai Y. (2007). Detection of Hot Event in the equatorial Indo-Pacific warm pool using advanced satellite sea surface temperature, solar radiation, and wind speed. *Journal of Geophysical Research: Oceans*, 112 (C7), DOI: 10.1029/2006JC003969.

Qin H., Kawamura H., Sakaida F., and Ando K. (2008). A case study of the tropical hot event in November 2006 (HE0611) using a geostationary meteorological satellite and the TAO/TRITON mooring array. *Journal of Geophysical Research: Oceans*, 113(C8), DOI: 10.1029/2007JC004640.

Qin H. and Kawamura H. (2009). Atmosphere response to a hot SST event in November 2006 as observed by AIRS instrument. *Advances in space research*, 44(3), 395-400, DOI: 10.1016/j.asr.2009.03.003.

Qin H. and Kawamura H. (2010). Air-sea interaction throughout the troposphere over a very high sea surface temperature. *Geophysical research letters*, 37(1), 1-4, DOI: 10.1029/2009GL041685.

Qin H., Chen G., Wang W., Wang D., and Zeng L. (2014). Validation and application of MODIS-derived SST in the South China Sea. *International journal of remote sensing*, 35(11-12), 4315-4328, DOI: 10.1080/01431161.2014.916439.

Ramanathan V. and Collins W. (1991). Thermodynamic regulation of ocean warming by cirrus clouds deduced from observations of the 1987 El Niño. *Nature* 351 (6321), 27-32, DOI: 10.1038/351027a0.

Setiawan R.Y., Wirasatriya A., Hernawan U., Leung S., and Iskandar I. (2020). Spatio-temporal variability of surface chlorophyll-*a* in the Halmahera Sea and its relation to ENSO and the Indian Ocean Dipole. *International Journal of Remote Sensing*, 41(1), 284-299, DOI: 10.1080/01431161.2019.1641244.

Swandiko M., Wirasatriya A., Marwoto J., Muslim, Indrayanti E., Subardjo P., Ismunarti D.H. (2021). Studi persistensi suhu permukaan laut tinggi (>30°C) di perairan Selat Malaka. *Buletin Oseanografi Marina*, 10(2), 162-170. [online]. Available at: <https://ejournal.undip.ac.id/index.php/buloma/article/download/31554/19352> [Accessed 25 October 2022] (in Indonesian).

De Garidel-Thoron T., Rosenthal Y., Bassinot F., and Beaufort, L. (2005). Stable sea surface temperatures in the western Pacific warm pool over the past 1.75 million years. *Nature*, 433(7023), 294-298, DOI: 10.1038/nature03189.

Tita A.D.C., Wirasatriya A., Sugianto D.N., Maslukah L., Handoyo G., Helmi M., and Avianto P. (2020). Persistence of high sea surface temperature (> 30°C) in Tomini Bay. In *IOP Conference Series: Earth and Environmental Science*, 530(1), 012-038, DOI: 10.1088/1755-1315/530/1/012038.

Waliser D.E. and Graham N.E. (1993). Convective cloud systems and warm-pool sea-surface temperatures: Coupled interactions and self-regulation. *Journal of Geophysical Research: Atmospheres*, 98(D7), 12881-12893, DOI: 10.1029/93JD00872.

Wallace J.M. (1992). Effect of deep convection on the regulation of tropical sea surface temperature. *Nature*, 357(6375), 230-231, DOI: 10.1038/357230a0.

Wirasatriya A., Kawamura H., Shimada T., and Hosoda K. (2015). Climatology of hot events in the western equatorial Pacific. *Journal of Oceanography*, 71, 77-90, DOI: 10.1007/s10872-014-0263-3.

Wirasatriya A., Kawamura H., Shimada T., Hosoda K. (2016). Atmospheric structure favoring high sea surface temperatures in the western equatorial Pacific. *Journal of Geophysical Research*, 121(19), 11-368, DOI: 10.1002/2016JD025268.

Wirasatriya A., Sugianto D.N., and Helmi M. (2017). The Influence of Madden Julian Oscillation on the Formation of the hot event in the western equatorial Pacific. In *IOP Conference Series: Earth and Environmental Science*, 55(1), 012-006, DOI: 10.1088/1755-1315/55/1/012006.

Wirasatriya A., Setiawan R.Y., and Subardjo P. (2017). The effect of ENSO on the variability of chlorophyll-a and sea surface temperature in the Maluku Sea, *IEEE Journal of Selected Topics in Applied Earth Observations and Remote Sensing*, 10(12), 5513-5518, DOI: 10.1109/JSTARS.2017.2745207.

Wirasatriya A., Prasetyawan I.B., Triyono C.D., Muslim, and Maslukah L. (2018). Effect of ENSO on the variability of SST and chlorophyll-a in Java Sea. In *IOP Conference Series: Earth and Environmental Science*, 116(1), 012-063, DOI: 10.1088/1755-1315/116/1/012063.

Wirasatriya A., Sugianto D.N., Helmi M., Setiawan R.Y., and Koch M. (2019). Distinct characteristics of SST variabilities in the Sulawesi Sea and the northern part of the Maluku Sea during the southeast monsoon. *IEEE Journal of Selected Topics in Applied Earth Observations and Remote Sensing*, 12(6), 1763-1770, DOI: 10.1109/JSTARS.2019.2913739.

Wirasatriya A., Kawamura H., Helmi M., Sugianto D.N., Shimada T., Hosoda K., Handoyo G., Putra Y.D.G., and Koch M. (2020). Thermal structure of Hot Events and their possible role in maintaining the warm isothermal layer in the Western Pacific warm pool. *Ocean Dynamics*, 70, 771-786, DOI: 10.1007/s10236-020-01362-8.

Wyrtki K. (1989). Some thoughts about the West Pacific warm pool. In *Proceedings of the Western Pacific International Meeting and Workshop on TOGA COARE*, Nouméa, New Caledonia: 99-109 [online]. Available at: https://horizon.documentation.ird.fr/exl-doc/pleins_textes/doc34-08/30195.pdf#page=114 [Accessed 26 October 2022].

Yan X.H., Ho C.R., Zheng Q., and Klemas V. (1992). Temperature and size variabilities of the western Pacific warm pool. *Science*, 258 (5088), 1643-1645, DOI: 10.1126/science.258.5088.1643.

Yu L. and Weller R.A. (2007). Objectively Analyzed air-sea heat fluxes for the global ice-free oceans (1981-2005). *Bulletin of the American Meteorological Society*, 88(4), 527-540, DOI: 10.1175/BAMS-88-4-527.



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CENDERAWASIH HOT POOL: THE FREQUENT HIGH SEA SURFACE TEMPERATURE PHENOMENA AT CENDERAWASIH BAY, PAPUA

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ABSTRACT. The term “warm pool” refers to a body of water with the characteristic of SST exceeding 28°C within a particular area and a relatively long period in an annual circle. However, there are regions with an annual mean SST measured above 30°C, and we classified them as hot pools because of the conditions of intense solar radiation and low wind speed. One of the Hot Pool spots was found in Indonesia, in Cenderawasih Bay. The present study examines the existence of the Cenderawasih Hot Pool using long-term observation of satellite SST data. In order to learn more about their mechanisms, we also analyzed surface wind, surface heat flux, and surface current data. The results show that SSTs in Cenderawasih Bay have a 50% chance of exceeding 30°C within the 13 years of study (2013-2015). Heat input comes from strong solar radiation, i.e., 50% of solar radiation is more than 200 W/m². The location is also dominated by low wind speed, i.e., 80% wind speed of lower than 4 m/s, which caused the low latent loss in Cenderawasih Bay. Cenderawasih Bay is fully separated from surface currents during the dry and wet seasons since the easterly subsurface water flow does not enter the bay. The absence of strong currents prevents the mixing process, maintaining the high temperature in the surface layer. Those processes are discovered and they serve as compelling evidence to support Cenderawasih Bay as one of the Hot Pool areas within the Indonesian seas.

KEYWORDS: global climate; sea surface temperature; Hot Pool spot; Cenderawasih Bay

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INTRODUCTION

The western equatorial Pacific has a big impact on the climate. The Earth's circulation is specifically affected by the warm pool, an area with average sea surface temperatures (SST) exceeding 28°C (e.g., Wyrтки 1989; Yan et al. 1992; Clement and Seager 1999; Chongyin et al. 1999;

Pierrehumbert 2000; Clement et al. 2005; Thoron et al. 2005; Herweijer et al. 2005).

On the other hand, high SST in tropical regions has attracted researchers to investigate the mechanisms since the formation of high SST requires a particular atmospheric process (e.g., Ramanathan and Collins 1991; Wallace 1992; Arking and Ziskin 1994). This process is depicted by (Waliser

and Graham 1993), which shows the relation between SSTs and deep convection. They utilized highly reflective cloud information from an arbitrary examination of monthly SST data with a grid spacing of 2° (produced from combined satellite observation and in situ data) and daily visible and infrared satellite image data. The highly reflective cloud increases along with the SST as it rises from 26°C to 29.5°C . In contrast, the highly reflective cloud diminishes with increased SST in the temperature from 29.5°C to 32°C . As a result, the analysis proved that several atmospheric processes impacted SSTs below and above 29.5°C .

By taking advantage of high temporal and spatial resolution SST products derived from satellite observations (i.e., daily and $\leq 25 \text{ km} \times 25 \text{ km}$), several studies (Kawamura et al. 2008; Qin et al. 2007, 2008; Qin and Kawamura 2009, 2010; Wirasatriya et al. 2015, 2016, 2017a, 2020) were able to identify high SST events (i.e., more than 30°C) in specific areas and at certain periods and define them as Hot Event (HE). In summary, they concluded that considerable daily heat gains characterize the production of HE under high solar radiation and low wind speed brought on by "remote convection" mechanisms.

(Wirasatriya et al. 2015) elaborated on the climatology of HEs in the western equatorial Pacific using the SST dataset derived from satellite microwave sensors. Throughout nine years of observation (2003-2011), they discovered 71 HE cases in the western equatorial Pacific, with the majority centered on the Solomon Islands and New Guinea Island's northern coasts, which extend eastward up to 160°W . According to the climatology, the region has solar radiation of more than 200 W/m^2 and wind speeds of less than 4 m/s . Low wind speeds heavily influence the mechanism for HE incidence in the western equatorial Pacific. Much of the equatorial region experiences sun radiation above 200 W/m^2 during the HE periods. Low wind speeds minimize latent heat loss, which results in high SSTs and HEs in specific locations. (Wirasatriya et al. 2015) also emphasized that high solar radiation and low wind speed are much more common during the development stage and less common during the decay stage. This study also demonstrated that a rise in the long-term mean SST in the western equatorial Pacific is correlated with an increase in the frequency of HE events. HEs were responsible for 51.5% of the SSTs $>30^\circ\text{C}$ in the warm pool region bounded by the 29.5°C isotherms of

the climatological SST. Thus, statistically, there is a relation between the occurrence of HE and the formation of the western Pacific warm pool. Moreover, (Wirasatriya et al. 2020) demonstrated the role of HE in maintaining the warm mixed layer in the western Pacific warm pool. The frequent occurrence of HE transports heat from the surface layer to the deeper layer above the thermocline to maintain the warm mixed layer in the warm pool. Since the Pacific warm pool influences the global climate's variability through coupled ocean-atmosphere dynamics and thermodynamics (e.g., Clement and Seager 1999; Herweijer et al. 2005), this becomes an example of the contribution of high SST phenomena to regulate the global climate.

Within the Indonesian seas, the frequent appearance of high SST $> 30^\circ\text{C}$ has been reported by (Tita et al. 2020) in Tomini Bay and by (Swandiko et al. 2021) in the Malacca Strait. Their appearances also the high solar radiation and weak wind. The morphology of the semi-enclosed waters causes the high-frequency occurrence of low wind speed by less than 4 m/s in both areas. Besides these two areas, there is no more study on the high SST phenomena within the Indonesian Seas.

The present study demonstrates the frequent high SST occurrence (more than 30°C) in Cenderawasih Bay. Cenderawasih Bay is in northern Papua Island, part of the western Pacific warm pool. It is semi-enclosed water with a deep basin, surrounded by mountain chains (Fig. 1). Cenderawasih Bay is the habitat for whale sharks (Ihsan et al. 2018). Since the area determines warm pool definition with SST climatology higher than 28°C , we call the high-frequency occurrence of SSTs higher than 30°C in the Cenderawasih Bay the Cenderawasih Hot Pool. (Wirasatriya et al. 2015) found that the duration of HE occurrence in the western equatorial Pacific is no longer than two months since HE requires a typical condition of high solar radiation and low wind speed. Thus, it is interesting to understand the mechanisms of the high-frequency occurrence of high SST in the Cenderawasih Hot Pool.

MATERIALS AND METHODS

We used the semi-daily $11 \mu\text{m}$ SST products from Moderate Resolution Imaging Spectroradiometer (MODIS) Aqua Lv3 with a spatial resolution of $0.04^\circ \times 0.04^\circ$ (Esaias

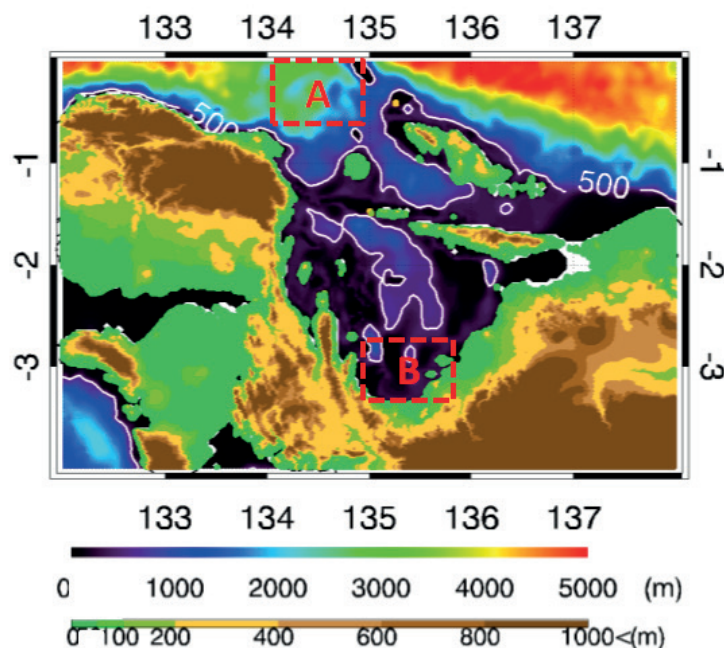


Fig. 1. Bathymetry and topography of Cenderawasih Bay. The dashed boxes A and B represent the sampling area for outside and inside Cenderawasih Bay, respectively, as shown in Fig. 5

et al. 1998) and observation period from 2003 to 2015. The MODIS SST 11 μm the Multi-Channel SST algorithm generates m by using brightness temperatures at 11 μm and 12 μm (Brown and Minnet 2009). The best accuracy was achieved by validating and testing this data against in-situ measurements (e.g., Lee et al. 2010; Qin et al. 2014; Ghanea et al. 2015).

We can obtain the Cross-Calibrated Multi-Platform (CCMP) gridded surface vector winds version 2.0 from the following website: www.remss.com, and use it to calculate surface winds. As a result, CCMP is regarded as a Level-3 ocean vector wind analysis product. They are created utilizing data from moored buoys, satellites, and model winds. The surface wind data for spatial and temporal resolutions is $0.25^\circ \times 0.25^\circ$ respectively and quarter daily. This wind product has greater accuracy than other wind reanalysis data.

We used a 0.5° grid from the Objectively Analyzed Air-Sea Fluxes project to analyze daily latent heat fluxes and solar radiation (Yu and Weller 2007). The period of observation for latent heat flux is from 2003 to 2015, but for solar radiation is only from 2003 to 2009 due to the availability of the data. Global 30 Arc-Second Elevation data was used to obtain topography. Global 30 Arc-Second Elevation (GTOPO30) data was used to obtain topography¹. ETOPO, a one-arc-minute global relief model of Earth's surface that combines land topography and ocean bathymetry², provides bathymetry. For oceanic parameters, we used the current reanalysis data from <http://marine.copernicus.eu/> to obtain the seasonal variation of the current pattern at 100 m depth in the Cenderawasih Hot Pool.

To calculate the frequency occurrence of SST within 13 years of observation, we calculate the percentage of SST > 30°C occurrence in each grid. Since high SST of more than 30°C occurs during the condition of low wind speed (i.e., < 4 m/s) and high solar radiation (i.e., > 200 W/m²) (Wirasatriya et al. 2015), we calculated the percentage of low wind speed and high solar radiation in the Cenderawasih Bay. Furthermore, we also calculated the percentage of low latent heat release (<120 W/m²) to explain how wind speed influences the variability of SST. The equation for calculating the percentage is as follows:

$$\% (x,y) = \frac{1}{n} \sum_{i=1}^n pi(x,y) \times 100\% \quad (1)$$

where $\%(x,y)$ is percentage the percentage of high SST > 30°C or percentage of weak wind speed < 4 m/s or percentage of solar radiation > 200 W/m² or percentage of latent heat flux < 120W/m² at position (x,y); pi is the amount of SST data > 30°C or the amount of weak wind speed < 4 m/sec or the amount solar radiation > 200 W/m² or the amount of latent heat flux < 120W/m² at position (x,y). Furthermore, we also analyze the climatological mean of each parameter. To create monthly and monthly climatology, all geophysical parameters are incorporated as follows (Wirasatriya et al. 2017b):

$$\bar{X}(x,y) = \frac{1}{n} \sum_{i=1}^n xi(x,y,t) \quad (2)$$

where \bar{X} is monthly mean value or monthly climatology value at position (x,y), xi is i^{th} value of the data at (x,y) position and time t . Next, n is the number of data in 1 month and the number of monthly data in 1 period of climatology (i.e., from 2003 to 2016 = 13 data) for monthly calculation and monthly climatology calculation, respectively. Moreover, if pixel xi is a hollow pixel, it is not included in the calculation.

RESULTS AND DISCUSSION

Cenderawasih Hot Pool

The definition of the Cenderawasih Hot Pool is described in Fig. 2, which shows the climatological mean of SST and the percentage of daily SSTs more than 30°C from 2003 to 2015 (13 years). In Cenderawasih Bay, the mean SST at the southern part of 2°S is more than 30°C, and the high SSTs are more than 50%. It means that the mean SST in Cenderawasih Bay is higher than the definition of the western Pacific warm pool which is 28°C (Wyrcki 1989), and more than 3.5 years within 2003-2015, SSTs in Cenderawasih Bay can reach more than 30°C. At the southernmost of the bay, the mean SST reaches 30.5°C, with the percentage of high SSTs being more than 70%. It has become the hottest area of Cenderawasih Bay.

In contrast, in the area outside Cenderawasih Bay, i.e., the western Pacific warm pool's offshore seas, the percentage is less than 30%. Thus, this evidence supports the definition of Cenderawasih Bay as the Cenderawasih Hot Pool. The high-frequency occurrence of high SST in Cenderawasih Bay may correspond to whale sharks' yearly appearance. (Ihsan et al. 2018) indicated the tolerance of whale sharks in high SST conditions.

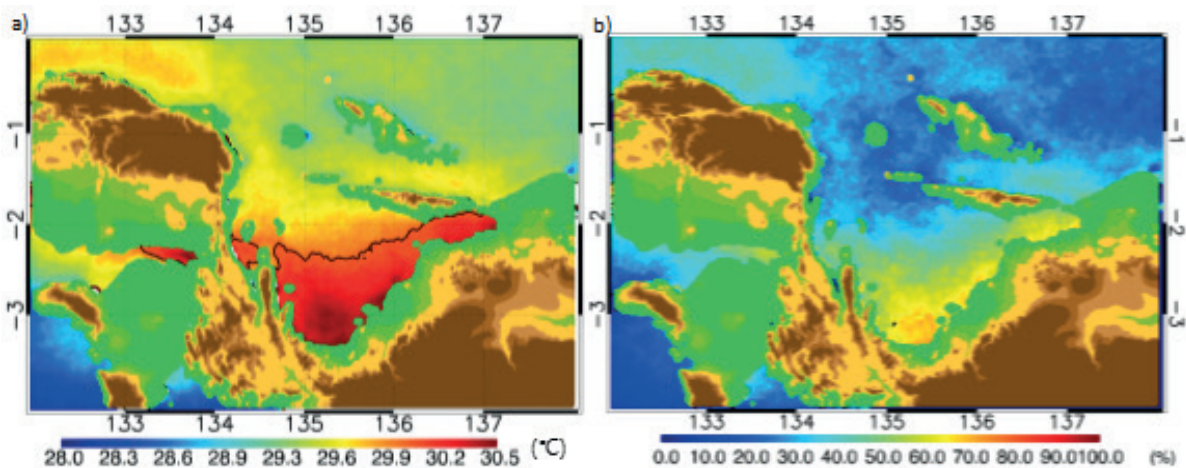


Fig. 2. Climatological mean of SST (a) and percentage of daily SSTs > 30°C during thirteen years of observation (2003-2015) (b). The black contour is 30°C

¹https://www.usgs.gov/centers/eros/science/usgs-eros-archive-digital-elevation-global-30-arc-second-elevation-gtopo30?qt-science_center_objects=0#qt-science_center_objects

² <https://www.ngdc.noaa.gov/mgg/global/>

Atmospheric Aspect of Cenderawasih Hot Pool

Fig. 3 clearly shows that low wind speed dominates Cenderawasih Bay, i.e., 80% wind speed < 4 m/s occurred during 2003-2015. This causes low latent loss in Cenderawasih Bay. For solar radiation, the percentage of high solar radiation inside and outside Cenderawasih Bay is almost similar i.e., nearly 50%, indicating that high solar radiation is not dominant in both areas. This means that, although high solar radiation does not frequently occur, the absence of strong wind speed may maintain the latent heat loss to keep the SST in Cenderawasih Bay higher than 30°C.

To investigate how low wind speed occurs in the study area, we show the monthly climatology map of surface wind during summer and winter (Fig. 4). For both seasons, Cenderawasih Bay is protected from high wind speed due to high mountain chains in the western, southern, and eastern parts. In the northern part, small islands at the mouth of the bay prevent the strong wind speed from entering the bay. Thus, the role of topography is crucial for the occurrence of constant high SSTs in Cenderawasih Bay. The same tendencies are also found at the Tomini Bay and Malacca Strait, as indicated by (Tita et al. 2020) and (Swandiko et al. 2021).

The influence of wind speed and solar radiation on the SST variation is further examined by plotting the monthly climatology of wind speed, solar radiation, and SST at areas A and B (Fig. 1), representing the areas outside and inside Cenderawasih Bay. Seasonal variations of SSTs are observed in both areas. The minimum SSTs occur during the wet and dry seasons, while the maximum SSTs during the transition season. This seasonal SST variability is similar to the other areas in the Indonesian seas, such as the Java Sea (Wirasatriya et al. 2018), Maluku Sea (Wirasatriya et al. 2019), Halmahera Sea (Setiawan et al. 2019), etc.

For the area outside Cenderawasih Bay, SST ranges from 28.2°C to 30.6°C. It is seen that the variability of SST is controlled by wind speed and solar radiation. From January to June, when wind speed decreases from 2.2 m/s to 0.3 m/s and solar radiation is more than 200 W/m², SST increases from 28.2°C to 30.5°C. From June to October, wind speed is lower than 1 m/s. In the absence of strong wind, the low SST in July is caused by the decrease of solar radiation to a minimum.

For the area inside Cenderawasih Bay, persistent high SST is observed as the SST ranges from 30.2°C to 31.1°C. It may be due to the low wind speed of fewer than 1 m/s for years. Thus, the variability of SST is seen as ruled by solar radiation. The minimum solar radiation causes the minimum SST in July. However, the absence of strong wind prevents latent heat loss that maintains the high SST inside Cenderawasih Bay.

The oceanic aspect of Cenderawasih Hot Pool

To investigate the oceanic aspect of the Cenderawasih Hot Pool, we plotted the current patterns at 100 m depth shown in Fig. 6. A northwestward current has been identified north of Cenderawasih Bay. During the wet season (January), the speed of this current ranges from 0.5 m/s to 0.7 m/s. During the dry season (August), when the northwesterly wind disappears (Fig. 4b), the speed of the northwestward current increases. This northwestward current is known as the New Guinea Coastal Under Current (NGCUC). NGCUC is a permanent current feature at 200 m depth regardless of the wind reversals (Tsuchiya et al. 1989). Inside Cenderawasih Bay, strong currents are absent in this area, preventing the mixing process from maintaining the high temperature in the surface layer. During the dry and wet seasons, Cenderawasih Bay is fully isolated since the westward subsurface water flow does not enter Cenderawasih Bay. Thus, this isolated basin causes the Cenderawasih Bay area to be influenced mainly by the air-sea interaction process mentioned in the previous section creating Cenderawasih Hot Pool.

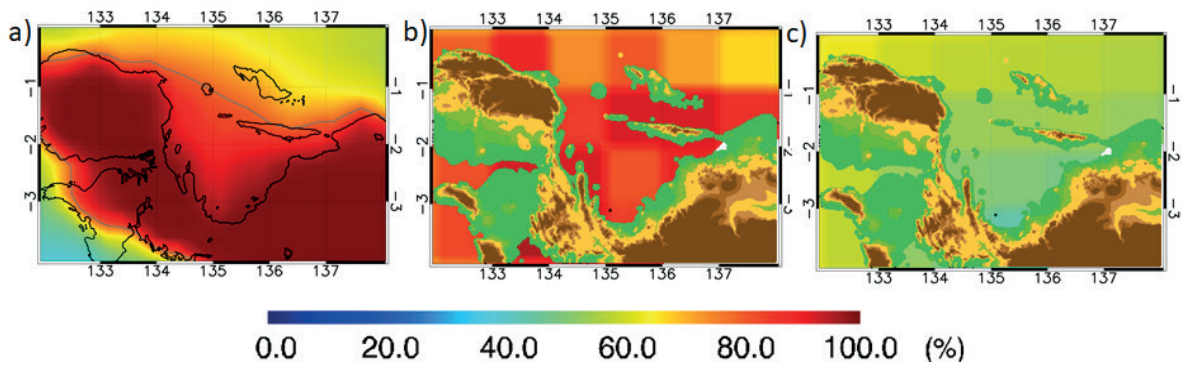


Fig. 3. The percentage of wind speed less than four m/s (a), latent heat release less than 120 W/m² (b), and solar radiation more than 200 W/m² (c)

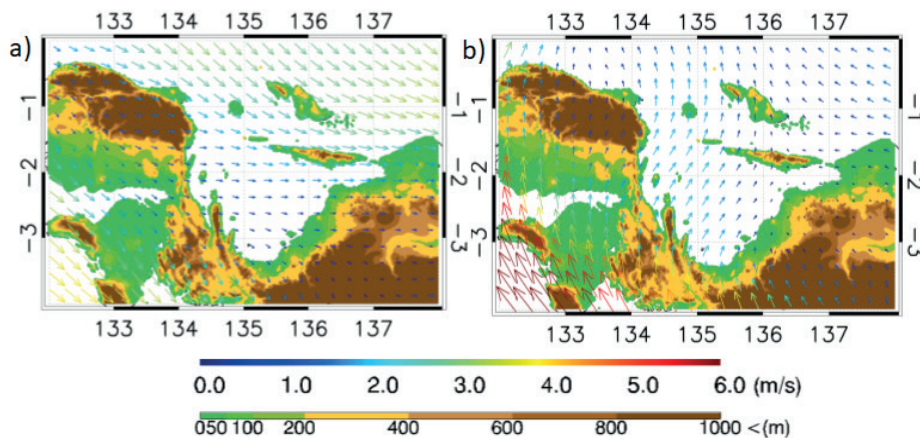


Fig. 4. Monthly climatology of surface wind speed during a) wet season (January) and b) dry season (August)

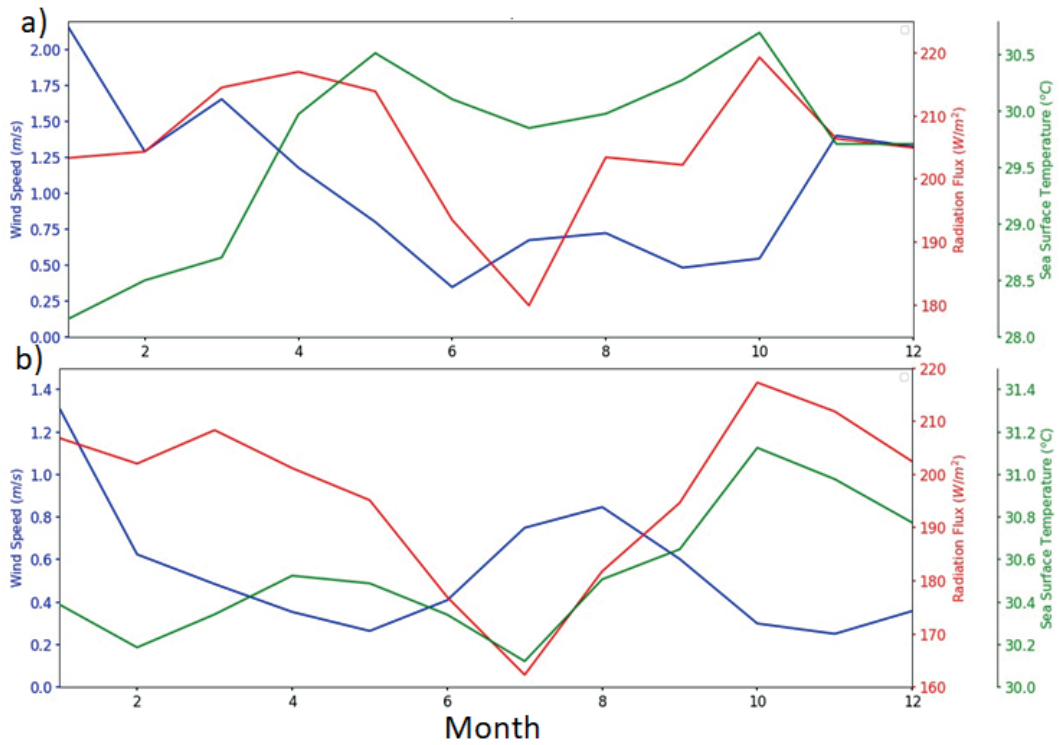


Fig. 5. Monthly climatology of SST, wind speed, and solar radiation at a) area outside and b) area inside Cenderawasih Bay as shown in Fig. 1

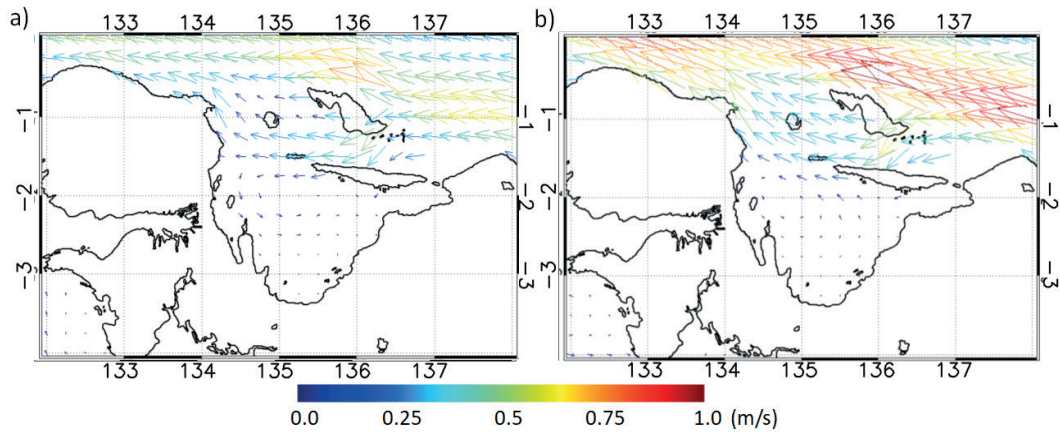


Fig. 6. Monthly climatology of current at 100 m depth in January (a) and August (b)

CONCLUSIONS

The hot Pool is used to categorize high SST episodes (>30°C) in particular areas and during specific periods (relatively long periods). Since the term “warm pool” has been used to define an area with the annual average SST distributed below 30°C in dominant, therefore “hot pool” was taken to represent the area with SST dominated above 30°C.

Under high solar radiation and low wind speed, this event is characterized by considerable daily heat gains.

The constant high SST occurrence in Cenderawasih Bay, which is more than 50% of 13 years, SSTs can reach more than 30°C and are dominated by the condition of low wind speed, i.e., 80% wind speed is less than m/s along the years, also 50% solar radiation measured is more than 200 W/m². The current pattern at 100 m depth also shows that NGCUC does not enter Cenderawasih Bay. Thus, this indicates that Cenderawasih Bay is an isolated water. Those are the solid evidence for defining Cenderawasih Bay as the “Cenderawasih Hot Pool.”

REFERENCES

- Arking A. and Ziskin D. (1994). Relationship between clouds and sea surface temperatures in the western tropical Pacific. *Journal of Climate*, 7(6), 988-1000, DOI: 10.1175/1520-0442(1994)007%3C0988:RBCASS%3E2.0.CO;2.
- Brown O.B. and Minnett P.J. (2009). MODIS infrared sea surface temperature algorithm theoretical basis document, ver 2.0. [online]. Available at: https://modis.gsfc.nasa.gov/data/atbd/atbd_mod25.pdf [Accessed 25 October 2022].
- Chongyin L., Mingquan M., and Guangqing Z. (1999). The variation of warm pool in the equatorial western Pacific and its impacts on climate. *Advances in Atmospheric Sciences*, 16(3), 378-394, DOI: 10.1007/s00376-999-0017-0.
- Clement A. and Seager R. (1999). Climate and the tropical oceans. *Journal of Climate*, 12(12), 3383-3401, DOI: 10.1175/1520-0442(1999)012%3C3383:CATTO%3E2.0.CO;2
- Clement A.C., Seager R., and Murtugudde R. (2005). Why are there tropical warm pools? *Journal of Climate*, 18(24), 5294-5311, DOI: 10.1175/JCLI3582.1.
- Esaías W.E., Abbott M.R., Barton I., Brown O.B., Campbell J.W., Carder K.L., ... and Minnett P.J. (1998). An overview of MODIS capabilities for ocean science observations, *IEEE Transactions on Geoscience and Remote Sensing*, 36(4), 1250-1265, DOI: 10.1109/36.701076.
- Ghanea M., Moradi M., Kabiri K., and Mehdinia A. (2015). Investigation and validation of MODIS SST in the northern Persian Gulf, *Advances in Space Research*, 57(1), 127-136, DOI: 10.1016/j.asr.2015.10.040.
- Harweijer C., Seager R., Winton M., and Clement A.M.Y. (2005). Why ocean heat transport warms the global mean climate. *Tellus*, 57(4), 662-675, DOI: 10.3402/tellusa.v57i4.14708.
- Hosoda K. (2013). Empirical method of diurnal correction for estimating sea surface temperature at dawn and noon. *Journal of Oceanography*, 69, 631-646, DOI: 10.1007/s10872-013-0194-4.
- Ihsan E.N., Enita S.Y., and Wirasatriya, A. (2018). Oceanographic factors in fishing ground location of Anchovy at Teluk Cenderawasih National Park, West Papua: Are these factors have an effect of Whale Sharks appearance frequencies? In *IOP Conference Series: Earth and Environmental Science*, 116(1), 012-017, DOI: 10.1088/1755-1315/116/1/012017.
- Kawamura H., Qin H., and Ando, K. (2008). In-situ diurnal sea surface temperature variations and near-surface thermal structure in the tropical hot event of the Indo-Pacific warm pool. *Journal of Oceanography*, 64, 847-857, DOI: 10.1007/s10872-008-0070-9.
- Ming-An L., Tzeng M. T., Hosoda K., Sakaida F., Kawamura H., Shieh W. J., Yang Y., and Chang, Y. (2010). Validation of JAXA/MODIS sea surface temperature in water around Taiwan using the Terra and Aqua satellites. *TAO: Terrestrial, Atmospheric and Oceanic Sciences*, 21(4), 727-736, DOI: 10.3319/TAO.2009.09.07.01(Oc).
- Pierrehumbert R.T. (2000). Climate change and the tropical Pacific: The sleeping dragon wakes. *Proceedings of the National Academy of Sciences*, 97(4), 1355-1358, DOI: 10.1073/pnas.97.4.1355.
- Qin H., Kawamura H., and Kawai Y. (2007). Detection of Hot Event in the equatorial Indo-Pacific warm pool using advanced satellite sea surface temperature, solar radiation, and wind speed. *Journal of Geophysical Research: Oceans*, 112 (C7), DOI: 10.1029/2006JC003969.
- Qin H., Kawamura H., Sakaida F., and Ando K. (2008). A case study of the tropical hot event in November 2006 (HE0611) using a geostationary meteorological satellite and the TAO/TRITON mooring array. *Journal of Geophysical Research: Oceans*, 113(C8), DOI: 10.1029/2007JC004640.
- Qin H. and Kawamura H. (2009). Atmosphere response to a hot SST event in November 2006 as observed by AIRS instrument. *Advances in space research*, 44(3), 395-400, DOI: 10.1016/j.asr.2009.03.003.
- Qin H. and Kawamura H. (2010). Air-sea interaction throughout the troposphere over a very high sea surface temperature. *Geophysical research letters*, 37(1), 1-4, DOI: 10.1029/2009GL041685.
- Qin H., Chen G., Wang W., Wang D., and Zeng L. (2014). Validation and application of MODIS-derived SST in the South China Sea. *International journal of remote sensing*, 35(11-12), 4315-4328, DOI: 10.1080/01431161.2014.916439.
- Ramanathan V. and Collins W. (1991). Thermodynamic regulation of ocean warming by cirrus clouds deduced from observations of the 1987 El Niño. *Nature* 351 (6321), 27-32, DOI: 10.1038/351027a0.
- Setiawan R.Y., Wirasatriya A., Hernawan U., Leung S., and Iskandar I. (2020). Spatio-temporal variability of surface chlorophyll-a in the Halmahera Sea and its relation to ENSO and the Indian Ocean Dipole. *International Journal of Remote Sensing*, 41(1), 284-299, DOI: 10.1080/01431161.2019.1641244.
- Swandiko M., Wirasatriya A., Marwoto J., Muslim, Indrayanti E., Subardjo P., Ismunarti D.H. (2021). Studi persistensi suhu permukaan laut tinggi (>30°C) di perairan Selat Malaka. *Buletin Oseanografi Marina*, 10(2), 162-170. [online]. Available at: <https://ejournal.undip.ac.id/index.php/buloma/article/download/31554/19352> [Accessed 25 October 2022] (in Indonesian).
- De Garidel-Thoron T., Rosenthal Y., Bassinot F., and Beaufort, L. (2005). Stable sea surface temperatures in the western Pacific warm pool over the past 1.75 million years. *Nature*, 433(7023), 294-298, DOI: 10.1038/nature03189.
- Tita A.D.C., Wirasatriya A., Sugianto D.N., Maslukah L., Handoyo G., Helmi M., and Avianto P. (2020). Persistence of high sea surface temperature (> 30°C) in Tomini Bay. In *IOP Conference Series: Earth and Environmental Science*, 530(1), 012-038, DOI: 10.1088/1755-1315/530/1/012038.
- Waliser D.E. and Graham N.E. (1993). Convective cloud systems and warm-pool sea-surface temperatures: Coupled interactions and self-regulation. *Journal of Geophysical Research: Atmospheres*, 98(D7), 12881-12893, DOI: 10.1029/93JD00872.
- Wallace J.M. (1992). Effect of deep convection on the regulation of tropical sea surface temperature. *Nature*, 357(6375), 230-231, DOI: 10.1038/357230a0.
- Wirasatriya A., Kawamura H., Shimada T., and Hosoda K. (2015). Climatology of hot events in the western equatorial Pacific. *Journal of Oceanography*, 71, 77-90, DOI: 10.1007/s10872-014-0263-3.
- Wirasatriya A., Kawamura H., Shimada T., Hosoda K. (2016). Atmospheric structure favoring high sea surface temperatures in the western equatorial Pacific. *Journal of Geophysical Research*, 121(19), 11-368, DOI: 10.1002/2016JD025268.
- Wirasatriya A., Sugianto D.N., and Helmi M. (2017). The Influence of Madden Julian Oscillation on the Formation of the hot event in the western equatorial Pacific. In *IOP Conference Series: Earth and Environmental Science*, 55(1), 012-006, DOI: 10.1088/1755-1315/55/1/012006.
- Wirasatriya A., Setiawan R.Y., and Subardjo P. (2017). The effect of ENSO on the variability of chlorophyll-a and sea surface temperature in the Maluku Sea, *IEEE Journal of Selected Topics in Applied Earth Observations and Remote Sensing*, 10(12), 5513-5518, DOI: 10.1109/JSTARS.2017.2745207.
- Wirasatriya A., Prasetyawan I.B., Triyono C.D., Muslim, and Maslukah L. (2018). Effect of ENSO on the variability of SST and chlorophyll-a in Java Sea. In *IOP Conference Series: Earth and Environmental Science*, 116(1), 012-063, DOI: 10.1088/1755-1315/116/1/012063.
- Wirasatriya A., Sugianto D.N., Helmi M., Setiawan R.Y., and Koch M. (2019). Distinct characteristics of SST variabilities in the Sulawesi Sea and the northern part of the Maluku Sea during the southeast monsoon. *IEEE Journal of Selected Topics in Applied Earth Observations and Remote Sensing*, 12(6), 1763-1770, DOI: 10.1109/JSTARS.2019.2913739.

Wirasatriya A., Kawamura H., Helmi M., Sugianto D.N., Shimada T., Hosoda K., Handoyo G., Putra Y.D.G., and Koch M. (2020). Thermal structure of Hot Events and their possible role in maintaining the warm isothermal layer in the Western Pacific warm pool. *Ocean Dynamics*, 70, 771–786, DOI: 10.1007/s10236-020-01362-8.

Wyrski K. (1989). Some thoughts about the West Pacific warm pool. In *Proceedings of the Western Pacific International Meeting and Workshop on TOGA COARE*, Noumèa, New Caledonia: 99-109 [online]. Available at: https://horizon.documentation.ird.fr/exl-doc/pleins_textes/doc34-08/30195.pdf#page=114 [Accessed 26 October 2022].

Yan X.H., Ho C.R., Zheng Q., and Klemas V. (1992). Temperature and size variabilities of the western Pacific warm pool. *Science*, 258 (5088), 1643-1645, DOI: 10.1126/science.258.5088.1643.

Yu L. and Weller R.A. (2007). Objectively Analyzed air-sea heat fluxes for the global ice-free oceans (1981–2005). *Bulletin of the American Meteorological Society*, 88(4), 527-540, DOI: 10.1175/BAMS-88-4-527.